Soil salinity and its distribution determined by soil sampling and electromagnetic techniques

J. Herrero*, A.A. Ba & R. Aragüés

Soils and Irrigation Dept., Laboratorio Asociado de Agronomía y Medioambiente (DGA-CSIC), 50080 Zaragoza, Spain.

*Corresponding author. Fax: +34 976 716 145. E-mail: jhi@eead.csic.es

Abstract. Diagnosis of soil salinity and its spatial variability is required to establish control measures in irrigated agriculture. This article shows the usefulness of electromagnetic (EM) and soil sampling techniques to map salinity. We analysed the salinity of a 1-ha plot of surface-irrigated olive plantation in Aragon, NE Spain, by measuring the electrical conductivity of the saturation extract (ECe) of soil samples taken at 22 points, and by reading the Geonics EM38 sensor at 141 points in the horizontal (EMH) and vertical (EMV) dipole positions. EMH and EMV values had asymmetrical bimodal distributions, with most readings in the non-saline range and a sharp transition to relatively high readings. Most salinity profiles were uniform (i.e. $EMH = EMV$), except in areas with high salinity and concurrent shallow water tables, where the profiles were inverted as shown by $EMH > EMV$, and by ECe being greater in shallow than in deeper layers. The regressions of ECe on EM readings predicted ECe with $R^2 > 84\%$ for the 0–100 to 0–150 cm soil depths. We then produced salinity contour maps from the 141 ECe values estimated from the electromagnetic readings and the 22 measured values of ECe. Owing to the high soil sampling density, the maps were similar (i.e. mean surface-weighted ECe values between 3.9 dSm⁻¹ and 4.2 dSm⁻¹), although the electromagnetically estimated ECe improved the mapping of details. Whereas soil sampling is preferred for analysing the vertical distribution of soil salinity, the electromagnetic sensor is ideal for mapping the lateral variability of soil salinity.

Keywords: Soil salinity, electromagnetic sensor, olive trees, irrigation, Spain.

INTRODUCTION

In the central valley of the river Ebro, one of the most arid areas in Europe, irrigation is required for profitable agricultural production. However, the aridity, coupled with the saliferous Miocene strata that are present in large areas of the central Ebro valley, has

been further compounded by improper soil and irrigation management, resulting in the development of about 250 000 ha of salt-affected soils (Herrero & Aragüés 1988). This salinity poses a severe limit to agriculture and a threat to its sustainability; so the need for identification, monitoring and control of salinity is increasing in the central Ebro valley, as well as in many similar arid and semiarid irrigated areas of the world.

Moreover, irrigated agriculture in the Ebro valley is adapting to changes in markets as well as new agricultural policies imposed by the European Union. The plot studied in this work is an example of this adaptation, where olive trees are being grown in a saline-sodic soil, which is an unusual environment for this crop. This soil has a complex salinity distribution, and an in-depth knowledge of the patterns will help to design adequate monitoring and agricultural practices.

The electromagnetic measurement of soil salinity is a non-destructive technique based on the emission of a primary electromagnetic wave and the concurrent measurement of a soil-induced secondary wave whose intensity depends, among other variables, on the electrical conductivity of the bulk soil (ECa). The portable electromagnetic sensor EM38 (Geonics Ltd, Canada) is suitable for diagnostic and agronomic purposes, since around 80% of the ECa response is due to the 0-100 cm (horizontal dipole disposition) and 0-200 cm (vertical dipole disposition) soil depths. The EM38 has been used to appraise soil salinity in many studies during the last 20 years (Rhoades et al. 1999). The EM38 has been widely used in the central Ebro valley to map soil salinity at the irrigation district level (Herrero & Bercero 1991; Tedeschi et al. 2001) and at the plot scale (Aragüés 1987; López-Bruna & Herrero 1996), to monitor soil salinity over time (Lesch et al. 1998) and to relate crop yield responses to salinity under natural (Bercero & Aragüés 1996) and artificial field conditions (Aragüés et al. 1992, 1999).

In recent work, Ba (2001) used the EM38 in the central Ebro valley to appraise soil salinity in areas, ranging from experimental plots to medium-sized irrigation districts, including salinity monitoring over time, to evaluate the advantages and limitations of this technique. The results were generally consistent and helpful. It was found that the instrument should be calibrated for each soil and date of measurement because of the influence of soil type, soil temperature, soil moisture and the vertical distribution of soil salinity on the EM38 readings. Although the calibration process requires some soil sampling and analysis, the field and laboratory workload is much less than for a traditional soil survey.

