

Textural characteristics and facies of sand-rich contourite depositional systems

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ABSTRACT

This work presents a detailed study of CONTOURIBER and Integrated Ocean Drilling Program 339 sediment data targeting sand-rich contourites in the Eastern Gulf of Cadiz. All of the collected sediments are interpreted as contourites (deposited or reworked by bottom currents) on the basis of oceanographic setting, seismic and morphometric features and facies characteristics. A variety of sandy and associated facies are found across the study area including: (i) bioturbated muddy contourites; (ii) mottled silty contourites; (iii) very fine mottled and fine-grained bioturbated sandy contourites; (iv) massive and laminated sandy contourites; and (v) coarse sandy/gravel contourites. The thickest sands occur within contourite channels and there is a marked reduction in sand content laterally away from channels. Complementary to the facies descriptions, grain-size analysis of 675 samples reveal distinctive trends in textural properties linked to depositional processes under the action of bottom currents. The finest muddy contourites (<20 µm) show normal grain-size

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distributions, poor to very poor sorting, and zero or low skewness. These are deposited by settling from weak bottom currents with a fine suspension load. Muddy to fine sandy contourites (20 to 200 μm) trend towards better sorting and initially finer and then coarser skew. These are typical depositional trends for contourites. As current velocity and carrying capacity increase, more of the finest fraction remains in suspension and bedload transport becomes more important. Clean sandy contourites ($>200 \mu\text{m}$) are better sorted. They result from the action of dominant bedload transport and winnowing at high current speeds. The results highlight the importance of bottom current velocity, sediment supply and bioturbational mixing in controlling contourite facies. Despite growing interest in their hydrocarbon exploration potential, contourite sands have remained poorly understood. This research therefore has important implications for developing current understanding of these deposits and aiding the correct interpretation of deep marine sands and depositional processes.

Keywords Contourites, deep-water sands, grain size, Gulf of Cadiz, sediment facies model.

INTRODUCTION

Contourites are those sediments deposited or substantially reworked by bottom currents (Stow et al., 2002). In areas with long-term deposition, they can construct large mounded drifts (contourite drifts) in a variety of shapes and sizes, which may be elongated up to several hundreds of kilometres along the continental margin (Rebesco et al. 2014; Esentia et al. 2017). The characteristics of the fine-grained sediments collected from these mounded drifts are now well-constrained thanks to early facies and textural studies (Stow & Lovell, 1979; Gonthier et al., 1984; Stow & Piper, 1984). More recently, the capability of bottom currents to deposit significant accumulations of sand-rich facies has been recognized. The characteristics of these sandy contourites are still poorly understood and, therefore, can be difficult to distinguish from other deep water sands, such as turbidites.

The eastern Gulf of Cadiz (Fig. 1) contains one of the largest modern sand-rich contourite depositional systems in the world. It covers an estimated area of 4000 km² and industry boreholes suggest that there is over 800 m of sand-rich contourite section (Buitrago et al. 2001; Antich et al. 2005). It therefore provides an ideal location for the study of contourite sands, their characteristics and controlling factors. These contourite sands form part of a much larger contourite depositional system in the Gulf of Cadiz (e.g. Llave et al., 2007, fig. 1; Hernández-Molina et al., 2016), for which a huge database exists (surface and subsurface). This database has been expanded with targeted data

acquisition over the sand-dominated region in the east of the Gulf of Cadiz, which is the focus of this study.

This sediment dataset is one of the largest for any sand-dominated contourite system acquired to date. This study develops: (i) the key features of the contourite facies recovered from the eastern Gulf of Cadiz; (ii) their textural characteristics and key insights on contourite depositional processes; and (iii) a new sedimentary facies model for contourite sands. Better characterization of these deposits in modern settings will aid their identification in the subsurface and allow their reservoir and hydrocarbon play potential to be evaluated (Viana & Rebesco 2007).

CONTOURITES: CURRENT KNOWLEDGE

Contourites are defined as any sediment deposited or substantially reworked by the persistent action of bottom currents (Stow et al., 2002; Rebesco et al., 2014). These bottom currents are capable of eroding, transporting and depositing sediments along the sea floor (Rebesco & Camerlenghi, 2008). Thermohaline-driven bottom currents, such as those encountered in this study, generally operate most effectively in deep water (i.e. off-shelf). They may also be present on the continental shelf, where they compete with other processes, such as wind- or tidal-driven circulation, which may have an even greater effect on sediment deposition. Thermohaline-driven currents generally exhibit persistent along-slope flow (following the local bathymetry) but can vary in direction and velocity due to the influence of other oceanographic processes (e.g. Shanmugam, 2012, 2013; Rebesco et al., 2014). Such processes form contourite facies in deep-water settings (Shanmugam, 2012; Stow et al., 2013a,b), and represent an entirely different deep-water deposit from turbidite and pelagic sediments (Viana, 2008; Stow et al., 2013b).

