

Physical–biological coupling in the coastal upwelling system of the Ría de Vigo (NW Spain). I: *In situ* approach

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ABSTRACT: The fate of the inorganic and organic nitrogen trapped in the coastal upwelling system of Ría de Vigo (NW Spain) was studied at the 2 to 4 d time-scale during July and September 2002. A transient geochemical model was applied to the measured residual currents and concentrations of inorganic nitrogen (N_T), and dissolved (DON) and particulate (PON) organic N to obtain (1) the net balance of inputs minus outputs ($i - o$); (2) the net accumulation; and (3) the net ecosystem production (NEP) of N_T , DON and PON. Previously unaccounted lateral variations in residual currents and N species concentrations in the ría were considered. The ría was autotrophic during July (average NEP, $107 \text{ mg N m}^{-2} \text{ d}^{-1}$). About 25 % of this material was exported to the shelf and the remaining 75 % was transferred to the sediments or promoted to higher trophic levels. During summer upwelling episodes, 30 to 70 % of the organic N exported to the shelf came from materials previously accumulated in the ría. By contrast, during summer relaxation, 60 to 70 % of the accumulated organic N came from *in situ* conversion of N_T and the remaining 30 to 40 % was allochthonous. As shown by other authors, NEP was rather low during intense upwelling and increased to high values during its subsequent relaxation. In September, the metabolism changed from heterotrophic to slightly autotrophic (average NEP, $26 \text{ mg N m}^{-2} \text{ d}^{-1}$). The DON and PON imported from the shelf during autumn downwelling experienced a different fate: DON was consumed and PON accumulated in the ría. By contrast, DON and PON were produced *in situ* at the expense of N nutrients previously accumulated into the system during autumn relaxation. The roles of vertical convection and turbulent mixing in the fertilization of the photic layer were also assessed. Mixing was the most important fertilization mechanism, ensuring transport of nutrients under upwelling and downwelling conditions.

KEY WORDS: Nitrogen fluxes · Geochemical model · Upwelling · Downwelling · Rías Baixas (NW Spain)

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INTRODUCTION

New production (NP; see Table 1 for list of abbreviations) is the fraction of primary production (P_g) that is supported by allochthonous nutrient salts (Dugdale & Goering 1967) and is a key variable for understanding marine ecosystems functioning (e.g. exploitation of marine resources global climate change). Therefore,

NP not only represents the threshold for biomass exploitation (Quiñones & Platt 1991), but also affects the efficiency of the biological pump as a trap for anthropogenic CO_2 (Longhurst 1991).

The most straightforward way to estimate NP is to consider the nutrient balance (input minus output) in a given ecosystem (Smith & Hollibaugh 1993, Wollast 1993). However, this conceptually simple method

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Table 1. Glossary of relevant terms

A	Free surface area of the ría
A_S, A_B	Surface and bottom areas of the transverse section
ADCP	Acoustic Doppler current profiler
chl a	Chlorophyll a
C_S, C_B	Average surface and bottom concentrations of C
DON	Dissolved organic nitrogen
DOM	Dissolved organic matter
ENACW	Eastern North Atlantic Central Water
F_x	Flux of any species carried by a horizontal convective flow (Q_S, Q_B)
F_z	Flux of any species carried by vertical convective flow (Q_z)
K_z	Turbulent mixing coefficient
LNM	Level of no movement
$\langle N \rangle$	Brunt-Väisälä frequency
N_T	Dissolved inorganic nitrogen
NEP	Net ecosystem production
NEP_T	Net ecosystem production of N_T
NEP_D	Net ecosystem production of DON
NEP_P	Net ecosystem production of PON
NEP_{SED}	Net balance of PON sedimentation, PON and resuspension and N_T -diffusion from the sediment
M_z	Vertical diffusive fluxes
$o - i$	Net budget of outputs minus inputs
$(o - i)_T$	Net budget of outputs minus inputs of N_T
$(o - i)_P$	Net budget of outputs minus inputs of PON
$(o - i)_D$	Net budget of outputs minus inputs of DON
PON	Particulate organic nitrogen
PAR	Photosynthetic available radiation
Q_B	Bottom horizontal flow
Q_S	Surface horizontal flow
Q_z	Vertical flow
R_i	Richardson number
ρ	Density
U_0	Tidal velocity amplitude
V	Volume
V_x	Horizontal convective velocity
V_z	Vertical convective velocity
VMADCP	Vessel-mounted acoustic Doppler current profiler
$\Delta(V \cdot N_T)/\Delta t$	Accumulation of N_T
$\Delta(V \cdot DON)/\Delta t$	Accumulation of DON
$\Delta(V \cdot PON)/\Delta t$	Accumulation of PON

involves operational difficulties: velocity and turbulent diffusion profiles, together with nutrient concentrations, have to be determined with a frequency higher than the flushing time of the study system. For this reason, different indirect methods have been developed over the last few decades to evaluate NP.

Gravitational sinking and subsequent mineralization is the main fate of NP in vast areas of the open ocean, where the horizontal circulation is quite slow and most of *Pg* is processed within the microbial loop (Legendre & Rassoulzadegan 1995, Cotner & Biddanda 2002). Therefore, assuming steady-state conditions, NP

should be equal to the sedimentation rate measured at the base of the photic layer (Eppley & Peterson 1979). *In vitro* approaches have also been used to estimate N ; these include the ^{15}N -labeled nitrate/ammonium+urea method (Dugdale & Goering 1967) or the light-dark oxygen incubation method (Strickland & Parsons 1972), in which it is assumed that the net community production (NCP) estimated in the light bottle is equal to the NP (Quiñones & Platt 1991). However, the application of the latter method in subtropical gyres has led to controversial results (autotrophic NCP > 0 or heterotrophic NCP < 0; Williams 1998, Serret et al. 2001, del Giorgio & Duarte 2002). In any case, the validity of shallow sediment traps (Buesseler 1991, Ducklow et al. 1995), ^{15}N -labeled incubations (Legendre & Gosselin 1989, Wafar et al. 1995), and light-dark oxygen incubations (Hansell et al. 2004) to estimate NP has been openly challenged.

The evaluation of NP is more difficult in coastal systems, due to the enhanced horizontal advection, and the transference of a significant fraction of *Pg* to metazoans (Legendre & Rassoulzadegan 1995, Cotner & Biddanda 2002). In the particular case of coastal upwelling systems at temperate latitudes, the short flushing times and the non-steady-state conditions increase the complexity of NP estimation (Álvarez-Salgado et al. 1996a, Smith & Hollibaugh 1997). The mass balance approach using sediment traps and/or *in vitro* incubations may not reflect an accurate NP estimate for these ecosystems (Smith & Hollibaugh 1997, Álvarez-Salgado et al. 2001).

In this study and its companion paper (Piedracoba et al. 2008 this volume) we estimate the NP in a coastal upwelling system (Ría de Vigo, NW Spain) under contrasting upwelling (July) and downwelling (September) conditions using mass balance, sediment trap, and *in vitro* incubation approaches. Part I is focused on the estimation of the inputs minus outputs balance of nitrogen species (total nutrients, N_T ; dissolved organic nitrogen, DON; particulate organic nitrogen, PON) on a short time-scale (<4 d) by combining dynamic and chemical determinations along a transverse section in the central segment of the Ría de Vigo. A mass balance box model was developed to evaluate the relative importance of net accumulation and net ecosystem production ($NEP = NP$) as a mechanism of organic and inorganic nitrogen trapping in this ecosystem. Previous geochemical balances applied in the Ría de Vigo (Álvarez-Salgado et al. 2000, Gilcoto et al. 2001) estimated the velocity field from salinity and temperature distributions. In the present study, NP was estimated from direct current measurements. In addition, the lateral variability of these currents was considered for the first time. In Part II, the metabolism of the Ría de Vigo is assessed while considering the oxygen production/

respiration balance. Sedimentation and microzooplankton grazing rates, estimated for the first time in the Ría de Vigo, are also considered. The NP estimations obtained by the *in situ* (Part I) and *in vitro* (Part II) approaches are compared and the possible causes of the observed discrepancies are discussed.

