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Title: Vegetation changes in the southern Pyrenean flank during the last millennium in relation to climate and human activities: the Montcortès lacustrine record

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Abstract: This paper reports the vegetation changes of the last millennium, as deduced from palynological analysis of a sediment core from lake Montcortès, situated around 1000 m elevation in the southern pre-Pyrenean flank. The record begins in the Middle Ages (~AD 800) and ends around AD1920, with an average resolution of about 30 years. The reconstructed vegetation sequence is complex and shows the influence of both climate and human activities in the shaping of landscape. Pre-feudal times were characterized by the presence of well-developed conifer forests, which were intensely burnt at the beginning of the feudalism (AD 1000) and were replaced by cereal (rye) and hemp cultivation, as well as meadows and pastures. In the 13th century, a relatively short warming likely corresponding to the Medieval Warm Period (MWP) was inferred by the presence of a low Mediterranean scrub community today restricted below 800 m elevation. This community disappeared during the Little Ice Age (LIA) cooling in the 15th century, coinciding with a decline in human activities around the lake. Forest recovery began around AD 1500, at the beginning of the Modern period, coinciding with wetter climates. However, forests retreated again during the 17th century coinciding with a maximum in olive and hemp cultivation. This situation was reverted in post-Modern times (19th century), characterized by an intense agricultural crisis and a significant decline of population that favored forest re-expansion. Correlations with the nearby Estanya lake, situated about 350 m below, provided a more regional picture. Besides some climatic forcing evident in both sequences, human activities seem to have the main drivers of landscape and vegetation change in the southern Pyrenean flank, in agreement with similar conclusions from other studies developed on high-mountain environments.

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Abstract This paper reports the vegetation changes of the last millennium, as deduced from palynological analysis of a sediment core from lake Montcortès, situated around 1000 m elevation in the southern pre-Pyrenean flank. The record begins in the Middle Ages (~AD 800) and ends around AD1920, with an average resolution of about 30 years. The reconstructed vegetation sequence is complex and shows the influence of both climate and human activities in the shaping of landscape. Pre-feudal times were characterized by the presence of well-developed conifer forests, which were intensely burnt at the beginning of the feudalism (AD 1000) and were replaced by cereal (rye) and hemp cultivation, as well as meadows and pastures. In the 13th century, a relatively short warming likely corresponding to the Medieval Warm Period (MWP) was inferred by the presence of a low Mediterranean scrub community today restricted below 800 m elevation. This community disappeared during the Little Ice Age (LIA) cooling in the 15th century, coinciding with a decline in human activities around the lake. Forest recovery began around AD 1500, at the beginning of the Modern period, coinciding with wetter climates. However, forests retreated again during the 17th century coinciding with a maximum in olive and hemp cultivation. This situation was reverted in post-Modern times (19th century), characterized by an intense agricultural crisis and a significant decline of population that favored forest re-expansion. Correlations with the nearby Estanya lake, situated about 350 m below, provided a more regional picture. Besides some climatic forcing evident in both sequences, human activities seem to have the main drivers of landscape and vegetation change in the southern Pyrenean flank, in agreement with similar conclusions from other studies developed on high-mountain environments.

Keywords

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Introduction

The paleoecological study of mountain ecosystems is useful to document the effects of climate warming and to forecast its potential future consequences. A key issue in these type of surveys is the need for disentangling the relative influence of either natural climatic variations or human disturbance on the observed ecological shifts. A strategy is to compare the ecological changes occurred before the development of human societies (for example, during glacial or postglacial times) with those of the last millennia, when human activities have intensely transformed mountain landscapes and ecology. Another possibility is to develop multiproxy studies including independent proxies for climatic, ecological and anthropic processes and compare the results. Due to the available paleoecological information and its recent human history, the Pyrenean range is a suitable setting to use both approaches. On the one hand, the Pyrenees have been recognized as an important glacial refuge area from which European forests expanded during the post-glacial and Holocene warming (Jalut et al. 1992; Montserrat 1992; Brewer et al. 2002; Terhürne-Berson et al. 2004; González-Sampériz et al. 2004, 2005; Leroy and arpe 2007; Benito et al. 2008; Magri 2008). On the other hand, this range has a long history of around 7000 years of human settlement and landscape utilization (Marsan and Utrilla 1996; Galop 2001), with particular incidence in the last millennium (Esteban 2003; Miras et al 2007).

Long-term and historical ecological changes are often reconstructed by pollen analysis, as one of the best indicators of vegetation and climatic change (Birks and Birks 1980). In the Pyrenees, the extension and intensity with which the northern slope has been studied palynologically (reviews by Jalut et al. 1992; Reille and Lowe 1993 and Visset et al. 1996) have not been paralleled in the southern slope, where studies by Montserrat (1992), González-Sampériz et al. (2006) and Miras et al. (2007), among others, are worth mentioning. Studies covering the last millennium with some detail are also scarce but have recently increased (Galop 2001; Riera et al. 2004, 2006; Morellón et al. 2009b; Ejarque et al. 2009; Pèlach et al. 2009a, b), and a first approach of the recent paleoecological history is starting to emerge. This paper reports the pollen analysis of the uppermost sediments from Lake Montcortès (1027 m elevation), which accounts for the vegetation changes occurred during the last millennium, and compares these results with previous reconstructions from the neighbor Lake Estanya (670 m

elevation) during the same period (Riera et al. 2004, 2006; Morellón et al. 2009b). Paleoclimatic trends are deduced from the multiproxy study (sedimentology, mineralogy, geochemistry and physical properties) of the laminated Montcortès sediments (Corella et al. this volume); whereas the potential influence of human activities in the ecological changes recorded is derived from more or less detailed historical records from the catchment area (Esteban 2003; Riera et al. 2004; Marugan and Rapalino 2005). Lake Montcortès sediments are an excellent target for high-resolution studies with an accurate time control due to its relatively high sedimentation rates and the presence of varves (Corella et al. this volume). These are the ideal conditions to study the last millennia and try to disentangle the effects of natural environmental changes and human activities in the shaping of present-day landscapes.