Fine-grained contourite facies (muds, silts and muddy sands) have been well-studied in previous works and their principal features are fairly well-constrained (Stow & Lovell, 1979; Gonthier et al., 1984; Stow & Piper, 1984; Alonso et al., 2016). They are distinguished from other fine-grained deposits (pelagic and turbiditic) on account of their bi-gradational sequences, distinctive uniform composition and textural characteristics, and their absence of primary sedimentary structures due to intense bioturbation. The now widely-recognized standard contourite facies model is based on an extensive compilation of data from modern and sub-recent sediments worldwide. The elements of this model, first developed by Stow (1977, 1982), were combined into the contourite sequence by Gonthier et al. (1984), Stow & Piper (1984) and Faugères et al. (1984), with the later addition of the

C1 to C5 notation (Stow et al., 2002; Stow & Faugères, 2008). The bi-gradational grading represents continuous sedimentation over an interval of several millennia and has been interpreted to be the result of the interaction of three independent controls (Stow et al. 1986): (i) fluctuation in mean bottom current velocity, with the coarsest division, C3, representing the maximum velocity; (ii) variation in lateral sediment supply, principally from turbidity currents; and (iii) variation in vertical sediment supply from primary biogenic productivity in surface waters. The complete contourite sequence is not always present, but is typically found as partial contourite sequences (Stow & Faugères, 2008).

Contourite *sands*, especially the clean medium and coarse sands, are much less well-known and are also, therefore, much more difficult to distinguish from other deep-water sands. Typically, most contourite sands have been assigned to the C3 division of the standard contourite facies model, although Stow & Faugères (2008) did recognize the need for a different facies model specifically for sand-rich contourites. Viana et al. (1998) had previously described ‘bottom-current controlled sand deposits’ from a variety of deep-water environments, but where they were uncertain of classifying the sediments as contourites. More recent work has elucidated the role of other processes, such as wind-driven currents, internal tides, internal waves, benthic storms and tsunami-generated bottom currents in depositing very similar sandy facies (Rebesco et al., 2014), as well as those deposited by mixed alongslope and downslope process interaction.

Stanley (1993) introduced the concept of bottom-current winnowed turbidite deposits. These occur where a downslope turbidite channel-levée system is formed under a bottom current. The current can pirate fines from the turbidity current overspill plume and redistribute them alongslope (Shanmugam, 2006). Evidence for interaction may also be seen at a facies-scale as modified turbidites (Stanley, 1993; Mulder et al., 2008). Examples of such deposits have been interpreted from the Columbia Channel offshore Brazil (Faugères et al., 2002) and Kazusa Forearc, Japan (Ito, 1996). Mixed downslope–alongslope and systems influenced by internal tides have been described in several cases (Shanmugam et al., 1995; Shanmugam, 2006, 2012; Youbin et al., 2008, 2011) although there is still considerable debate over what features accurately reflect this process interaction.

At higher bottom current velocities, the complete removal and redistribution of incoming downslope sediment, as well as widespread erosion, may occur as is seen along the Pliocene of the Algarve margin of the Gulf of Cadiz (Brackenridge et al., 2013a). The pirated and eroded sediment will be re-deposited elsewhere in the basin where the transport capacity of the current is diminished. According to most accepted definitions, these are true contourites. No comprehensive sediment facies and textural study has been completed to date in a setting that is contourite dominated such as the Gulf of Cadiz.

GULF OF CADIZ: GEOLOGICAL AND OCEANOGRAPHIC SETTING

Geological setting

The Gulf of Cadiz has had a complex and dynamic tectonic history as a result of its location on the boundary between the African and Eurasian continental plates. Three key events have been identified as crucial geological factors responsible for the modern set-up of the margin (Maldonado et al., 1999; Hernández-Molina et al., 2006): (i) Late Mesozoic passive margin development related to rifting, sea floor spreading and the opening of the Tethys and Atlantic Oceans; (ii) Cenozoic convergence of the African and Eurasian plates in a north–south direction; and (iii) foredeep formation relating to Betic–Rifean orogenesis east of the Gulf of Cadiz. Today, a compressional tectonic regime coupled with oblique WNW–ESE motion gives an apparently slowing rate of convergence and a move to tectonic stability (Rosenbaum et al., 2002).

The continental margin of the Gulf of Cadiz is broadly arcuate, *ca* 350 km long, and runs from the Strait of Gibraltar to the south-western tip of Portugal (Cape St Vincent; Fig. 1). The shelf break is at a water depth of 120 to 140 m and the continental slope joins the abyssal plain at depths of over 4300 m. Morphological features of the continental slope include both downslope and alongslope depositional and erosional elements (Hernández-Molina et al., 2006; Medialdea et al., 2009). The emplacement of an unstable allocthonous unit in the Miocene has led to intense mud-diapirism and halokinesis along the margin. The resulting mud volcanoes, diapiric ridges and adjacent depocentres modify the modern day oceanic conditions (Somoza et al., 2003; Hernández-Molina et al., 2006; Llave et al., 2007; Tasiánas 2010).