MATERIALS AND METHODS

Survey area. The Ría de Vigo is a large (2.76 km^3) V-shaped indentation in the western coast of the Iberian Peninsula (Fig. 1). It experiences wind-driven upwelling of cold, salty and nutrient-rich Eastern North Atlantic Central Water (ENACW) from April–May to September–October (Wooster et al. 1976, Fraga 1981). Upwelling events occur with a 1 to 2 wk frequency (Blanton et al. 1987, Álvarez-Salgado et al. 1993). Downwelling is the dominant process the rest of the year; it is characterized by the advection of warm and nutrient-poor shelf surface water to the ría (Castro et al. 1997). The main tributary is the river Oitabén–

Verdugo, which drains an average annual flow of $15 \text{ m}^3 \text{ s}^{-1}$ (Nogueira et al. 1997) into San Simón Bay, a semi-enclosed basin of $\sim 20 \text{ km}^2$ (Fig. 1).

Sampling strategy. Five stations in the middle segment of the Ría de Vigo ($41 \pm 2 \text{ m}$ max. depth) were sampled about 1 h before sunrise on 15, 18, 22 and 26 July and 17, 19, 23 and 26 September 2002. Water samples from Stn 00 (Fig. 1) were taken with a rosette sampler equipped with twelve 10 l PVC Niskin bottles with stainless-steel internal springs, while water samples from Stns R01 and R03 located at the northern and southern ends of the segment, were taken with 5 l PVC Niskin bottles. At Stn 00, salinity and temperature were recorded with an SBE 9/11 CTD probe attached to the rosette sampler. Conductivity measurements were converted into practical salinity scale values with the equation of UNESCO (1985). Water samples were collected from 5 depths: at the surface or at the depth of 50% of the photosynthetically available radiation (PAR, $\mu\text{E m}^{-2} \text{ s}^{-1}$) penetration ($2.5 \pm 0.3 \text{ m}$), at 25% of PAR penetration ($7 \pm 1 \text{ m}$), at 1% of PAR penetration ($14.2 \pm 2.2 \text{ m}$), at $26.2 \pm 1.2 \text{ m}$, and at the bottom.

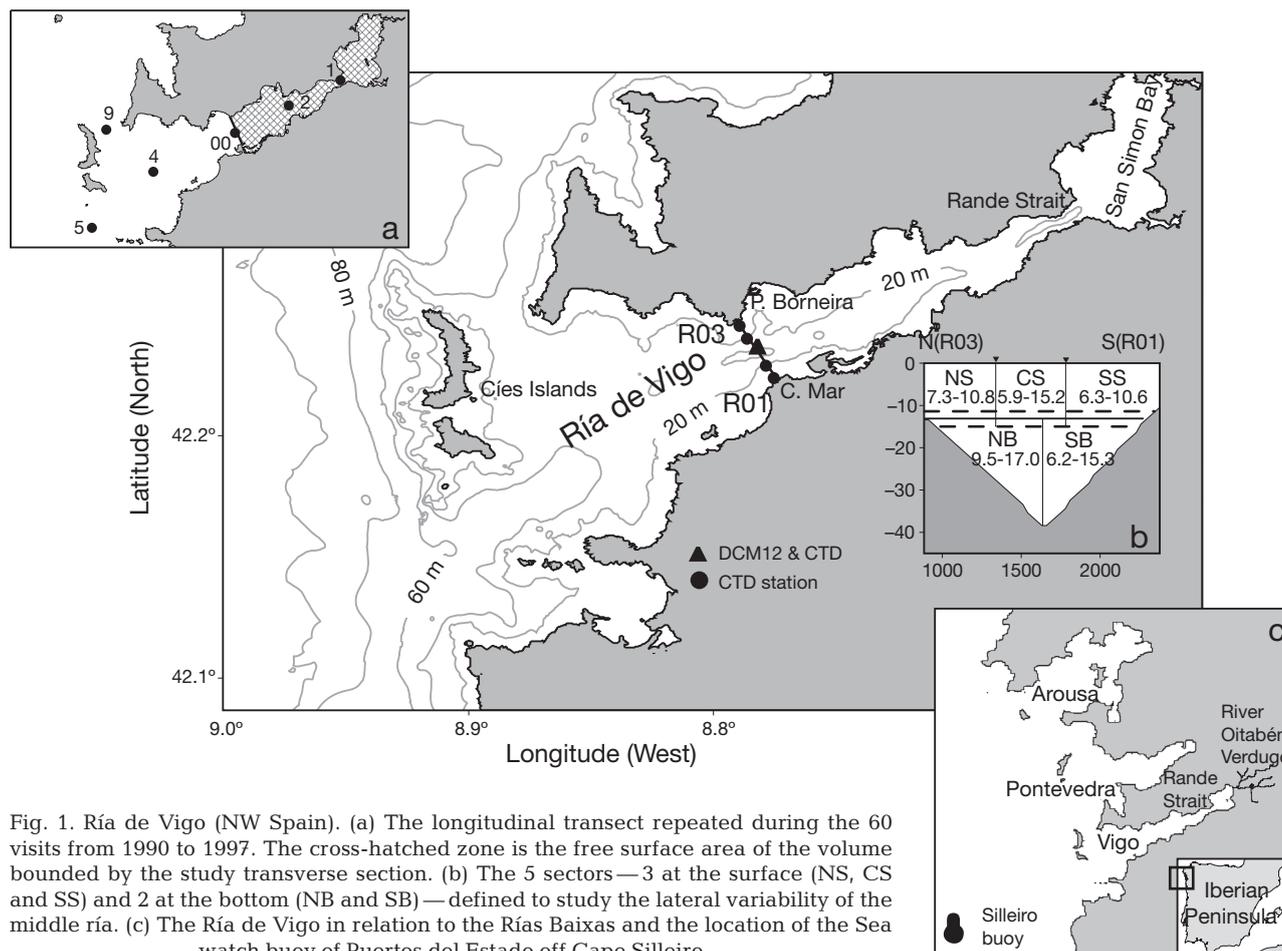


Fig. 1. Ría de Vigo (NW Spain). (a) The longitudinal transect repeated during the 60 visits from 1990 to 1997. The cross-hatched zone is the free surface area of the volume bounded by the study transverse section. (b) The 5 sectors—3 at the surface (NS, CS and SS) and 2 at the bottom (NB and SB)—defined to study the lateral variability of the middle ría. (c) The Ría de Vigo in relation to the Rías Baixas and the location of the Sea watch buoy of Puertos del Estado off Cape Silleiro

At Stns R01 and R03, salinity and temperature were recorded with a CTD SBE-25 and water samples were collected from 2 depths: surface and bottom.

Wind and current measurements. Coastal wind data were obtained from the Seawatch buoy off Cape Silleiro (www.puertos.es; our Fig. 1), a representative site for winds blowing off the Ría de Vigo (Herrera et al. 2005). Low-pass filtered winds are used to define the prevailing oceanographic scenario in the ría during each period. The heat balance and freshwater flux for each sampling date were taken from Piedracoba et al. (2005).

Current velocity profiles were measured aboard RV 'Mytilus' with an RDI vessel-mounted broadband acoustic Doppler current profiler (VMADCP) of 300 kHz along the transverse section joining Stns R01 and R03 on 18 and 22 July and 19 September 2002. In addition, current profiles were measured with an Aandera DCM12 acoustic Doppler current profiler (ADCP) moored at Stn 00. The mooring was deployed from 10 to 29 July and from 17 to 26 September 2002. The DCM12 lay on the seabed, recording, at 30 min intervals, the current integrated in 5 vertical cells 11 m thick. DCM12 and VMADCP data were rotated counterclockwise 30° to produce a longitudinal and a transversal component relative to the main axis of the ría.

A C++ version of the HAMSOM model (3D, *z*-coordinate Arakawa-C grid), baroclinic, semi-implicit finite-difference numerical prediction model), recently validated in the outer (Souto et al. 2003) and inner (Diz et al. 2004) Ría de Vigo, was used to reproduce the lateral variability of the subtidal current field on 15 and 26 July and 17, 23 and 26 September 2002. The accuracy of the model to reproduce the lateral current field data has been successfully tested by Piedracoba et al. (unpubl.).

Horizontal advection. Five sectors—3 at the surface (NS, CS and SS) and 2 at the bottom (NB and SB)—were defined to account for the lateral variability in the middle ría (Fig. 1b). The criterion to divide the surface and bottom sectors was based on the level of no horizontal motion, taking into account the data from the DCM12 and the VMADCP, and the outputs of the HAMSOM model. Horizontal advection velocities were averaged in the 5 sectors.