Study area

Geographical setting and climate

Lake Montcortès is situated in the southern flank of the Central Pyrenees, in the Pallars region, at 42° 19' N - 0° 59' E and 1027 m altitude (Fig. 1). This lake lies on karstic terrains characterized by the presence of Triassic limestones, marls and evaporites (Keuper and Muscheskalk facies), and Oligocene carbonatic conglomerates (Rosell 1994; Gutiérrez et al. 2008). The catchment is small and the lake is mainly fed mainly by groundwater, with intermittent small creeks and scattered springs. The main water outputs are due to evaporation and a small seasonal outlet at the north. The lake perimeter is roughly circular, with a diameter between 400 and 500 m and a maximum water depth of 30 m near the center (Camps et al. 1976; Modamio et al. 1988). The annual average temperature is 10.6 °C, ranging from 1.9 °C in January to 20.3 °C in July. The total annual precipitation is 860 mm, with March as the driest month (46.6 mm) and May as the wettest one (99.2 mm) (Corella et al. this volume).

Vegetation

The lake is around the altitudinal boundary between the Mediterranean lowlands and the Middle Montane belt, which in the Pyrenees is situated around 800-1000 m elevation, depending on local conditions (Vigo 2008). Three major forest formations coincide

1 around the lake, as representative of this boundary condition (Fig. 2). The evergreen oak
2 forests dominated by *Quercus rotundifolia* are representative of the Mediterranean
3 lowlands, while the deciduous oak forests dominated by *Quercus pubescens* represent
4 the Middle Montane belt, with higher total annual precipitation. The conifer forests of
5 *Pinus nigra* subsp. *salzmanii* are mostly secondary and colonize, or eventually replace,
6 the degraded deciduous oak forests (Folch 1981). Carreras et al. (2005-2006), using the
7 CORINE system (CBM 1991), distinguished twelve vegetation types in the catchment
8 (Fig. 2, Table 1). Around the lake, there is a belt of littoral vegetation dominated by
9 *Juncus*, *Scirpus*, *Phragmites*, *Typha* and *Sparganium* (Camps et al. 1976).

17 Historical background

21 A summary of the main historical traits of the Pallars region is shown in Figure 3. The
22 beginning of the Medieval period was marked by the defeat of the Roman Empire by
23 the Visigoths (around AD 400), who dominated until the early 8th century. The region
24 was then under the dominance of the Muslims until about AD 800, when the
25 Carolingians established. The last millennium has been subdivided into three main
26 periods (Esteban 2003; Marugan and Rapalino 2005): Medieval Ages (8th to 15th
27 centuries), Modern Ages (16th to 18th centuries) and Post-modern Ages (19th century
28 onwards). Until the 10th century, the southern Pyrenean flank was a poor and
29 overpopulated region, due to the pressure of the northward Muslim expansion. The land
30 was subdivided into small properties ruled by autonomous landholders. Forest clearance
31 by fire was a common practice to prepare the terrains for shifting cultivation, the main
32 economic activity at the time. After AD 1000, the individual properties were confiscated
33 by the nobility and the church, thus determining the establishment of the feudal system,
34 which endured until the end of the Middle Ages. Fire and shifting cultivation were
35 replaced by a more permanent land exploitation system. The increase in cattle, mainly
36 sheep, led to deforestation and the development of meadows and pastures. The more
37 common crops were cereals (wheat, barley and oats), olive and vine. The last two,
38 however, were restricted to the southern lowlands (less than about 1000 m elevation),
39 due to climatic constraints. Towards the end of the Medieval period, a significant
40 depopulation of the Pallars region was caused by internal wars, disease epidemics and
41 the Little Ice Age cooling (Marugan and Oliver 2005).

1 The feudalism collapsed and was replaced by the monarchy around AD 1500
2 (Modern Ages). This determined the return to a private property system, but this time
3 the land owners were rich families and an oligarchic system established. The Pallars
4 region became a marquisate and was newly populated by immigrants from the north
5 (presently France). Agriculture and cattle raising experienced a significant expansion
6 and diversification. The development of extensive livestock practices and the common
7 use of animals for agriculture purposes (such as plowing and fertilizing) were especially
8 significant. Forests were exploited mainly for wood, other common activities were
9 mining, hunting and fishing. At the end of the Modern period, the commerce began to
10 be important (Bringué 2005).
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20 The post-modern period began with the nascent capitalist society typical of the
21 19th century, characterized by the centralization of the production system and the
22 existence of a well-developed route system for commercial interchange. The Pallars
23 region, however, remained isolated and self-sufficient subsistence economy prevailed
24 during the 19th century. At the beginning of the 20th century, the general agricultural
25 crisis caused a new significant depopulation in the region, but the development of an
26 extensive hydroelectrical power program soon promoted a socio-economic revival. The
27 Civil War (AD 1936-1939) interrupted again the modernization trend, but a second
28 hydroelectric program and a new mining development, revitalized the economy.
29 Between 1960 and 1980, the industrialization caused a global crisis manifested in the
30 massive emigration from rural to urban areas. Today, the main economic activities of
31 the Pallars region are tourism and related services (Farràs 2005).
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43 **Material and methods**

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47 In April 2004, four sediment cores were obtained in the deepest part (about 30 m water
48 depth) of the Montcortés lake following a NW-SE transect using a Kullenberg corer and
49 platform of the Limnological research Centre (LRC), University of Minnesota. All the
50 cores were split, imaged and their physical properties measured with a Geotek (Corella
51 et al. this volume). Core MON-04-1A-1K, 6.69 m long, was selected for palynological
52 analysis and sampled every 10 cm. In this paper, we report the results of the upper 3.60
53 m, roughly corresponding to the last millennium. Chronology was taken from the age-
54 depth model of Corella et al. (this volume), based on radiocarbon dating of
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macroremains and bulk sediment samples. The average resolution of this pollen record is around 33 years per sampling interval, ranging between about 14 and 59 years.