The objectives of the present study were (i) to describe the vertical and lateral distribution of soil salinity in a plot that has undergone several changes in crops and irrigation systems, and (ii) to evaluate the usefulness of the EM38 to map soil salinity, as compared to conventional soil sampling methods.

MATERIALS AND METHODS

Location and description of the plot under study

The plot is located in the middle Ebro valley, close to the town of Callén in the Flumen irrigation district (Figure 1). The climate is characterized by mean annual values of 15.3°C, 434 mm precipitation and 1188mm ET₀, based on records at the Almuniente weather station, located 10 km from the plot. The soil moisture regime is on the border of the xeric and the aridic regimes defined by Soil Survey Staff (1999). The irrigation water is of excellent quality for crop production ($EC < 0.4 \text{ dS m}^{-1}$, $SAR < 1$), although its low EC may cause structural (i.e. clay dispersion) and water penetration problems in these illitic soils. Inadequate water and soil management, the high evaporative demand and the saliferous underlying rock strata, all contribute to explain the widespread occurrence of saline-sodic soils in the Flumen-Monegros district (Vizcayno et al. 1995; Nogués et al. 2000).

The plot is part of the 9-ha Agro-Callén farm, previously irrigated with solid-set sprinklers and currently flood-irrigated, using the concrete ditches built when this area was first irrigated. The sprinkler system was installed in 1980, and at the same time gravel was applied in those areas prone to water ponding. Forage crops were grown initially, but the rise in soil salinity and the low crop prices led the owner to substitute them with rice. The plots were levelled, the soil puddled and flooded, and rice was cropped from 1992 to 1996. In the spring of 1997, the land was graded with a gentle slope to avoid water logging and planted with olive trees (*Olea europaea* L. cv. arbequina). Plot No. 6 (250340 m) was selected for study because of the salinity gradient along it, as evidenced by salt efflorescence, water logging, the presence of halophytes and the variable growth of the olive trees. The trees were well developed on the southern half of the plot with a transition zone of stunted trees leading to dead trees in the remaining 80-100m of the northern part of the plot. This salinity gradient was confirmed by shallow soil samples taken from the three areas in February 1999. The E_{Ce} (saturation extract EC) values were 1.7 dS m^{-1} in the non-saline area, 9.5 dS m^{-1} in the transition area, and 15.6 dS m^{-1} in the saline area.

Field measurements

The EM38 readings (16 June 1999) were made every 10 m along five transects parallel to the nine tree lines. The transects formed an orthogonal grid whose cells measured 10 m $\times$ 8 m giving a total of 141 points for the EM38 readings in the horizontal (EM_H) and vertical (EM_V) dipole positions. In addition, soil temperatures were measured at depths of 20 cm and 40 cm with a digital soil thermometer in order to convert the EM38 readings to a reference temperature of 25°C. These readings were performed several days after an irrigation (i.e., at relatively high soil water contents), as soon as the field was trafficable.

Twenty-two of the 141 points, covering both the entire studied area and the full range of EM38 readings, were selected for soil sampling and EM38 calibration purposes. Within 4 hours of EM38 runs, six soil samples were taken in each of the 22 sampling points at 25-cm depth increments using an Edelman auger. Soil moisture increased with depth at all the sampling points. The 132 samples were air-dried, ground and sieved (< 2 mm). ECe was measured in the soil saturation extract (United States Salinity Laboratory Staff 1954), and the average ECe values were calculated for each of the six soil depths (0-25, 0-50, 0-75, 0-100, 0-125 and 0-150 cm) for EM38 calibration.

The trunk diameter of 341 olive trees, numbered 1-70 of rows 5-9, were measured in September 1999 and 2000. The difference in diameter was plotted against the average ECa measured on both dates.

Data analysis

Both EM38 readings and ECe values were submitted to exploratory data analysis by histograms, measures of central tendency (mean, median and quartiles), and dispersion (standard deviation, variance, and coefficient of variation).

Calibration of EM38

The EM38 readings were calibrated against the ECe to obtain simple and multiple linear regressions of ECe on EM_H and on EM_V . The simple linear regressions were computed for the six increasing sample depths, irrespective of the distribution of their variables.

In addition, multiple linear regressions were also obtained using the log-transformed variables to give a gaussian distribution (Isaaks & Srivastava 1989). Since the co-linearity between EM_H and EM_V is a constraint when computing the regressions of ECe on EM_H and EM_V , we reduced it by taking the difference between the log-transformed EM_H and EM_V values as the second independent variable instead of EM_V

(Lesch *et al.* 1992). We also explored the fourth root transformation of E_{Ce}, EM_H and EM_V, as used by Rhoades *et al.* (1989) for normalization purposes.