Vertical advection. The Ría de Vigo was visited for the purpose of measuring mass balances 60 times from 1990 to 1997. Fig. 1a shows the longitudinal transect repeated during the 60 visits. The good relationship between the bottom horizontal flow (Q_B) and the vertical flow (Q_Z) obtained from the mass balances (Fig. 1b) permits calculation of the unknown Q_Z (in $\text{m}^3 \text{s}^{-1}$) from the measured Q_B (in $\text{m}^3 \text{s}^{-1}$) in this study (n are the pairs of consecutive sampling):

$$Q_Z = -212(\pm 43) + 0.73(\pm 0.03) \cdot Q_B \quad (1)$$

$$r = +0.92, p < 0.001, n = 47$$

The vertical convective velocity (V_Z) was calculated considering the surface area of the shaded zone in the inset of Fig. 1a. The vertical displacement of the level of no movement (LNM) between 2 consecutive surveys was also calculated.

Vertical mixing. The turbulent mixing coefficient, K_Z ($\text{m}^2 \text{d}^{-1}$), was parameterized according to Munk & Anderson (1948), using the ADCP and CTD data:

$$K_Z = K_0 \cdot (1 + 10 \cdot R_i)^{1/2} \quad (2)$$

$$K_0 = 3 \times 10^{-3} \cdot U_0 \cdot Z \quad (3)$$

$$R_i = -\frac{g \cdot \frac{\partial \rho}{\partial z}}{\left(\frac{\partial u}{\partial z}\right)^2} \quad (4)$$

where K_0 is the eddy coefficient for neutral stability, U_0 is the amplitude of the tidal velocity from 0 to Z m depth, R_i is the Richardson number, ρ is the average density, and g is the gravity acceleration. Velocity dimensions were obtained by dividing K_Z by $Z/2$, the distance between the gravity centers of the upper and lower layer.

Stability. The water column average Brunt-Väisälä frequency, $\langle N \rangle$ (in min^{-1}), at Stn 00 (Fig. 1) was calculated as follows: $\langle N \rangle = \{4g(\rho_l - \rho_u)/(\rho_l + \rho_u)Z\}^{1/2}$, where ρ_l and ρ_u are the average density of the upper and lower layers.

Chemical variables. Salinity (S , accuracy ± 0.003), N-nutrients (NH_4^+ , $\pm 0.05 \mu\text{M}$; NO_2^- , $\pm 0.02 \mu\text{M}$; and NO_3^- , $\pm 0.1 \mu\text{M}$), DON ($\pm 0.2 \mu\text{mol N l}^{-1}$), suspended PON ($\pm 0.1 \mu\text{mol N l}^{-1}$), and chl *a* ($\pm 0.05 \text{mg m}^{-3}$) were measured as described elsewhere.

Horizontal/vertical fluxes. Average concentrations ($\mu\text{mol kg}^{-1}$) of all species were calculated from the 3 stations located along the middle segment of the ría by means of an inverse distance interpolation method. The horizontal flux of any nitrogen species (F_X , in $\mu\text{mol m}^{-2} \text{d}^{-1}$) was calculated as follows:

$$F_X = V_X C_X \rho_X \quad (5)$$

where V_X is the horizontal velocity, and C_X and ρ_X are the average concentration and density, respectively, at each of 5 sectors defined above.

Vertical convective fluxes (F_Z , in $\mu\text{mol m}^{-2} \text{d}^{-1}$) were obtained as follows:

$$F_Z = V_Z C_Z \rho_Z \quad (6)$$

where V_Z is the vertical convective velocity, and C_Z and ρ_Z are the depth-average concentration and density, respectively, in the study section.

Finally, vertical diffusive fluxes (M_z , in $\mu\text{mol m}^{-2} \text{d}^{-1}$), were obtained as:

$$M_z = \frac{K_z}{\Delta z} (C_B - C_S) \quad (7)$$

where C_S and C_B are the average concentrations of any species in the surface and bottom layer, respectively.

Geochemical budget. The net budget of inputs minus outputs, $i - o$, of any species (C) into a given volume (in $\mu\text{mol C m}^{-2} \text{d}^{-1}$) results from the accumulation, $\Delta(C \cdot V)/\Delta t$, and net ecosystem production, NEP_C , terms:

$$i - o = \frac{\Delta(C \cdot V)}{\Delta t} - \text{NEP}_C \quad (8)$$

The accumulation term (in $\mu\text{mol C m}^{-2} \text{d}^{-1}$) can be separated in 2 terms:

$$\frac{\Delta(C \cdot V)}{\Delta t} = \frac{V \Delta C}{\Delta t} + C \cdot \frac{\Delta V}{\Delta t} \quad (9)$$

The first term on the right of Eq. (9) indicates the net accumulation due to changes in concentration and the second term due to changes in volume. The average NEP_C between 2 consecutive surveys can be estimated as:

$$\text{NEP}_C = o - i + \frac{\Delta(V \cdot C)}{\Delta t} \quad (10)$$

where the balance of $o - i$ (in $\text{mol C m}^{-2} \text{d}^{-1}$) is calculated as:

$$o - i = (F_S \cdot A_S - F_B \cdot A_B)/A \quad (11)$$

F_S and F_B are the average surface and bottom horizontal fluxes (in $\mu\text{mol m}^{-2} \text{d}^{-1}$); A_S and A_B are the areas of the surface and bottom layer of the transverse section (in m^2); and A is the free surface area of the ría bounded by the transverse section ($35.6 \times 10^6 \text{m}^2$).

RESULTS

Water flows

Northerly winds prevailed over the shelf in July (Fig. 2a), while southerly winds were dominant in September (Fig. 3a). Since continental runoff was very low in July ($<9 \text{m}^3 \text{s}^{-1}$; Fig. 2b) and September (from 5.7 to $14.5 \text{m}^3 \text{s}^{-1}$; Fig. 3b), 6 periods were defined according to the dominant shelf wind. The first, from 15 to 18 July, was characterized by an intense upwelling episode that reached a maximum on 15 July. A positive residual circulation pattern with a strong surface outflow of -12cm s^{-1} was recorded (Fig. 2d). During the second period, from 18 to 22 July, northerly winds abated, producing a reversal of the residual circulation pattern (surface inflow and bottom outflow; Fig. 2d). The third period (22 to 26 July) was dominated by

northerly winds that promoted an intense upwelling on 26 July. A positive residual circulation pattern resumed (Fig. 2d), with surface and bottom currents similar to the first period. The fourth period, from 17 to 19 September, was dominated by southeasterly winds that decreased gradually and caused a negative residual current pattern with a maximum surface inflow of 10cm s^{-1} and bottom outflow of -10cm s^{-1} (Fig. 3d). The fifth period, from 19 to 23 September, was characterized by the transition from southerly to northerly winds (Fig. 3a), when the reversal of the residual pattern persisted but current velocities diminished. On 23 September, a 3-layer circulation developed, with outflow through the surface and bottom layers and inflow through the intermediate layer. Finally, the sixth period, from 23 to 26 September, was characterized by moderate northeasterly winds blowing over the shelf with lower intensity than in July (Figs. 2a, 3a). As a result, the 2-layer positive circulation pattern was re-established.

On 15 July, the salinity and temperature distributions showed the typical response to an upwelling event. Despite the subsequent relaxation of shelf winds, the effect of coastal upwelling still remained in the bottom layer 3 d later (Fig. 2e,f). The reversal of the circulation in response to the wind calm (Fig. 2a) affected the water column stratification: salinity was more homogeneous and the halocline was difficult to distinguish, while surface temperature increased and the thermocline was found at 20 m depth. The thermal structure found on 17 September corresponded to a typical autumn downwelling, with warm ($>17^\circ\text{C}$) and relatively salty (>34.8) shelf surface waters in the surface layer (Fig. 3e,f). The persistence of shelf southerly winds produced a deepening of the surface salinity and temperature as a consequence of downwelling. Southerly winds changed to northeasterly on 22 September (Fig. 3a), producing a slight uplift of the isohalines and isotherms, suggesting the entry of oceanic ENACW through the bottom layers on 26 September.

The analysis of the longitudinal component of the residual current shows typical 2-layered positive and negative circulation patterns under upwelling conditions on 15 July (Fig. 4a) and downwelling conditions on 19 September (Fig. 4f), respectively. A variety of intermediate cases were observed on the other sampling dates, showing the complexity of the lateral residual circulation of the ría.