Pollen samples consisted of ~3-5 g of sediment, which were processed with standard palynological methods, including NaOH, HCl and HF digestions and density gradient centrifugation, without acetolysis (Bennett and Willis 2001). Two *Lycopodium* tablets (batch n° 483216; 18,583 spores/tablet) were added to each sample before chemical processing. Residues were suspended in glycerine and slides were mounted with glycerine-jelly. Pollen and pteridohyte spores were identified and counted until diversity saturation (Rull 1987), obtaining total counts between 289 and 873 (average 427). Algae remains were also identified to the genus level and counted. Fungal spores and charcoal particles above 5 µm were counted as a group. The percentage diagram was based on the pollen sum (ΣP), which included all pollen types except those from aquatic plants (Cyperaceae, *Cladium*, *Myriophyllum*, *Pinguicula*, *Potamogeton*, *Ranunculus*, *Typha/Sparganium*-type and *Utricularia*), ranging from 266 to 866 (average 421) counts. Pollen and spore types below 0.1% of the ΣP were not represented, these are: *Acer*, *Androsace*, *Asphodelus albus/ramosus*-t, *Centaurea*, *Convolvulus*, *Cornus*, *Cupressus*, *Draba*-t, *Echium*, *Gentiana*, *Geranium*, *Linum*, *Myriophyllum*, *Paronychia*-t, *Pedicularis*, *Pinguicula*, *Polygonum aviculare*-t, *Populus*, *Potamogeton*, *Prunus avium*-t, *Ranunculus*, *Rhamnus*, *Sideritis*, *Tilia*, *Utricularia*, *Valeriana*, *Huperzia*, *Isoëtes* and *Lycopodium cernuum*. Zonation was performed using the Optimum Splitting by Information Content (OSIC) and the broken-stick methods (Bennett 1996), considering only pollen types. Pollen groups (Table 2) were defined according to the present-day vegetation types. Percentages for non-pollen palynorphs or NPP (algae, fungi, etc.) were referred to pollen sum.

Results and interpretation

The pollen diagram is dominated alternatively by trees and herbs, while shrubs are relatively scarce (Fig. 4). Among trees, the more abundant are *Pinus* and *Quercus* (evergreen) followed by *Olea*, whereas the dominant herbs are Poaceae (others), *Plantago*, *Cannabis*-type and *Artemisia*. Cyperaceae dominate the aquatic assemblage, in which *Typha/Sparganium* is important only in the upper part. Fern spores are

comparatively scarce, in general. Total pollen and charcoal particles have their maxima near the base. The dominant vegetation types are conifer forests at the beginning, meadows/pastures in the middle, and herbaceous crops at the end. Oak forests show a progressive increase towards the top, while ruderal plants and weeds decline from the base to the top (Fig. 5). The following pollen zones have been identified:

Zone M-1 (360-325 cm, 1172-1000 cal BP, AD 778-950, 4 samples, 45 years/sampling interval)

This zone is dominated by *Pinus*, followed by *Quercus* (evergreen) and Poaceae (others), which show a decreasing trend towards the top. On the contrary, *Cannabis*-type, *Artemisia* and psilate monoletes increase in the same sense. These patterns coincide with an increment in charcoal concentration. Conifer forests are the dominant vegetation type (Figures 4 and 5). Meadows/pastures are important at the base but they decline around the middle, where herbaceous crops and ruderal/weeds start to increase. *Pseudoschizaea* has its maximum development and *Botryococcus* is poorly represented (Fig. 6). Fungi are also at maximum values and increasing. Among them, *Glomus* is well represented. The whole picture suggests an initial pine woodland landscape around AD 800, being progressively reduced by fires and replaced by crops and associated elements (ruderal plants and weeds). The abundance of *Pseudoschizaea* and the scarcity of planktonic algae is consistent with low lake levels or marshy environments, indicative of dry climates (Scott 1992; Carrión 2002, Carrión et al. 2004, 2007). The presence of *Glomus* has been interpreted in terms of increased soil erosion (Anderson et al. 1984, van Geel et al. 1989, 2003; Leroy et al. 2009), and associated to intensification of human activities in the basin (López-Sáez et al. 2009, Argant et al. 2006). It should be also noticed that *Glomus* is a genus of mycorrhizal fungi commonly associated to forest vegetation. Forest management experiments have shown that mycorrhizal colonization of roots in pine and oak forests is significantly greater if half of the forest is removed, as compared to uncut and clearcut treatments (Zhou et al. 1997). Therefore, a medium-impact clearance of *Pinus* forests is suggested for the time interval between ca AD 780 and 940.

Zone M-2 (325-203 cm, 1000-834 cal BP, AD 950-1116, 12 samples, 14 years/sampling interval)

A dramatic decline in conifer forests, due basically to *Pinus*, marks the beginning of this zone (Fig. 4). Oak forests (both evergreen and deciduous) and shrubs slightly increase. The dominance is now shared by meadows/pastures (mainly Poaceae and *Plantago*) and herbaceous crops (*Cannabis*-type). The first dominate in the middle of the zone, coinciding with a maximum of ruderals/weeds, while the second peak at the beginning and at the end. Cultivated trees (mostly *Olea*) start to increase in the middle of the zone. These shifts in the pollen assemblages occur at the same time than an abrupt increase in both pollen and charcoal concentration, which attain their respective maxima. This pollen zone is interpreted in terms of intense land use, mainly cereal cultivation alternating with grazing -as indicated by the increase of meadows and pastures- after an abrupt and severe reduction of conifer forests using fire. Among non-pollen palynomorphs (NPP), *Pseudoschizaea* and *Glomus* drastically decline, while *Tetraedron* shows a conspicuous peak coinciding with the main charcoal maximum (Figs. 4 and 6). The decline in *Glomus* is consistent with a strongly reduced cover of *Pinus* forests (Zhou et al. 1997). The charcoal peak at the beginning of the zone coincides with a peak in *Cannabis*-type and *Secale*, whereas the dramatic charcoal decrease occurred around 290 cm (~AD 1000) coincides with the increase of meadows/pastures and ruderal/weeds. This suggests that, at that time, the use of fire was linked to herbaceous cropping rather than to pastoralism. The *Tetraedron* peak would be indicative of lake fertilization due to increased nutrient releasing linked to higher organic-rich sediment input.

Zone M-3 (2203-112 cm, 834-494 cal BP, AD 1116-1456, 9 samples, 39 years/sampling interval)

Evergreen and deciduous forests do not show significant changes with respect to zone M-2. The main differences occur in the formerly dominant meadows/pastures and herbaceous crops, which start a maintained decreasing trend, and the low scrubs (*Rosmarinus*-type, *Helianthemum*, *Hedysarum*-type and *Ephedra*), which peak in the middle of the zone and decrease around 140 cm (~AD 1360). This increase in low scrubs occurs shortly after a secondary charcoal peak, which decreases in the middle of the zone and attain minimum values at the end, where cultivated trees (*Olea*) start to increase. Ruderal/weeds show a maximum around the middle of the zone, just after the

low scrub peak and before the *Olea* maximum. Total pollen attains its minimal values in this zone. Today, the same low scrub taxonomic association that characterized this zone corresponds to the *Rosmarinetalia officinalis* order (27.C), a thermophilous community typical of Mediterranean calcareous lowland dry environments below 800 m elevation (de Bolós 2001). This association is absent today at Lake Montcortès elevations, which suggests an upper displacement of at least 200 m of this vegetation type due to a warming between about AD 1100 and 1350. The small charcoal peak at the beginning of the zone is not followed by an increase of any taxa associated to human activities and would therefore support the occurrence of drier climatic conditions, favorable for wild fires. In spite of the low or null fire incidence, *Pinus* forests did not recover and only evergreen oak forests and shrubs expanded, suggesting drier climates and/or an intensification of seasonality. The appearance of the salinity-tolerant sedge *Cladium* would also support drier climates favoring increased evaporation. Also noteworthy is the cessation of cereal and hemp cultivation, and the increase of olive crops, since about AD 1280.