In order to choose the best equations for EM38 calibration, all the above regressions were evaluated through the statistical significance of their coefficients of determination and by comparing their slopes and intercepts, and their standard errors.

Spatial variability of soil salinity

The vertical distribution of soil salinity was studied by examining the 0-150 cm E_{Ce} profiles obtained at each of the 22 sample points. We considered that a profile was inverted when the E_{Ce} of the surface layers was greater than the E_{Ce} of the deeper layers. This is an important consideration, since inverted profiles are usually developed in the presence of shallow water tables. The horizontal distribution of soil salinity was studied by examining the E_{Ce} contour lines obtained through interpolation by kriging using the Surfer program. The contour maps were obtained for the 0-100 cm and the 0-150 cm soil profiles from the 22 E_{Ce} measured values and from these measured values plus the 119 E_{Ce} values estimated from the EM38 readings.

RESULTS AND DISCUSSION

Exploratory data analysis and salinity profiles

Frequency histograms of EM_H and EM_V (141 EM38 points and 132 E_{Ce} values measured at the 22 soil sampling points) indicated a bimodal distribution (Figure 2), with most readings below 1.0 dS m⁻¹ and above 1.5 dS m⁻¹. This suggests a narrow transition zone between relatively low and relatively high soil salinity values. However, the distributions were asymmetrical in that most EM38 readings were low (i.e., a large proportion of the plot is non-saline).

The distribution of EM_V was slightly more uniform than the distribution of EM_H (Figure 2), suggesting that the variability of the apparent EC was smaller for the deeper layers. This result agrees with the presence of a shallow water table at depths above 1.5 m in the first 100 m of the plot, which smoothed somewhat the EM_V readings due to over-saturation of the deeper layers in contact with the water table. The EM38 histograms obtained for the 141 EM38 points and for the 22 points selected for soil sampling were similar (Figure 2, and median and mean values in Table 1), indicating that these sampling points properly represented the salinity distribution of the entire plot. This is also evidenced by the mean and the median EM_H and EM_V values

computed for the 141 reading points and the 22 sampling points (Table 1), which were not significantly different ($P > 0.05$).

Table 1 displays relevant statistics of EM_H and EM_V , and of the average 0-100 cm and 0-150 cm EC_e profiles at the 22 sampling points. Both EC_e profiles had similar statistics indicating that, as shown later, the vertical variability of salinity in the deeper horizons was low. The substantial differences observed between the mean and the median is in agreement with the histograms shown in Figure 2. The coefficients of variation of EM_H , EM_V and EC_e were very high (in particular for EC_e) confirming the large variability in the lateral soil salinity.

The EM_H and the EM_V readings were linearly correlated ($P < 0.0001$), although some dispersion was evident for the larger EM_{38} values (Figure 3). Both EM_H and EM_V readings were similar (i.e., slopes close to one, and intercepts close to zero), although for values above 2 dS m^{-1} the EM_H readings were somewhat higher than the EM_V readings. This result suggests that most salinity profiles were uniform (i.e., $EM_H = EM_V$), except at points with high salinity values where the profiles were inverted (i.e., $EM_H > EM_V$).

This conclusion is supported by the EC_e salinity profiles obtained at the 22 sampling points, which show that they were quite uniform for EC_e values below 5 dS m^{-1} , but were more variable and had in general an inverse configuration for EC_e values above 10 dS m^{-1} (Figure 4). Thus, the most inverted profiles occurred at points A2, A5, B6, F4 and F5, located in the northern 60 m of the plot, which is the most saline area with the highest water table. The water table depth in this area varied between 0.6 m and 0.8 m during the irrigation season, against values of 1.5 m or deeper in the rest of the plot. These inverted profiles were therefore the consequence of the upward capillary flow of water from the shallow water table and the subsequent evapo-concentration of water and salts at the soil surface. The use of the EM_{38} to identify shallow water tables should be further pursued, since this application could be of interest for diagnostic and salinity control purposes.

Analysis of the EM_{38} - EC_e calibration equations

The simple linear regressions obtained between the various average EC_e depths (i.e., from 0-25 to 0-150 cm soil depths) and the EM_H and EM_V values measured at the 22 soil sampling points were highly significant (i.e., all R^2 were significant at $P < 0.0001$), with intercept values not significantly different from zero ($P > 0.05$) and slopes significantly different from zero ($P < 0.001$) (Table 2, equations 1 to 12). With

increasing soil depth R^2 values increased, especially for the vertical dipole configuration (EM_V). On the other hand, the slope values decreased with increasing soil depths, from values of around 7.2 for the 0-25 cm soil depth (equations 1 and 2) to values of around 5.3 for the 0-150 cm soil depth (equations 11 and 12, Table 2). These decreases in slopes are attributed to the greater soil water content of the deeper layers in the presence of a relatively shallow water table.