Vertical convective velocities were upwards during the upwelling events of 15 and 26 July (Fig. 4a,d) and 23 and 26 September (Fig. 4g,h). They were nearly zero during the relaxation of upwelling (Fig. 4b) and reversed to downwards on 22 July (Fig. 4c), and on 17 and 19 September (Fig. 4e,f). In July, mixing velocities

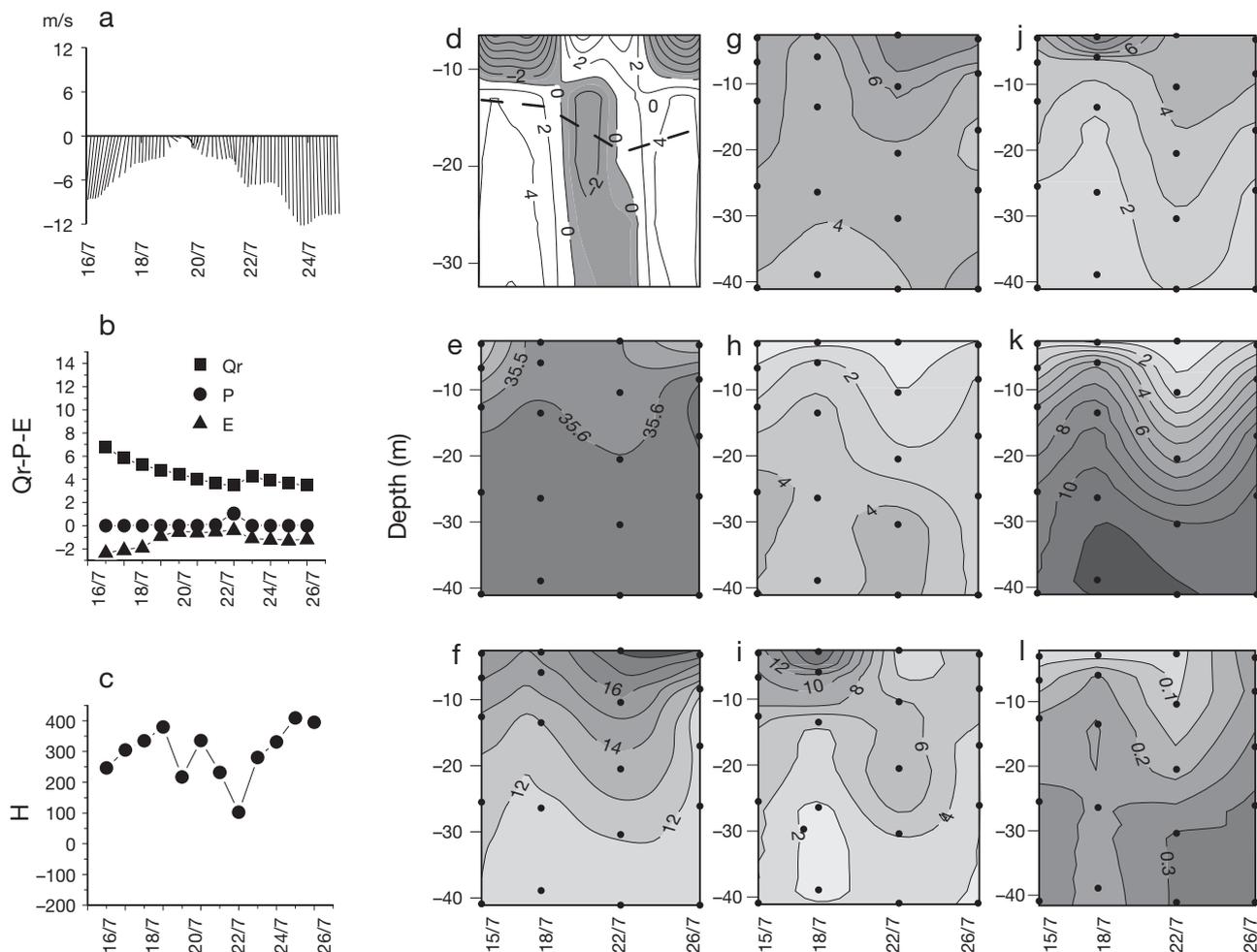


Fig. 2. July 2002. (a) Time course of shelf winds, in m s^{-1} ; (b) continental runoff (Q_r), precipitation (P) and evaporation (E), in $\text{m}^3 \text{s}^{-1}$; (c) heat balance (H), in $\text{cal cm}^{-2} \text{d}^{-1}$; (d) residual currents, in cm s^{-1} ; dashed line represents the pycnocline; (e) salinity; (f) temperature, in $^{\circ}\text{C}$; (g) dissolved organic nitrogen, in $\mu\text{mol l}^{-1}$; (h) ammonium, in $\mu\text{mol kg}^{-1}$; (i) chl *a*, in $\mu\text{g l}^{-1}$; (j) particulate organic nitrogen, in $\mu\text{mol l}^{-1}$; (k) nitrate, in $\mu\text{mol kg}^{-1}$; and (l) nitrite, in $\mu\text{mol kg}^{-1}$. Heat and freshwater fluxes taken from Piedracoba et al. (2005)

reached a maximum during upwelling events, being twice the vertical convective velocity on 15 July (Fig. 4a). Throughout September, mixing velocity was higher than the convective vertical velocity, except on 17 September, and it reached a maximum on 19 September (Fig. 4f). The stratification that took place during the relaxation after the first upwelling episode caused mixing velocity that was half the vertical convective velocity on 26 July (Fig. 4d). The vertical displacements of the LNM were negligible compared to the vertical convective and mixing velocities in July and September.

According to our previous hydrographic and dynamic description, the 6 periods were named as follows: (i) spin-down of strong summer upwelling (15 to 18 July), (ii) summer relaxation (18 to 22 July), and (iii) spin-up of weak summer upwelling (22 to 26 July), (iv) autumn downwelling (17 to 19 September),

(v) autumn transition (19 to 23 September), and (vi) autumn upwelling (23 to 26 September).

Nitrogen fluxes

NO_3^- (Fig. 2k), the most abundant N-nutrient in July (69% of $\text{N}_T = \text{NO}_3^- + \text{NO}_2^- + \text{NH}_4^+$), was at a maximum at the bottom layer during the spin-down of the strong summer upwelling (15 to 18 July) because of the entry of ENACW. In the surface layer, maximum concentrations of chl *a* (Fig. 2i) and PON (Fig. 2j) were recorded. During the summer relaxation (18 to 22 July), lower chl *a* and PON concentrations were recorded at the surface compared with the previous period (15 to 18 July) and a subsurface chl *a* maximum developed at 20 m depth (the nutricline). DON (Fig. 2g) accumulation in the surface layer took place at the end of the