Zone M-4 (112-72 cm, 494-337 cal BP, AD 1456-1613, 4 samples, 29 years/sampling interval)

The main changes in this zone are the sharp increase of conifer forests, which peak around the middle, and the equally abrupt decline of meadows/pastures and ruderal/weeds group (mostly *Artemisia*). Oak forests, especially the deciduous ones, and scrublands also increase. Among cultivated plants, *Olea* shows a slight decrease and cereals are nearly absent, whereas hemp is very low at the beginning, but starts to increase around the middle of the zone, after the *Pinus* peak. There is no significant change in charcoal influx, which is at low values. The general decline of all cultivated elements and others associated to human activities (ruderal/weeds, meadows/pastures) and the low charcoal values indicate the low human pressure between about AD 1480 and 1590, though hemp cultivation seems to restart around AD 1550. The increase in other elements, especially conifer forests and *Quercus* deciduous forests, together with the increase of *Corylus* and the abrupt *Artemisia* decline, suggest a moderate increase in moisture. Low scrubs are at minimum values, suggesting colder climates.

Zone M-5 (72-20 cm, 337-117 cal BP, AD 1613-1833, 5 samples, 34 years/sampling interval)

This zone is characterized by the decline of almost all pollen groups, except for the spectacular increase of *Cannabis*-type, which largely dominates the assemblage. This coincides with the increase in charcoal and fungi spores (including *Glomus*). The whole picture indicates the reinitiation of human activities around the lake since about AD 1630. Fires slightly reactivated, affecting primarily conifer and oak forests, and cultivation was largely centered on hemp, while *Olea* maintained the importance attained since the end of zone M-3 (AD 1280). The increase of *Typha/Sparganium* and the disappearance of *Cladium* suggest changes in lake shore environments or water salinity, probably freshening.

Zone M-6 (20-5 cm, 117-29 cal BP, AD 1833-1921, 2 samples, 30 years/sampling interval)

This zone is characterized by a conspicuous reversal of both *Pinus* and *Cannabis*-type pollen curves, the other pollen types remaining similar to the former zone. This indicates that hemp cultivation sharply declined around AD 1860, and conifer forests reexpanded.

Discussion

Comparison with sedimentological and historical records

The multiproxy study of the same core analyzed here provided the depositional and environmental history of the lake catchment during the last 6000 years (Corella et al. this volume). The time interval recorded in this study was subdivided into three main sedimentological phases, represented by the lithological units IV, III and II.

Sedimentary facies indicate that the lake has been deep and meromictic during the last 1500 years. During the Medieval epoch, represented by the lithological unit IV (AD 690 to 1460), the lake had a high clastic input. This unit was subdivided into two subunits (IVa and IVb), with the boundary around AD 1000. Unit IVb is dominated by turbiditic Facies 4, reflecting events of high terrigenous input, while Unit IVa is characterized by