The absence of large residuals for the simple linear regressions leads us to conclude that the calibration equations allow accurate prediction of ECe from EM38 values, especially for soil depths of 100 cm or greater (i.e., $R^2 \geq 84\%$).

The multiple linear regressions of ECe on EM_H and EM_V were calculated for the six consecutive soil depths, and the log-transformed variables were also used to compute the multiple linear regressions of $\ln ECe$ on $\ln EM_H$ and on the $\ln EM_H - \ln EM_V$ difference (following Lesch *et al.* 1992). Similar equations were obtained using the one-fourth power-transformed variables (following Rhoades *et al.* 1989). However, none of these models significantly improved the results obtained using the simple linear regressions. So, we used the simpler interpolation for delineating and mapping the salinity of the studied plot.

Soil salinity maps

Two soil salinity maps (Figure 5) were drawn based on the average ECe profiles for the 0-100 cm (A) and 0-150 cm (B) soil depths obtained from the salinity values measured at the 22 soil sampling points. Two additional maps were based on these measured ECe profiles plus the estimated ECe profiles for the 0-100 cm (C) and 0-150 cm (D) soil depths obtained from the remaining 119 EM38 reading points using calibration equations 7 and 12 (Table 2), respectively. The 0-100 cm profile is relevant for assessing crop response or development of the natural vegetation, whereas the 0-150 cm profile is relevant for assessing the stock of salts in the soil and designing reclamation measures.

The maps of measured ECe (A and B in Figure 5) were similar, and agree with the low variability in vertical soil salinity shown in Figure 4. The only significant difference in these maps is that salinity values $> 12 \text{ dS m}^{-1}$ covered a larger area in the 0-100 cm map than in the 0-150 cm map (i.e., the 14 dS m^{-1} contour line was not present in the 0-150 cm map). Both maps show that the south half of the plot was non-saline (i.e., $ECe < 2 \text{ dS m}^{-1}$ in the area above 140 m in the longitudinal direction) or slightly saline (i.e., $ECe < 4 \text{ dS m}^{-1}$ in the area between 100 m and 140 m), and that there was a sharp

transition zone from this low-saline area to the saline area in the north end of the plot. This salinity pattern is in agreement with the unsymmetrical bimodal distributions of the EM38 readings previously presented in Figure 2.

Also, the distribution of the growth of the olive trees planted in the plot agreed with this salinity distribution. Thus, the one-year (1999 to 2000) growth in trunk diameter of the olive trees was $> 2 \text{ cm yr}^{-1}$ for the area of low salinity (i.e., area beyond 100 m in Figure 6), whereas in the highly saline area (i.e., 0-100 m of the left corner of the plot in Figure 6) most trees had grown very little or were already dead. In fact, of the 341 trees planted in 1997, 173 trees were dead in 1999, increasing to 187 dead trees in 2000.

The ECe-estimated contour maps (C and D in Figure 5) were quite similar to each other, while closely resembling the previous ECe-measured contour maps (A and B in Figure 5). The similarity of these maps is objectively tested later (Table 3). As expected, the introduction of the 119 ECe estimates significantly improved map detail, so that the contour lines had a smoother appearance, and some patches with ECe values above 2 dS m^{-1} were found in an area previously classified as non-saline on the basis of the ECe-measured maps. Obviously, the similarity of the ECe-measured and the ECe-estimated maps was due to the high soil sampling density performed in this study. It should be emphasized that using classical soil survey methodology, this sample density would be orders of magnitude lower, and the EM38 maps would then greatly improve the delineation of the spatial variability of soil salinity.

The surface areas within the different salinity intervals in Figure 5 were in general similar for A-D (Table 3). However, (i) the non-saline area (i.e., $\text{ECe} < 2 \text{ dS m}^{-1}$) was around 54% of the total area in the ECe-measured maps, and decreased to values of 37% (0-100 cm) and 49% (0-150 cm) in the ECe-estimated maps, (ii) the areas in the $2\text{-}4 \text{ dS m}^{-1}$ interval were greater in the estimated than in the measured ECe maps, and (iii) the areas for $\text{ECe} > 12 \text{ dS m}^{-1}$ were negligible in the ECe-estimated maps, whereas they were around 4% (0-150 cm) and 6% (0-100 cm) of the total area in the ECe-measured maps (Table 3). In any case, the surface-weighted average ECe values were almost identical in the four maps (i.e., ECe interval from 3.9 dS m^{-1} to 4.2 dS m^{-1}), indicating that the proposed methodology gave similar and consistent values of soil salinity in the studied plot.