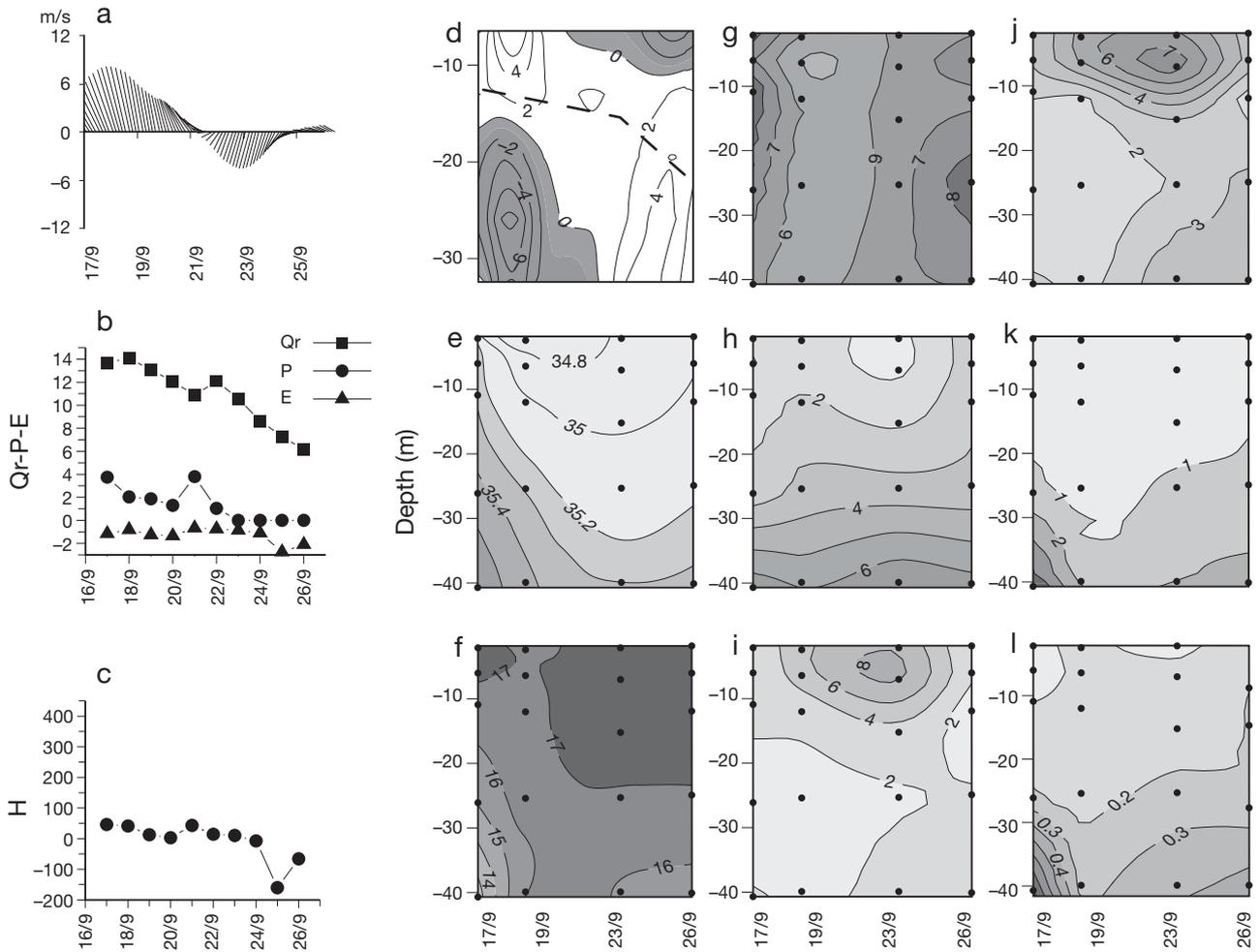


Fig. 3. September 2002. (a) Time evolution of shelf winds, in m s^{-1} ; (b) continental runoff (Q_r), precipitation (P) and evaporation (E), in $\text{m}^3 \text{s}^{-1}$; (c) heat balance (H), in $\text{cal cm}^{-2} \text{d}^{-1}$; (d) residual currents, in cm s^{-1} (dashed line: pycnocline); (e) salinity; (f) temperature, in $^{\circ}\text{C}$; (g) dissolved organic nitrogen, in $\mu\text{mol l}^{-1}$; (h) ammonium, in $\mu\text{mol kg}^{-1}$; (i) chl a , in $\mu\text{g l}^{-1}$; (j) particulate organic nitrogen, in $\mu\text{mol l}^{-1}$; (k) nitrate, in $\mu\text{mol kg}^{-1}$; and (l) nitrite, in $\mu\text{mol kg}^{-1}$. Heat and freshwater fluxes taken from Piedracoba et al. (2005)

summer relaxation, reaching concentrations $>7 \mu\text{mol kg}^{-1}$. During the spin-up of weak summer upwelling (22 to 26 July), nutrient-rich water entered through the bottom layers, but this input was weaker compared with the first upwelling period (bottom $\text{NO}_3^- < 11 \mu\text{mol kg}^{-1}$). In contrast with July, N-nutrients levels were reduced by half during September and NH_4^+ (Fig. 3h) was the most abundant form, representing 72% of N_T . During the autumn downwelling (17 to 19 September), nutrient-poor waters entered from the shelf. Concomitantly, lower PON ($\sim 4 \mu\text{mol kg}^{-1}$) and chl a ($\sim 4 \mu\text{g kg}^{-1}$) concentrations were measured at the surface layer compared to July (Fig. 3i,j). DON was higher than PON during this period and presented a homogeneous distribution throughout the water column (Fig. 3j). At the beginning of the autumn transition (19 to 23 September), nutrients levels were still very low in the surface layer ($\text{NH}_4^+ < 1.6 \mu\text{mol l}^{-1}$ and $\text{NO}_3^- < 0.5 \mu\text{mol l}^{-1}$), but

the wind change caused a progressive entry of ENACW in the bottom layer. The levels of NO_3^- and NH_4^+ (Fig. 3h,k) increased especially at the beginning of the autumn upwelling (23 to 26 September), although the concentration of NO_3^- at the bottom was 5 times lower than the levels measured during the weakest summer upwelling (22 to 26 July). As a consequence of increasing nutrient levels, high surface chl a ($>8 \mu\text{g l}^{-1}$) and PON ($>7 \mu\text{mol kg}^{-1}$) concentrations were measured (Fig. 3i,j). As for July, the DON maximum (Fig. 3g) was delayed 3 to 4 d compared with the chl a and PON maxima.

In July, no significant differences in either nutrient stocks or fluxes were observed between the 2 bottom sectors. By contrast, significant differences ($p < 0.05$) in nutrient fluxes between the 3 surface sectors were observed on most sampling dates (Table 2). DON and PON stocks were very homogeneous through the sur-

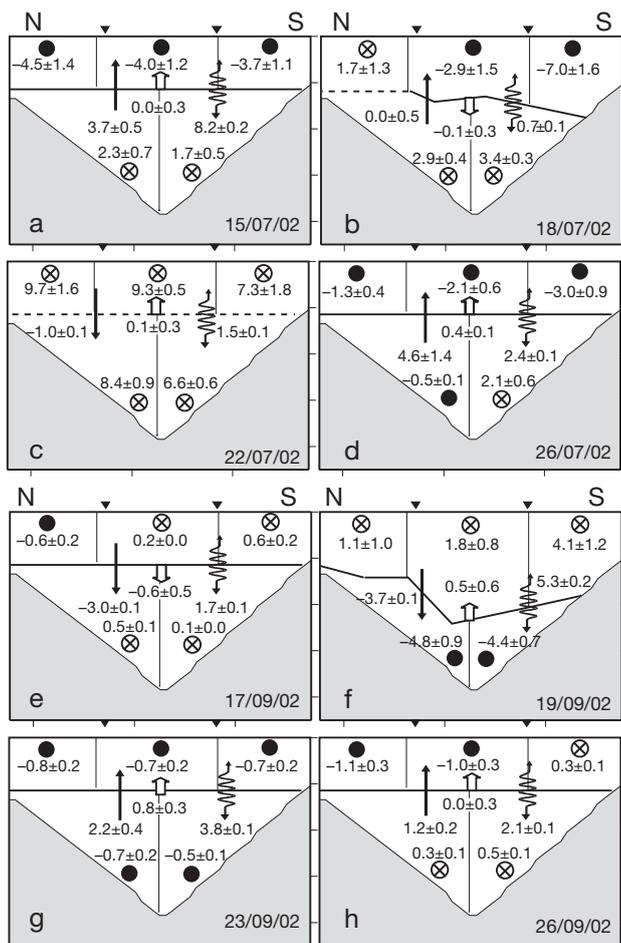


Fig. 4. Time course of horizontal (km d^{-1}) and vertical (m d^{-1}) velocities for each sampling date (dd/mm/yy). Solid circle, outflow ($V < 0$ or westwards) and crossed circle, inflow ($V > 0$ or eastwards) into the ría; solid arrow, vertical advective velocity (upward, >0 ; downward, <0); zigzag arrow, vertical mixing velocity ($K_z/\Delta z$); white arrow, vertical displacement of the LNM. Sectors are described in Fig. 1b

face and bottom layers, whereas DON and PON fluxes frequently exhibited significant differences ($p < 0.05$) between the 3 surface sectors, especially on 26 July. In September, most of the concentration differences along the transverse section occurred in organic nitrogen. As in July, the fluxes of inorganic and organic species were significantly different between sectors ($p < 0.05$), especially on 17 September (Table 3).

Nitrogen budget

A budget of $o - i$ (in $\text{mg N m}^{-2} \text{d}^{-1}$) for N_T , PON and DON was estimated considering the fluxes per m^2 of free surface area from the middle section to the inner reaches of the ría (Table 4). The lateral variability reported in both the N species concentrations and the

longitudinal component of the residual current was considered in the calculation of N-nutrients, PON and DON fluxes and, therefore, in the corresponding $i - o$ balances. In addition, the accumulation terms of all nitrogen species were also calculated ($V \cdot \Delta N_T / \Delta t$), assuming that the concentrations of N_T , PON and DON in the volume of the ría bounded by the middle section do not differ from the concentrations in the middle section. NEP was obtained by summing up $o - i$ and $V \cdot \Delta N_T / \Delta t$ (see Eq. 10) for N_T (NEP_T), DON (NEP_D) and PON (NEP_P) between 2 consecutive surveys.