1 the dominance of laminated Facies 1, indicating less sediment delivery to the lake and
2 low bottom bioturbation. This sedimentological unit 4 includes embraces pollen zones
3 M-1 to M-3, and coincides with the Medieval epoch. The boundary between
4 sedimentological units IVb and IVa coincides with the end of the main charcoal peak
5 (Fig. 4) and the disappearance of *Pseudoschizaea* (Fig. 6), indicators of forest fires and
6 increased terrigenous inputs to the lake, respectively. According to the historical
7 records, this date (AD 1000) is the end of the pre-feudal epoch, when the predominant
8 land use was forest burning and shifting cultivation, and the beginning of the feudalism,
9 characterized by a more controlled use of fire and a more or less permanent land
10 exploitation (Marugan and Oliver 2005). The same historical records also report the
11 onset of vineyard and olive cultivation around the same times. In the Lake Montcortès
12 pollen diagram, the *Olea* pollen consistently appear at 310 cm (AD 970), but *Vitis* is
13 absent in the whole sequence. A possible explanation is that, during that time, olive and
14 vineyard were cultivated mainly in lowlands below 1000 m elevation for climatic
15 reasons (Marugan and Oliver 2005), and *Olea* pollen was better dispersed upwards than
16 *Vitis*. In support to this, a recent work in a nearby area has shown the high upward
17 dispersal efficiency of *Olea* pollen, which has been found in significant percentages in
18 surface samples from 800 to >2600 m elevation, in spite of the absence of the parent
19 plant along the whole transect (Cañellas-Boltà et al. 2009). Another apparent
20 inconsistency between historical and pollen records is the high amount of hemp pollen
21 in the diagram and the absence of any reference to its cultivation in the Pallars region
22 since the Medieval Ages to the present (Esteban 2003; Marugan and Rapalino 2005).
23 One possibility is that hemp cultivation was restricted to certain favorable areas and this
24 is not commonly reflected in historical accounts. But it is also possible that these crops
25 were not autochthonous and the lake was used for hemp retting, as it occurred in the
26 neighbor Lake Estanya during the same epoch (Riera et al. 2004). However, this is not
27 reported in the writings reviewed so far and more work is still needed in this respect.
28 Among the cereals mentioned in written documents, the better represented in the pollen
29 diagram are wheat (*Triticum*) and rye (*Secale*), but the Poaceae curve should include
30 other types less characteristic to be identified with confidence.
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56 Concerning forests, pine woodlands were drastically reduced by AD 950, during
57 the pre-feudalism, and did not recover during the whole Medieval epoch, when the
58 landscape was likely a mosaic of predominantly crops and pastures, with smaller
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1 pine/oak forest patches. This human induced vegetation change also coincides with an
2 increase in total pollen concentration, and the dominance of turbidites in the
3 sedimentary sequence indicating enhanced runoff and siliciclastic inputs to the lake
4 (Corella et al. this volume). The dilution effect of higher sedimentation rates on pollen
5 concentrations, often observed elsewhere (e.g. Leroy et al. 2009), was not recorded
6 here. The maximum in pollen concentration corresponds to increases in herbaceous
7 crops, meadows/pastures and ruderal/weeds, and could be explained by enhanced pollen
8 production due to the intensification of human activities linked to these vegetation
9 types.
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18 The vegetation changes recorded until around AD 1100 were mainly triggered
19 by human activities thus potential paleoclimatic signals are likely obscured. Historical
20 records indicate the existence of a phase of “favorable weather” during 10th and 11th
21 centuries (Fig. 3) that favored the expansion of crops and pastures in river margins and
22 wetlands (Esteban 2003; Marugan and Oliver 2005), which is coherent with the
23 interpretation of pollen zone M-2 (Fig. 4). Clearer palynological indications of warming
24 and aridity were found afterwards, between about AD 1100 and 1350, by the presence
25 of low Mediterranean scrub communities in pollen zone M-3. Similar warmer and drier
26 phases have been recorded during the Middle Ages in Estanya and other lakes from the
27 Iberian Peninsula (Morellón et al. 2009a), coinciding with the so-called Medieval Warm
28 Period (MWP) of the Northern Hemisphere (Seager et al. 2007; Mann et al. 2009). The
29 Medieval period ended with a cooling, as deduced from the decrease of the
30 thermophilous low scrubs around AD 1400 and its disappearance from the site around
31 AD 1460 (Fig. 5). Between this date and the end of the 15th century, the Pallars region
32 was significantly depopulated due to local wars, black pest epidemics and climate
33 cooling. This cooling is coeval with the onset of the Little Ice Age (LIA) (Mann et al.
34 2009), which in Europe was cool and wet phase, also recorded in Estanya and other
35 lakes of the Iberian Peninsula (Morellón et al. 2009a). As a result, many villages
36 disappeared and the population concentrated in larger, lowland towns (Marugan and
37 Oliver 2005). This land abandonment is reflected in the pollen diagram in the decline of
38 charcoal and herbaceous crops at the end of zone M-3. Curiously, *Olea*, a lowland
39 element, increases at the same time, attaining its maximum values (~8%). It could be
40 hypothesized that olive cultivation increased in lowlands as a result of the population
41 increase.
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Lithological Unit III was deposited during the Modern Ages, between AD 1460 and 1770, and is characterized by the predominance of Facies 1, indicating deep meromictic lake conditions with high bioproductivity and carbonate precipitation, and reduced terrigenous inputs (Corella et al. this volume). This is consistent with the palynological interpretation of zone M-4, as a period of forest expansion (mainly pines), reduced human activity around the lake (low values of charcoal, and crop and ruderal elements), and decreased terrigenous inputs to the lake (low values of *Pseudoschizaea* and fungi spores in Figure 6). Climate was still cool and probably wet, as it was common during the LIA, a period characterised by moisture fluctuations, as it is shown in the nearby Estanya record (Morellón et al. 2009a). In this case, it is likely that climate was dominant with respect to human activities in the shaping of the Montcortès landscape. A similar situation has been observed in other regional sequences such as Taravilla lake (Moreno et al. 2008), when climate and not human action seems the main cause of hydrological, sedimentological and vegetation changes recorded. The expansion of pastures documented in historical records after the feudalism in the Pyrenees is not recorded in the Montcortès pollen diagram, suggesting that once more, it occurred primarily in lowlands. A slight increase in clastic inputs was recorded at AD 1660, coinciding with the returning to a phase of forest clearance and farming, as suggested by the decline of forests and the increase of charcoal and hemp pollen.

A significant sedimentological shift occurred at AD 1770, close to the end of the Modern epoch (Fig. 3), when turbiditic sedimentation (Facies 4) appeared again and sedimentation rates increased and remained so until the top of the sequence (Unit II). This coincides with pollen zone M-5, when the maximum of hemp cultivation and/or retting is observed (Fig. 4) suggesting a strong human activity in the lake watershed. Historically, this phase corresponds to the pre-capitalist or pre-industrial period, when subsistence agriculture was dominant in the Pallars region due to its geographic isolation. In the lowlands, below Lake Montcortès, crops expanded, whereas in the mountains above the lake, cattle raising and forest exploitation were the dominant practices (Bringué 2005). This situation is clearly reflected in the pollen diagram by the reduction of trees (especially *Pinus*) and the increase of herbaceous crops at the end of zone M-5. Once more, the large amount of hemp pollen contrasts with the lack of documentary evidence on its cultivation or the use of the lake for retting. The large

amounts of *Cannabis* pollen in zone M-5 would be considered to favor the second option, but further work is needed for a definite assessment. Pollen zone M-6 coincides with a crisis of the subsistence agriculture and the significant decline in population of the Pallars region, coeval with the general European agricultural crisis of the mid-19th century (Farràs 2005). Both climate (floods) and agriculture pests have been invoked to explain this phenomenon. Despite the low number of samples in pollen zone M-6, the decline of herbaceous crops and the forest recovery are evident (Figs. 5) and consistent with historical accounts.

Correlations with the Estanya records: vegetation and land use

Before comparing the two records, it should be realized that Lake Montcortès (LM) and Lake Estanya (LE) settings possess different environmental characteristics due to their different altitudes. LE is situated ~50 km SW of LM and 350 m elevation below it. Therefore, LM is in the boundary between the Mediterranean lowlands and the Middle Montane belt, within the Submediterranean bioclimatic domain; whereas LE is in the Mediterranean lowlands, around the transition between the Submediterranean and the Mediterranean bioclimatic regime (Morellón et al. 2009b; Corella et al. in press). The climate around LE is continental Mediterranean, with a mean annual temperature of 14 °C, ranging from 4 (January) to 24 °C (July); and an average annual precipitation of 470 mm, ranging from 18 (July) to 50 mm (October) (Morellón et al. 2009b). At present, the vegetation is a mosaic of shrublands (such as *Buxus*, *Juniperus*, and *Pistacia*), oak forests (*Q. rotundifolia*), pastures, and barley (*Hordeum*) crops. Littoral vegetation around the lake is dominated by *Tamarix*, *Phragmites*, *Juncus*, *Typha* and *Scirpus* (Avila et al. 1984). Two pollen records of the last millennium are available for this lake: one in the deepest part, at 20 m water depth (Morellón et al. 2009b) and another littoral, under 1.5 m water depth (Riera et al. 2004). The first covers the last 800 years, whereas the second accounts for the last two millennia, both with a similar average resolution of around 50 years per sampling interval. The main results of these studies and Lake Montcortès, in terms of vegetation and land use, are depicted in Fig. 7 for comparison.