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Table 1. Relevant statistics of the EM38 (EM_H and EM_V , $dS\ m^{-1}$) and the EC_e ($dS\ m^{-1}$) values measured for the 141 EM38 points and the 22 soil sampling points.

	EM38 values at the:				ECe values at the 22 sampling	
	141 points		22 sampling points		points	
	EM_H	EM_V	EM_H	EM_V	0-100 cm	0-150 cm
Mean	0.87	0.92	0.96	1.01	4.94	4.76
Median	0.49	0.51	0.59	0.68	2.36	2.45
Minimum	0.33	0.36	0.27	0.29	0.77	0.72
Maximum	2.50	2.34	2.89	2.67	15.1	13.1
Standard deviation	0.62	0.63	0.78	0.75	4.71	4.40
Coefficient of variation %	71	69	81	74	95	92

Table 2. Simple linear regressions of ECe (dS m⁻¹) on EM_H and on EM_V (ECe = a + b EM38) obtained for six average ECe soil depths measured at 22 sampling points.

Computed sampling depth (cm)	EM38 readings	Intercept (a)	Slope (b)	R ² (%)	Eq.
0-25	EM _H	-0.67	7.10	74.0	1
	EM _V	-1.20	7.27	71.2	2
0-50	EM _H	-0.49	6.15	81.3	3
	EM _V	-0.94	6.29	78.0	4
0-75	EM _H	-0.46	5.78	84.6	5
	EM _V	-0.88	5.91	81.1	6
0-100	EM _H	-0.42	5.57	85.9	7
	EM _V	-0.87	5.74	83.7	8
0-125	EM _H	-0.36	5.35	85.3	9
	EM _V	-0.84	5.56	84.5	10
0-150	EM _H	0.23	5.19	85.3	11
	EM _V	-0.74	5.44	86.0	12

Table 3. Surface areas in each of the ECe intervals given in the first column, obtained by planimetry of the contour maps depicted in Figure 5.

ECe interval (dS m ⁻¹)	Measured ECe in 22 soil sampling points; average for depths of:		Measured ECe in 22 soil sampling points plus estimated ECe in 119 EM38 reading points; average for depths of:					
	0-100 cm (Fig. 5A)		0-150 cm (Fig. 5B)		0-100 cm (Fig. 5C)		0-150 cm (Fig. 5D)	
	m ²	%	m ²	%	m ²	%	m ²	%
0-2	4642	54.4	4576	53.6	3143	36.8	4195	49.1
2-4	971	11.4	970	11.4	2538	29.7	1379	16.2
4-6	445	5.2	485	5.7	472	5.5	537	6.3
6-8	385	4.5	438	5.1	527	6.2	575	6.7
8-10	685	8.0	808	9.5	1056	12.4	1044	12.2
10-12	744	8.7	946	11.1	685	8.0	781	9.1
12-14	528	6.2	313	3.7	113	1.3	25	0.3
14-16	136	1.6	0	0.0	1	0.0	0	0.0
>16	0	0.0	0	0.0	0	0.0	0	0.0
Surface-weighted ECe	4.0		4.1		4.2		3.9

Figure 1. Location of the Callén farm and the plot number 6 under study.

Figure 2. Frequency histograms of the EM_H (horizontal dipole configuration) and EM_V (vertical dipole configuration) readings taken in the plot at all the EM38 monitoring points (141) and from soil sampling points (22).

EM38 points and (b) the 22 soil sampling points.

Figure 4. Salinity (EC_e, dS m⁻¹) profiles (0-150 cm) of the 22 soil sampling points.

Figure 5. ECe contour maps of plot number 6 obtained from the measured ECe values at the 22 soil sampling points [average ECe of the 0-100 cm (A) and 0-150 cm (B) profiles] and from these values plus those estimated for the remaining 119 EM38 measurement points [average ECe of the 0-100 cm (C) and 0-150 cm (D) profiles]. The top end of the four maps correspond to the south of the plot.

Figure 6. One-year trunk diameter growth of the olive trees planted in 1997 in an area comprising the first 130 m $\times$ 16 m of the studied plot.

TABLE AND FIGURE CAPTIONS

Table 1. Relevant statistics of the EM38 (EM_H and EM_V , $dS\ m^{-1}$) and the ECe ($dS\ m^{-1}$) values measured for the 141 EM38 points and the 22 soil sampling points.

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Figure 6. One-year trunk diameter growth of the olive trees planted in 1997 in an area comprising the first $130\ m \times 16\ m$ of the studied plot.