During the strong summer upwelling, the volume of the ría bounded by the middle section trapped $-289 \pm 17 \text{ mg N m}^{-2} \text{d}^{-1}$, i.e. 62% of the N_T entry, but only 44% of $-289 \pm 17 \text{ mg N m}^{-2} \text{d}^{-1}$ were transformed into organic matter ($NEP_T = -127 \pm 37 \text{ mg N m}^{-2} \text{d}^{-1}$), while $162 \text{ mg N m}^{-2} \text{d}^{-1}$ accumulated into the volume. By contrast, the ría was a nutrient source during the subsequent upwelling relaxation ($o - i = 101 \pm 5 \text{ mg N m}^{-2} \text{d}^{-1}$) at the expense of the nutrients accumulated during the previous period. During the following weak upwelling episode, nutrient salts were trapped again, with 57% utilized for organic matter production and 43% accumulated into the study volume. During the autumn downwelling the ría acted as a nutrient source ($o - i = 119 \pm 9 \text{ mg N m}^{-2} \text{d}^{-1}$) at the expense of the nutrients previously accumulated in the system, whereas NEP_T was balanced ($-18 \pm 12 \text{ mg N m}^{-2} \text{d}^{-1}$). However, during the subsequent relaxation period, the ría showed a heterotrophic metabolism ($NEP_T > 0$), exporting 28% of the regenerated N_T to the shelf ($o - i = 19 \pm 1$) and accumulating the remaining 72% into the ría. Finally, during the weak autumn upwelling, the ría acted as a nutrient trap, $-96 \pm 8 \text{ mg N m}^{-2} \text{d}^{-1}$, one third lower than during the weak summer upwelling. Only 27% of N_T was consumed to produce organic nitrogen and the remaining 73% accumulated into the system. As a result, the system changed from heterotrophic to balanced conditions ($NEP_T \sim 0$). During July, the proportion of NEP_T due to NO_3^- and NH_4^+ was 67% and 33% respectively, whereas during September, the values were 45% and 55% respectively.

The fraction of NEP_T transferred to the sediments (NEP_{SED}) can be inferred indirectly from the following mass balance equation:

$$NEP_T + NEP_D + NEP_P + NEP_{SED} = 0 \pm \delta \quad (12)$$

NEP_{SED} is the balance between PON sedimentation and PON resuspension + N_T diffusion from the sediment (Álvarez-Salgado et al. 1996a). It should be noted that NEP_{SED} is calculated with a large uncertainty ($\pm \delta$), resulting from the individual uncertainties associated with the calculation of NEP_T , NEP_D and NEP_P . During the strong and weak upwelling periods, a large frac-

Table 2. *t*-test of the concentrations and horizontal fluxes of N_T , DON, PON, NO_3^- , NO_2^- and NH_4^+ between the 3 surface (NS, CS and SS) and the 2 bottom (NB and SB) sectors, for each sampling date in July 2002. The *p*-level was calculated as follows: if $(\bar{x} - \bar{y})/\sqrt{\sigma_x^2 + \sigma_y^2} > 2.015$, the differences between 2 sectors are significant at $p < 0.05$ (indicated by *), whereas if $(\bar{x} - \bar{y})/\sqrt{\sigma_x^2 + \sigma_y^2} < 2.015$, the difference is nonsignificant (indicated by ns). \bar{x} , σ_x^2 and \bar{y} , σ_y^2 the mean and variance of the concentrations or fluxes in 2 sectors respectively

Date	Sectors	Concentrations						Fluxes					
		N_T	DON	PON	NO_3^-	NO_2^-	NH_4^+	N_T	DON	PON	NO_3^-	NO_2^-	NH_4^+
15 July	SS-CS	*	*	ns	ns	*	ns	*	ns	ns	*	ns	*
	CS-NS	*	*	*	ns	ns	*	*	*	ns	*	ns	*
	SS-NS	*	*	ns	ns	*	*	*	*	ns	ns	ns	*
	NB-SB	*	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns
18 July	SS-CS	*	ns	*	*	ns	*	*	*	*	*	*	*
	CS-NS	*	ns	ns	*	*	*	*	*	*	*	*	*
	SS-NS	*	ns	*	*	*	*	*	*	*	*	*	*
	NB-SB	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns
22 July	SS-CS	*	ns	ns	*	ns	*	*	*	*	*	*	*
	CS-NS	*	ns	ns	*	ns	*	*	ns	ns	*	*	*
	SS-NS	ns	ns	ns	ns	ns	ns	ns	*	*	ns	ns	ns
	NB-SB	*	ns	ns	ns	ns	ns	*	*	*	*	ns	*
26 July	SS-CS	ns	ns	ns	*	*	*	*	*	*	ns	ns	*
	CS-NS	*	ns	ns	ns	*	*	*	*	*	*	*	*
	SS-NS	*	ns	ns	ns	*	*	*	*	*	*	*	*
	NB-SB	ns	ns	ns	ns	ns	ns	*	*	*	*	*	*

Table 3. *t*-test of the concentrations and horizontal fluxes as in Table 2, but for September 2002

Date	Sectors	Concentrations						Fluxes					
		N_T	DON	PON	NO_3^-	NO_2^-	NH_4^+	N_T	DON	PON	NO_3^-	NO_2^-	NH_4^+
17 Sept	SS-CS	*	*	*	ns	*	ns	*	*	*	*	*	*
	CS-NS	*	*	*	ns	ns	ns	*	*	*	*	*	*
	SS-NS	*	ns	ns	ns	ns	ns	*	*	*	*	*	*
	NB-SB	ns	*	ns	ns	ns	ns	*	*	*	*	*	*
19 Sept	SS-CS	ns	*	*	ns	ns	ns	*	*	*	ns	*	*
	CS-NS	ns	*	ns	ns	ns	ns	ns	ns	ns	*	ns	ns
	SS-NS	ns	*	*	ns	*	ns	*	*	*	*	*	*
	NB-SB	ns	*	*	ns	ns	ns	ns	ns	*	*	ns	ns
23 Sept	SS-CS	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns	ns
	CS-NS	ns	ns	ns	ns	*	ns	ns	ns	ns	*	ns	*
	SS-NS	ns	ns	ns	ns	ns	ns	*	ns	ns	*	ns	*
	NB-SB	*	ns	*	ns	ns	ns	*	ns	ns	*	*	*
26 Sept	SS-CS	ns	*	*	ns	ns	ns	*	*	*	*	*	*
	CS-NS	ns	*	*	ns	ns	ns	ns	ns	ns	ns	ns	ns
	SS-NS	ns	*	ns	ns	ns	ns	*	*	*	*	*	*
	NB-SB	ns	ns	ns	ns	ns	ns	ns	ns	ns	*	*	ns

tion of NEP_T sank to the bottom ($75 \pm 35\%$ on 15 to 18 July and $97 \pm 27\%$ on 22 to 26 July). During the autumn downwelling, the net sedimentation (NEP_{SED}) estimated with Eq. (12) is balanced by the loss of the DON previously accumulated in the system, since NEP_T was nearly zero. Transformation of DON into sinking particles implies the use of DON as a nutrient source for microheterotrophs (mainly bacterioplank-

ton) and its transference to higher trophic levels (zooplankton, hanging mussels) to become sinking particles. The autumn transition was the only period when NEP_{SED} was < 0 , which means that resuspension of PON and/or diffusion of N_T from the sediments was larger than POM deposition over the bottom.

The relative importance of accumulation versus export of DON and PON was also assessed in each period.