There is an initial phase of apparently undisturbed woodland around both lakes. In Montcortès, these forests were dominated by *Pinus* while in Estanya, *Quercus* (both deciduous and evergreen) were more abundant, likely because its lower altitude.

Subsequent deforestation took place also in both localities, as recorded by sudden tree-pollen declines and major charcoal peaks, but it occurred first at Estanya (around AD 900; Riera et al. 2004) than at Montcortès, where it was not completed until the end of the pre-feudal period, slightly before AD 1000. Although dating errors could not be disregarded, this northward trend would be consistent with historical records that report a conspicuous migratory trend towards the mountains caused by the northward Muslim expansion (Marugan and Oliver 2005). In this context, it is worth noting that Muslims occupied the region around Estanya between around AD 720 and AD 1030, but they did not reach the Montcortès area (Riera et al. 2004; Esteban 2003). Major deforestations also coincide with the end of shifting cultivation and the beginning of a more or less permanent and extensive land exploitation, which marks the beginning of the feudal epoch. This is manifested in the development of cereal and hemp cultivation in both localities studied. As stated before, occasional *Olea* and *Vitis* crops are mentioned in written documents, thus the scarcity of the first and the absence of the second in the Montcortès diagram is interpreted in terms of inadequate altitude for their cultivation. In Estanya, *Olea* pollen is also very scarce at that time, suggesting wind transport from lower altitudes, but *Vitis* is already present, though in low quantities (Riera et al. 2004).

The warming recorded at Montcortès between AD 1100 and 1350, correlated with the MWP, is paralleled at Estanya by equally warmer and drier climates, manifested in the increase of thermophilic elements in the pollen diagrams (Riera et al. 2004; Morellón et al. 2009a). During this warm and dry episode, fire events reappear at both Montcortès, around AD 1100, and Estanya, between about AD 1250 and 1300 (Morellón et al. 2009a). The first was of medium intensity and is not linked to extensive deforestation, possibly related to pastoralism. The second fire event was deduced from a barren, charcoal-rich layer from the deeper Estanya core (Morellón et al. 2009a), which is absent in the shallow one (Riera et al. 2004), and coincides with a conspicuous forest clearance event. A possible explanation for this discrepancy is that water levels were probably lower during that time (which is consistent with warmer and drier climates) and the shallow coring site was exposed; so charcoal was not deposited in the littoral zone or was removed by erosion. During this phase, agricultural activities did not experience significant changes in the lake catchments, but a vineyard expansion around Estanya is worth mentioning (Morellón et al. 2009a).

1 The LIA cooling initiated around AD 1350 in Montcortès was also identified in
2 Estanya, where several moisture fluctuations occurred until the 19th century. Morellón et
3 al. (2009a) suggested several humidity oscillations, while Riera et al. (2004) proposed
4 the occurrence of a dry episode between about AD 1600 and AD 1750, and two humid
5 phases before and after, respectively (Fig. 7). The LIA onset was characterized, in both
6 lake records, by a decline of cultivated plants suggesting an agriculture depression. This
7 coincides with the so called “low medieval crisis”, characterized by intense
8 depopulation and land abandonment (it is estimated that around one third of the
9 population emigrated and concentrated in larger towns located to the southern lowlands
10 [Marugan and Oliver 2005]) attributed to the combined effect of climate deterioration, a
11 “black death” epidemic occurred in AD 1348, local wars, and probably the collapse of
12 the feudal system (Fig. 3) (Esteban 2003; Marugan and Rapaino 2005). Despite this
13 agricultural crisis, both Montcortès and Estanya records document a progressive
14 increase of *Olea* pollen culminating around AD 1500, which is interpreted as an
15 expansion of olive crops in the region (Riera et al. 2004; Morellón et al. 2009a).
16 According to written documents, an expansion of olive and vineyard crops occurred
17 during this time at the southern lowlands, so it is probable that most of the recorded
18 *Olea* pollen was transported upslope to the lake catchments by wind (Cañellas-Boltà et
19 al. 2009).

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36 The end of feudalism and the beginning of the Modern epoch (AD 1500) is
37 characterized in Montcortès by a forest recovery, also documented in Estanya, where a
38 still cooler but wetter climate has been suggested (Riera et al. 2004). During the 16th
39 century, pollen records show some spatial heterogeneity concerning land use (Fig. 7).
40 This probably due to the contrasting histories of lowlands (used mainly for agriculture)
41 and highlands (mostly dedicated to pastures and cattle raising) since the beginning of
42 the Modern epoch, characterized by land privatization (Bringué 2005). In the 17th
43 century, a dramatic increase of hemp pollen was recorded in Montcortès and in the
44 Estanya shallow core, coinciding with a dry episode and another forest clearance event
45 (Riera et al. 2004). As noted before, the lack of written documentation about hemp
46 cultivation around Montcortès prevents to know if the *Cannabis* pollen peak may be due
47 to cultivation and/or retting, or both. High amounts of hemp pollen (10-80%) together
48 with the presence of a suitable environment are considered to be indicative of retting
49 (Mercuri et al. 2002). In Montcortès, the suitable environment is provided by the lake
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and the hemp pollen is around 40-60% at that time, which strongly suggests retting. Anyway, it seems that hemp industry was a predominant activity in the region at that time. However, the hemp peak was not recorded in the Estanya deeper core, where this pollen type is comparatively less abundant and does not experience significant changes in time (Morellón et al. 2009a). Sedimentological and taphonomical factors cannot be disregarded but more records are still needed for a satisfactory explanation. The post-Modern crisis of subsistence agriculture and the coeval forest recovery, were also documented in the three available pollen records.

General conclusions and final remarks

After a detailed comparison of the Montcortès and Estanya records for the last millennium, it seems clear that human activities have played a major role in vegetation change history and have been crucial for the shaping of southern Pyrenean landscapes (Ejarque et al. 2009; Pèlachs et al. 2009a, b). The more intense climatic shifts (i. e. the MWP and the LIA) are still recognizable through particular plant associations, as for example the low Mediterranean scrubs in Montcortès; however, minor oscillations are likely obscured by the consequences of human land use. This is especially true for moisture changes, as the appearance or disappearance of plant communities commonly used as humidity proxies –as for example forests– have been mainly determined by human activities rather than by climatic reasons. As it is well known, biological records usually need independent evidence to interpret them in terms of climatic or anthropic factors, avoiding circular reasoning (e.g. Leroy 2010). Fortunately, in the case of Montcortès and Estanya lakes, abundant independent evidence is available from both sedimentological and documentary perspectives, which greatly strengthens paleoecological interpretation. In this study, both lines of independent evidence have been used to try to disentangle the potential forcing factors behind the vegetation changes recorded by pollen analyses. In general, vegetation patterns in time and human land-use changes have shown a high degree of coincidence but climate-humans-landscape relationships are complex and there is still room for improvement in our interpretations. For example, climate has not only a direct influence of natural communities but also an indirect one by affecting human practices, thus modifying the anthropic element of change. It would be interesting to further investigate this aspect. Another interesting point, which is already in progress, is the detailed paleolimnological

1 reconstruction using aquatic proxies combined with sedimentological and physico-
2 chemical evidence (Scussolini et al., in prep). Montcortès is especially suited for such
3 studies because its annually laminated sediments (Corella et al. in press). Also, the
4 comparison of pollen and aquatic remains records would be useful to know the
5 magnitude and time lag of the response of the involved organisms to environmental
6 changes and, in this way, to evaluate their potential usefulness as paleoclimatic
7 indicators.
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Table 1 Vegetation types present in the Montcortés lake area, following Carreras et al. (2005-2006) and using the CORINE system (CBM 1991).

CORINE UNITS	VEGETATION TYPES	MAIN COMPONENTS	ENVIRONMENTAL OBSERVATIONS	CLIMATE
<i>Unit 42v</i>	Submontane conifer forests	<i>Pinus nigra salzmanii</i> , poorly developed understory	diverse substrates	moderately wet
<i>Unit 45f</i>	Evergreen oak forests	<i>Quercus rotundifolia</i> , well developed understory with abundant sclerophyllous shrubs (<i>Quercus coccifera</i> , <i>Rhamnus alaternus</i> , <i>Rhamnus saxatilis</i> , <i>Prunus spinosa</i> , <i>Buxus sempervirens</i> , <i>Lonicera etrusca</i>). Herbaceous stratum with nemoral (shadow-adapted) species: <i>Rubia peregrina</i> , <i>Teucrium chamaedrys</i> , <i>Asparagus acutifolius</i> and <i>Brachypodium retusum</i> .	well exposed sites of lowland plains, hills and slopes	continental Mediterranean
<i>Unit 41k</i>	Deciduous oak forests	<i>Quercus pubescens</i> , sometimes with <i>Pinus sylvestris</i> or, less frequently, <i>Fagus sylvatica</i> or <i>Tilia cordata</i> . Other trees: <i>Quercus cerritoides</i> , <i>Quercus subpyrenaica</i> , <i>Acer monspessulanum</i> , <i>Acer opalus</i> subsp. <i>Opalus</i> ; Shrubs: <i>Buxus sempervirens</i> , <i>Coronilla emerus</i> , <i>Amelanchier ovalis</i> , <i>Colutea arborescens</i> , <i>Cytisophyllum sessilifolium</i> , <i>Viburnum lantana</i> ; Herbs: <i>Primula veris</i> subsp. <i>Columnae</i> , <i>Hepatica nobilis</i> , <i>Brachypodium phoenicoides</i> , <i>Campanula persicifolia</i> .	well exposed warm slopes, submontane stage, up to 1500 m a.s.l., calcareous substrates	Submediterranean or mountain Mediterranean
<i>Unit 24a</i>	Non-forested river margins	Pioneer plants frequent. Trees: <i>Salix</i> spp.; Shrubs: <i>Myricaria germanica</i> ; Herbs: <i>Mentha longifolia</i> , <i>Juncus</i> spp., <i>Saponaria officinalis</i> , <i>Polygonum</i> spp., <i>Bidens</i> spp., <i>Xanthium italicum</i> , <i>Agrostis stolonifera</i> , <i>Polypogon viridis</i> , <i>Paspalum distinctum</i> , <i>Andryala ragusina</i> , <i>Glaucium flavum</i> .	azonal vegetation type due to natural or anthropic causes, present from the lowlands to the subalpine zone	
<i>Unit 31q</i>	Shrubland	<i>Amelanchier ovalis</i> , <i>Buxus sempervirens</i> and <i>Rhamnus saxatilis</i> . Up to 2 m high, with	well exposed rocky	Submediterranean

		an irregular herbaceous stratum of heliophyllous plants. Other shrubs: <i>Coronilla emerus</i> , <i>Ligustrum vulgare</i> , <i>Cytisophyllum sessilifolium</i> ; Herbs: <i>Tanacetum corymbosum</i> , <i>Viola alba</i> , <i>Teucrium chamaedrys</i> , <i>Carex halleriana</i> .	slopes, on calcareous terrains. Secondary colonizer where the forest has been removed	
<i>Unit 32 ab</i>	Shrubland	<i>Arctostaphylos uva-ursi</i> , with some emergent shrubs (<i>Buxus</i>) or trees (<i>Pinus</i> , <i>Quercus</i>). Other shrubs: <i>Juniperus communis</i> , <i>Amelanchier ovalis</i> ; Herbs: <i>Primula veris</i> subsp. <i>Columnae</i> , <i>Lavandula angustifolia</i> subsp. <i>Pyrenaica</i> , <i>Hepatica nobilis</i> , <i>Cruciata glabra</i> , <i>Carex humilis</i> , <i>Avenula pratensis</i> subsp. <i>Iberica</i> .	deforested areas and forest clearings on calcareous soils	Submediterranean
<i>Unit 34n</i>	Meadows	<i>Aphyllanthes monspelliensis</i> . Shrubs: <i>Salvia lavandulifolia</i> , <i>Santolina chamaecyparissus</i> , <i>Ononis natrix</i> , <i>Teurcium polium</i> , <i>Genista scorpius</i> , <i>Helianthemum italicum</i> , <i>Euphorbia nicaensis</i> , <i>Thymus vulgaris</i> ; Herbs: <i>Avenula pratensis</i> subsp. <i>Iberica</i> , <i>Koeleria vallesiana</i> , <i>Brachypodium phoenicoides</i> , <i>Leuzea conifera</i> , <i>Linum narbonense</i> , <i>Carduncellus monspeliensis</i> , <i>Catananchae caerulea</i> , <i>Asphodelus cerasiferus</i> , <i>Asperula cynanchica</i> .	light pastures from 100 to 1400 m a.s.l., on calcareous, generally unstable terrains	Mediterranean to Submediterranean
<i>Unit 38b</i>	Mowing meadows	<i>Arrhenaterum elatius</i> . Other herbs: <i>Trisetum flavescens</i> , <i>Trifolium pretensis</i> , <i>Dactylis glomerata</i> , <i>Rhinanthus mediterraneus</i> , <i>Geniana lutea</i> , <i>Plantago lanceolata</i> , <i>Lotus corniculatus</i> , <i>Daucus carota</i> , <i>Rumex acetosa</i> .	plains and gentle slopes with deep and eutrophic soils	Submediterranean
<i>Unit 82a</i>	Extensive herbaceous crops	Cereals, rarely forages, with abundant weeds. Herbs: <i>Hordeum</i> sp., <i>Avena sativa</i> , <i>Triticum</i> sp., <i>Secale cereale</i> , <i>Medicago sativa</i> ; Weeds: <i>Lolium rigidum</i> , <i>Papaver rhoeas</i> , <i>Bromus</i> sp., <i>Polygonum aviculare</i> .	open valleys and gentle slopes, on clayey deep soils	Mediterranean
<i>Unit 87a</i>	Abandoned crops	Pasture-like terrains colonized by shrubs and ruderal species like <i>Inula viscosa</i> and <i>Oryzopsis miliacea</i> .	lowlands to medium elevations, on a wide range of soils	wide range of climate conditions