Table 4. Net budget of outputs minus inputs ($o - i$), accumulation ($V \cdot \Delta N_T / \Delta t$), and net ecosystem production (NEP) and sedimentation (SED) of N_T , DON and PON in July and September 2002 (in $\text{mg N m}^{-2} \text{d}^{-1}$)

Period	Interval	$(o - i)_T$	$V \cdot \Delta N_T / \Delta t$	NEP_T	$(o - i)_D$	$V \cdot \Delta DON / \Delta t$	NEP_D	$(o - i)_P$	$V \cdot \Delta PON / \Delta t$	NEP_P	NEP_{SED}
Strong summer upwelling	15–18 Jul	-289 ± 17	162 ± 21	-127 ± 37	61 ± 8	-85 ± 2	-24 ± 6	91 ± 6	-34 ± 18	57 ± 24	95 ± 44
Summer relaxation	18–22 Jul	101 ± 5	-136 ± 1	-35 ± 5	-13 ± 4	41 ± 11	28 ± 15	-21 ± 3	35 ± 22	14 ± 25	-6 ± 30
Weak summer upwelling	22–26 Jul	-280 ± 14	121 ± 20	-159 ± 34	21 ± 7	-7 ± 8	15 ± 14	36 ± 6	-46 ± 17	-10 ± 24	154 ± 44
Autumn downwelling	17–19 Sep	119 ± 9	-137 ± 4	-18 ± 12	-9 ± 13	-155 ± 5	-163 ± 8	-94 ± 8	110 ± 16	16 ± 25	165 ± 29
Autumn transition	19–23 Sep	19 ± 1	47 ± 14	66 ± 14	1 ± 1	49 ± 10	50 ± 11	-15 ± 2	32 ± 22	16 ± 24	-132 ± 30
Autumn upwelling	23–26 Sep	-96 ± 8	70 ± 15	-26 ± 23	-29 ± 9	68 ± 9	39 ± 18	39 ± 6	-68 ± 14	-29 ± 21	15 ± 36

The balance of $o - i$ showed that the system acts normally as a source of organic matter under upwelling conditions ($o - i > 0$) and as a trap under downwelling or relaxation conditions ($o - i < 0$). However, it is necessary to compare this with the accumulation of DON and PON to establish the metabolic status of the system independent of the hydrodynamic conditions (Table 4). During the strong summer upwelling (15 to 18 July) the decrease in DON was due to export (71%) and oxidation (29%), whereas 63% of the PON exported to shelf surface waters was produced *in situ* and the remaining 37% came from materials previously accumulated in the ría. During the summer relaxation (18 to 22 July) progressive accumulation of DON and PON occurred: 68% of the DON and 61% of the PON accumulated in the study volume came from *in situ* consumption of N_T and the remaining 32% and 39% of DON and PON came from allochthonous materials. During the weak summer upwelling (22 to 26 July) 69% of the DON exported was produced *in situ* and the remaining 31% of DON exported to the shelf came from the materials previously accumulated in the ría. In contrast, 78% of the PON accumulated previously was exported and the remaining 22% was oxidized inside the system. During the autumn downwelling (17 to 19 September), there was a net consumption of DON; only 5% was imported and 95% came from the material previously accumulated into the system. However, the particulate fraction was accumulated (15% of PON was produced *in situ* and 85% of PON was imported from the shelf). During the autumn relaxation (19 to 23 September) nearly 100% of the DON and 50% of the PON accumulated in the study volume was produced *in situ*. Finally, during the weak autumn upwelling (23 to 26 September), DON accumulated in the study volume: 42% imported from the shelf and 58% produced *in situ*. In contrast, the PON losses were due to PON consumption (42%) and export (58%).

DISCUSSION

Vertical convection and diffusion in the fertilization of the photic layer

The impact of shelf winds on the horizontal circulation of estuaries and coastal inlets has been discussed thoroughly in the literature (Carter et al. 1979, Wong & Moses-Hall 1998, Janzen & Wong 2002). In the case of the Galician rías, the issue has also been considered (Rosón et al. 1997, Álvarez-Salgado et al. 2000, Piedra-coba et al. 2005). The vertical circulation is also relevant because the fluxes resulting from the combination of convective/diffusive velocities and the gradient of nutrient concentrations are responsible for the fertilization of the photic layer. However, few studies have dealt with the vertical fluxes, mainly because vertical currents are difficult to obtain directly from current meters compared with horizontal currents. In fact, most studies discussing vertical fluxes were carried out in the Galician rías (Álvarez-Salgado et al. 1996a, Rosón et al. 1999, Pérez et al. 2000).

Water flows and nutrient fluxes in the rías were originally calculated with a 2D steady-state box model using salinity as a tracer (Prego & Fraga 1992). This model did not adequately reproduce the transient upwelling events occurring in the rías. Box models improved when a transient and salt-heat-weighted version was applied to the rías of Arousa (Rosón et al. 1997), Vigo (Álvarez-Salgado et al. 2000) and Pontevedra (Pardo et al. 2001). Finally, a 3D version has been recently developed by Gilcoto et al. (2007), after empirical and numerical confirmation of a 3D residual circulation pattern in the outer part of the Ría de Vigo (Souto et al. 2003). Since the turbulent diffusion velocities obtained by box modeling are as representative for comparable spatial and temporal scales as the turbulent mixing velocities of this study, direct comparisons can be made.

The mixing velocity ($K_Z/\Delta z$) was nearly twice the vertical convective velocity (V_Z) during strong summer upwelling (Fig. 4a). The weakness of coastal winds together with thermal stratification ($\langle N \rangle$ increased by 0.04 min^{-1} between 18 and 22 July) led to nearly no vertical displacements during the subsequent summer relaxation (Fig. 4b,c). Finally, vertical advection became more relevant with the re-establishment of upwelling conditions, although previous stratification reduced the mixing between layers: $(K_Z/\Delta z)/V_Z \sim 0.5$ (Fig. 4d).

Our velocity values are in agreement with the results obtained by Míguez et al. (2001), who applied a 2D transient box model to the middle Ría de Vigo under similar hydrodynamic conditions: an intense upwelling episode after a wind calm in March 1994. However, the $(K_Z/\Delta z)/V_Z$ ratios obtained by Míguez et al. (2001) ranged from 0.5 to 0.7; values of $K_Z/\Delta z > V_Z$ were not recorded. The 3D transient box model of Gilcoto et al. (2007) also produced an average ratio $(K_Z/\Delta z)/V_Z$ of ~ 0.5 for the upwelling period in the middle Ría de Vigo.

In September, the magnitude of the downward convective velocity during the initial downwelling was comparable to the mixing velocity (Fig. 4e), which reached a maximum on 19 September (Fig. 4f) due to the progressive decrease of stratification ($\langle N \rangle$ decreased by -0.05 min^{-1} between 17 and 19 September). $(K_Z/\Delta z)/V_Z$ was comparable with the ratio calculated by Míguez et al. (2001) under downwelling conditions in October 1994: V_Z ranged from -3.7 to -4.7 m d^{-1} and $(K_Z/\Delta z)$ increased from 0.3 to 3.6 m d^{-1} in agreement with the progressive decrease of water column stability. Gilcoto et al. (2007) obtained a ratio > 0.7 during a downwelling period in December.

Subsequently, stratification ($\langle N \rangle$ increased by 0.01 min^{-1} from 19 to 23 September) reduced the mixing velocity during the autumn transition (Fig. 4g), and at the beginning of the autumn upwelling (Fig. 4h), although the decrease was not as intense as in July. In fact, excluding 17 September, mixing velocities were always larger than vertical convective velocities.

Average $(K_Z/\Delta z)/V_Z$ ratios obtained by Álvarez-Salgado et al. (2000) with a 2D transient box model for the whole Ría de Vigo were in agreement with the ratios obtained in the present study. The multiparameter inverse method applied to the Ría de Vigo during September 1990 and 1991 by Álvarez-Salgado & Gilcoto (2004) also reproduced similar average $(K_Z/\Delta z)$ and V_Z in both periods: $< 1.5 \text{ m d}^{-1}$, i.e. low compared with the vertical velocities calculated in the present study.

Our vertical velocities were also comparable with the values obtained in the middle Ría de Arousa for the upwelling and downwelling season of 1989 (Rosón et al. 1999). $(K_Z/\Delta z)/V_Z$ was 0.8 for the up-

welling season, indicating that advection was the prevalent mechanism, whereas in the Ría de Vigo, mixing was sometimes up to twice the advection. During an autumn downwelling period in October 1989, the absolute value of $(K_Z/\Delta z)/V_Z$ was ~ 0.3 in the middle Ría de Arousa. By contrast, $(K_Z/\Delta z)$ was nearly equal to V_Z in the Ría de Vigo under the same conditions. These observations suggest that turbulent mixing is more important in Vigo than in Arousa. This is probably due to the longer flushing time (60%) of the Ría de Arousa, as derived from the relationship between the volume (Arousa, 4.34 km^3 ; Vigo, 3.12 km^3) and the width of the open boundary (Arousa, 6.7 km; Vigo, 7.7 km) in both rías: $(4.34 \times 7.7)/(3.12 \times 6.7)$. A longer flushing time together with a 30% larger free surface area (230 vs. 176 km^2) favors stratification in the Ría de Arousa compared with the Ría de Vigo.

Vertical nutrient fluxes depend not only on the magnitude of $(K_Z/\Delta z)$ and V_Z , but also on the vertical profiles of nutrient concentrations. In July, mixing controlled the vertical transport of N_T ; it was 6 times larger than that by advection during the strong summer upwelling (15 to 18 July). During the upwelling relaxation (18 to 22 July), vertical mixing fertilized the photic layer in spite of the downward advection of N_T . By contrast, the transport of inorganic nitrogen by mixing was one third of that by advection over the upwelling season in the Ría de Arousa (Álvarez-Salgado et al. 1996a). Although mixing fluxes decreased considerably during the autumn downwelling (17 to 19 September) and autumn transition (19 to 23 September) compared with July, they contributed to the fertilization of the photic layer despite the dominant downwelling conditions. Vertical advection and mixing induced by the prevailing external forces (shelf winds, continental runoff) conditions the stability of the water column. An elevated stability can make vertical mixing negligible.

The vertical transport of N_T , i.e. $N_T \cdot V_Z + K_Z \cdot (\Delta N_T/\Delta z)$, depended more on the nutrient profile ($60 \pm 20\%$, $p < 0.08$) than on the vertical velocity ($40 \pm 20\%$, $p < 0.03$) with $r^2 = 0.85$, $n = 6$ and $p < 0.03$ for the multiple linear correlation of $N_T \cdot V_Z + K_Z \cdot (\Delta N_T/\Delta z)$ with the average N_T concentration (\bar{N}_T) and the total (advective + diffusive) vertical velocity ($V_Z + K_Z/\Delta z$).

The hydrography (stratification vs. homogenization) and dynamics (upwelling vs. downwelling) condition the relative importance of vertical convection and turbulent mixing in the fertilization of the photic layer. Mixing became the most important fertilization mechanism under strong upwelling conditions and compensated the downward vertical convection during downwelling episodes, ensuring the fertilization of the photic layer.

A geochemical approach to the origin and fate of biogenic materials

The Ría de Vigo acts as a transient biogeochemical reactor driven by a succession of upwelling/downwelling episodes separated by short periods of wind calm. Therefore, the metabolic condition of this marine ecosystem can be evaluated with a transient geochemical box model. The transient regime of the ría increases its efficiency as a nutrient trap and allows the accumulation of organic matter recently synthesized within the system. The efficiency as a nutrient trap is calculated from the ratio $(o - i)/i$. It was higher under weak ($66 \pm 3\%$) and strong summer upwelling ($61 \pm 3\%$) than under weak autumn upwelling ($56 \pm 3\%$). Similar efficiencies (66%) were obtained during the upwelling season in the adjacent Ría de Arousa (Álvarez-Salgado et al. 1996b). The trapped nutrients can accumulate or be consumed: 56%, 43% and 73% of the net entry of nutrients accumulated in the system during strong and weak summer upwelling and weak autumn upwelling, respectively (Table 4). In this situation, net uptake by phytoplankton was delayed due to the low biomass in newly upwelled waters and the lag time for adaptation to the light conditions. Nutrients accumulated during upwelling support the productivity of the ría during the subsequent relaxation, when no entry of new nutrients from the shelf occurs. In the Ría de Arousa, most of the nutrients trapped during the upwelling season are transformed into organic matter (87%). Hydrodynamic accumulation is more important in the outer ría: up to 40% (Álvarez-Salgado et al. 1996b).

During downwelling/relaxation events, the ría acts as a nutrient source. During the summer relaxation, there was a net loss of the previously accumulated N-nutrients (74% was exported to the shelf and the remaining 26% was consumed by phytoplankton). During the autumn downwelling, there was also a net N-nutrients loss (87% exported, 13% consumed). In contrast, during the autumn relaxation, 28% of the nutrients regenerated into the system were exported and the remaining 72% were accumulated (Table 4).

Under strong upwelling, phytoplankton in the photic layer is exported to the shelf. It is replaced by the slow-growing cells in the upwelled subsurface water, not adapted to the light conditions of the photic layer (Zimmerman et al. 1987), and most of the upwelled N-nutrients accumulate in the photic layer rather than being consumed. For this reason, nutrient uptake by phytoplankton and, therefore, chl *a* and PON accumulation, occurred during the spin-down phase of upwelling events, when nutrients are abundant, horizontal transports are reduced and phytoplankton cells are acclimatized (Zimmerman et al. 1987, Álvarez-Salgado et

al. 1996b, Pérez et al. 2000). The sampling frequency (twice a week) showed that DON and PON accumulated during upwelling relaxation, in agreement with the observations of dissolved organic matter (DOM) accumulation during the decline of phytoplankton blooms observed in other marine systems (Kirchman et al. 1994, Normann et al. 1995).

Average NEP_T calculated with transient box models is suitable to quantify the net trophic balance of a transient ecosystem (Smith & Hollibaugh 1997, Pérez et al. 2000, Álvarez-Salgado et al. 2001). According to our calculations, NEP_T was negative during July, showing an autotrophic status at the ecosystem level. In September, only during autumn transition was the system strictly heterotrophic ($NEP_T > 0$), whereas the rest of the study period was nearly balanced ($NEP_T \sim 0$). The geochemical budget for July 2002 showed a net autotrophic metabolism similar to results of Álvarez-Salgado et al. (2001) during summer 1997. The differences between the 2 periods derive from the relevance of accumulation versus export of organic matter during July 2002, whereas in July 1997 there was a close coupling between hydrodynamics and biogeochemistry resulting in a near-zero accumulation term. Autumn 1997 (Álvarez-Salgado et al. 2001) was comparable with September 2002. In both periods, the NEP_T depended mostly on the nutrient stock previously accumulated in the ría.

The periods 15 to 26 July and 17 to 26 September can be considered representative for the typical upwelling-relaxation and downwelling-relaxation events off NW Spain, respectively. Average NEP during July was $107 \text{ mg N m}^{-2} \text{ d}^{-1}$ ($= -NEP_T$), very close to the estimation of Gilcoto et al. (2001) for the same area ($140 \text{ mg N m}^{-2} \text{ d}^{-1}$, assuming $C/N = 6.7$) during an upwelling relaxation in September 1990. About 25% of this material was exported to the shelf. In September, only the weak autumn upwelling period was productive ($\overline{NEP_T} = -26 \text{ mg N m}^{-2} \text{ d}^{-1}$), and the fraction exported to the shelf was about 40%. These results are consistent with those of Gago et al. (2003), who estimated that 32% of the organic carbon was exported to the shelf. The estimations are also comparable with the values given by Waldron et al. (1998) for the Benguela upwelling system, where 34% of the new production was exported offshore. In the Ría de Arousa, 80% of NEP was exported to the shelf during the upwelling season of 1989 (Rosón et al. 1999), and 50% when only the middle ría is considered (Álvarez-Salgado et al. 1996a). Regarding the POM:DOM ratio of NEP, the estimations of Gago et al. (2003) and our own results yield a PON:DON ratio of 60:40. By contrast, the massive culture of mussels on hanging ropes in the Ría de Arousa causes an inverse PON:DON ratio of 30:70 (Álvarez-Salgado et al. 1996a). The results of our study indicate that the PON:DON

ratio of the exported material was independent of the upwelling intensity. However, Gago et al. (2003) found that this ratio varied within the upwelling intensity during July 1997 in the Ría de Vigo. During the weak autumn upwelling, only PON was exported to the shelf. In contrast, DON was imported from the shelf and accumulated into the system.

The NEP of DON (NEP_D) is coupled with the balance of outputs minus inputs of DON ($o - i$)_D only under weak summer upwelling. The ría acts as a DON source during moderate upwelling, as reported by Gago et al. (2003). The importance of the accumulation term during the rest of July and September leads to a decoupling of the production and export of DON and PON. Our results show that most of the material exported to the shelf comes from losses of the material previously accumulated rather than recently produced.

The need to assess the relative importance of the production and accumulation of organic matter in a transient marine ecosystem makes the geochemical approach presented in this paper the most suitable to characterize the metabolic status of the Ría de Vigo at the ecosystem level. Although several previous studies estimated the NP of the Ría de Vigo from mass balances (Tilstone et al. 2000, Gilcoto et al. 2001, Gago et al. 2003, Álvarez-Salgado & Gilcoto 2004), none of them either used current meter records or considered the lateral variability of the currents and the chemical species concentrations. In addition, the mass balances were obtained simultaneously with oxygen production/respiration, microzooplankton grazing and sedimentation rates in the present study (see Piedracoba et al. 2008, this volume), which allowed, for the first time, a direct comparison of the *in situ* and *in vitro* approaches for the estimation of NP.

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