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Unit 61f	Badlands	Devoid of vegetation or with few scattered plants like <i>Achnatherum calamagrostis</i> , <i>Brachypodium phoenicoides</i> , <i>Erucastrum nasturtifolium</i> , <i>Tussilago farfara</i> , <i>Genista scorpius</i> , <i>Thymus vulgaris</i> or <i>Lithospermum fruticosum</i>	highly eroded terrains, generally on steep slopes	
Unit 86a	Urban areas	diverse ruderal and cultivated flora		

Table 2 Pollen groups according to the present-day vegetation types, as defined by Carreras et al. (2005-2006), with additional information from Muenscher (1980), Folch (1981), Vigo (1983, 1996, 2008), de Bolòs et al. (1990), and de Bolòs (2001).

VEGETATION GROUPS	POLLEN TAXA
Conifer forests	<i>Pinus, Abies</i>
Evergreen oak forests	<i>Quercus</i> -evergreen type
Deciduous oak forests	<i>Quercus</i> -deciduous type, <i>Tilia, Acer, Betula, Cornus, Corylus, Fagus</i>
Riverine forests	<i>Alnus, Populus, Salix, Ulmus</i>
Cultivated trees	<i>Cupressus, Juglans, Olea, Prunus avium</i> -type
Shrubs from both shrublands and the understory of light oak forests	<i>Buxus, Ericaceae, Juniperus, Phillyrea, Rhamnus</i>
Low Scrubs	<i>Rosmarinus</i> -type, <i>Hedysarum</i> -type, <i>Helianthemum, Ephedra</i>
Meadows and/or pastures	<i>Asphodelus albus/ramosus</i> -type, <i>Linum, Plantago, Poaceae, Sideritis, Pedicularis</i>
Herbaceous Crops	<i>Cannabis</i> -type, <i>Triticum/Avena</i> -type, other <i>Fabaceae, Secale</i>
Ruderal plants and/or weeds	<i>Artemisia, Centaurea, Chenopodium</i> -type, <i>Convolvulus, Echium, Polygonum aviculare</i> -type, <i>Rumex</i>
Aquatic plants	<i>Cladium</i> , other <i>Cyperaceae, Myriophyllum, Pinguicula, Potamogeton, Typha/Sparganium, Utricularia</i>
Others or taxa occurring in several vegetation types or with insufficient taxonomic resolution	<i>Ambrosia</i> -type, <i>Androsace, Apiaceae, Asteraceae, Bupleurum</i> -type, Psilate tricolporate (~ <i>Castanea</i>), <i>Cerastium</i> -type, <i>Draba</i> -type, <i>Gentiana</i> -type, <i>Geranium, Paronychia</i> -type, <i>Ranunculus, Sanguisorba</i> minor-type, <i>Valeriana</i>

Figure captions

Figure 1. Location map of Lakes Montcortès and Estanya. The Pallars region is highlighted (grey area).

Figure 2. Vegetation map according to the classification of Carreras et al. (2005-2006), using the CORINE system (CBM 1991). L = Lake Montcortès.

Figure 3. Schematic account of the main historical periods and events in the Pallars region, especially those related to economic models and land use. Summary after Esteban (2003), Marugan and Oliver (2005), Bringué (2005), and Farràs (2005). LM = Lake Montcortès, N = North, S = South.

Figure 4. Percentage sporomorph diagram, including the total pollen and charcoal influx. Elements included in the pollen sum: CF – conifer forests, EOF – evergreen oak forests, DOF – deciduous oak forests, RF – riverine forests, CT – cultivated trees, S – shrublands, MS – Low Mediterranean scrubs, M/P – meadows/pastures, HC – herbaceous crops, R/W – ruderal/weeds. Elements outside the pollen sum: A – aquatic plants, F – ferns. Zonation was made using the Optimum Splitting by Information Content (OSIC) and the broken stick methods (Bennett 1996). Solid lines indicate x10 exaggeration.

Figure 5. Percentage pollen diagram grouped by vegetation types, according to Carreras et al. (2005-2006), with additional information from Muenscher (1980), Folch (1981), Vigo (1983, 1996, 2008), de Bolòs et al. (1990), and de Bolòs (2001). The zonation is based on percentage diagram (Fig. 3).

Figure 6. Percentage diagram for non-pollen palynomorphs (NPP) with respect to the pollen sum. The scales of *Tetraedron* and *Botryococcus* have been reduced for more clarity. Solid lines indicate x10 exaggeration. Zonation as in Figure 3.

Figure 7. Interpretation of Montcortès pollen diagram and correlations with the existing Estanya records. LIA = Little Ice Age, MWP = Medieval Warm Period.

Figure 1

Figure 2

Figure 3

Lake Montcortès (core MON-04-1A-1K)
Analyst: V. Rull

Figure 5

Lake Montcortès (core MON-04-1A-1K)
Analyst: V. Rull

Figure 6

Figure 7