Oceanographic setting and impact on sedimentary processes

The Gulf of Cadiz is dominated by the water exchange between the Atlantic Ocean and Mediterranean Sea. The Strait of Gibraltar allows for the exchange between the warm saline Mediterranean Outflow Water (MOW) and the overlying Atlantic Inflow Water (AIW) (Howe, 1982; Baringer & Price, 1999; Lobo et al., 2000; Ambar et al., 2002; Cabecadas et al., 2002; Garcia-Lafuente et al., 2006; Llave et al., 2007). The MOW is composed of a mixture of water masses sourced from the Mediterranean Basin (Millot, 2009). It accelerates through the narrow Strait of Gibraltar – locally reaching velocities up to 300 cm sec^{-1} (Ambar & Howe, 1979; Mulder et al., 2003) as it cascades down the shallow sill within the Strait. It forms a turbulent density-driven flow before veering towards the north-west along the mid-continental slope of the Gulf of Cadiz in multiple distinct cores of differing properties (Fig. 1). Density-driven descent (Legg et al., 2009) and mixing with overlying Atlantic waters results in decreasing salinity along the margin from south-east to north-west (Baringer & Price, 1997). Eventually the MOW leaves the sea bed and reaches a neutral buoyancy at a depth of 1200 to 1400 m offshore of Cape St Vincent, where it begins to raft above the North Atlantic Deep Water (Zenk, 1975; Ambar et al., 1999).

Mediterranean Outflow Water-influenced contourite sedimentation in the Gulf of Cadiz is estimated to have begun in the early Pliocene (van der Schee et al., 2016). Hernández-Molina et al. (2003, 2006) and Llave et al. (2007) have identified five distinct ‘morphological sectors’ (Fig. 1) that have developed in direct response to the broadly diminishing bottom current velocities from the proximal sector close to the Strait of Gibraltar in the east, to the canyons sector in the west where bottom current velocities are so reduced that downslope processes begin to dominate. Sector 1, the most proximal to the Strait of Gibraltar, is the area of interest in this study on account of the abundant contourite sands deposited there.

This study uses the terminology and mapped extent of morphological features as described in detail by (Hernández-Molina et al., 2014). Proximal to the Strait of Gibraltar, the turbulent cascade of the MOW over the Strait of Gibraltar (Legg et al., 2009) has resulted in two large and deep erosional channels with areas of rocky outcrop (Hernández-Molina et al., 2014) (Fig. 2). These channels open onto broad erosional terraces (over 2000 km^2 in area) that form a major abraded surface between 500 m and 800 m water depth. This surface is dominated by erosional processes; seismic reflections show truncation and aligned features that are V-shaped in cross-section, interpreted as erosive scours or giant furrows (Hernández-Molina et al., 2006, fig. 9). Erosive features pass to gravel and sand lag deposits and subsequently give way to sand ribbon and wave deposits. Plastered drifts are found along

the upper slope, and a mounded drift has developed at the north-western limit of the erosional terraces (Fig. 2) (Hernández-Molina et al., 2014, fig. 2C). Moving downslope, a large contourite drift separates the erosional terraces from the Lower Continental Slope (Mulder et al., 2003; Hernández-Molina et al., 2014). The lower slope consists of wave fields, channels and depositional lobes (Habgood et al., 2003). This distribution of features represents the reduction of MOW velocity (and thus sediment transport capacity) moving away from the Strait of Gibraltar to the north-west as well as laterally away from the main MOW flow (Mulder et al., 2003; Llave et al., 2006, 2007).

Some work has been carried out on the adjacent shallow water sediments on the continental shelf (Nelson et al., 1993, 1999; Maldonado et al., 2003; Lobo et al., 2010), the adjacent channel and sediment waves sectors (Akhmetzhanov et al., 2007; Hanquiez et al., 2007), as well as within the Strait of Gibraltar (Kelling & Stanley, 1972; Esteras et al., 2000). Focused studies on the region most proximal to the Strait of Gibraltar, however, remain limited to a very few. These include: two industry studies (Buitrago et al., 2001; Antich et al., 2005); some superficial and remotely-sensed studies (Kenyon & Belderson, 1973; Habgood et al., 2003; Mulder et al., 2003), and some recent publications relating to the data used in this study (Alejo et al., 2012; Brackenridge et al., 2013b; Hernández-Molina et al., 2014).

DATA AND METHODOLOGY

Data

This study utilizes core data of differing vintages (Table 1; Fig. 2). Three piston cores, two trigger cores and one box core were acquired on the Scientific Cruise CONTOURIBER-1 on the *R/V Sarmiento de Gamboa*, October 2010. The three 12 cm diameter piston cores (PC06, PC04 and PC08) were divided into sections approximately 1 m in length for ease of transportation. The box core (BC05) was sub-cored and sliced at a resolution of 1 cm for sampling. In total, *ca* 16 m of gravity cores were retrieved.

Six sites were drilled across the Gulf of Cadiz during Integrated Ocean Drilling Program (IODP) Expedition 339, which sailed from 16 November 2011 to 16 January 2012, and in total 5.5 km of cores were recovered (Expedition 339 Scientists, 2012; Stow et al., 2013a). These sites targeted a variety of depositional environments across the Gulf of Cadiz, and the key site of interest in this study is U1388, located in the east (Fig. 2). This site penetrated 253 m and recovered a total of 121 m of

core (Expedition 339 Scientists, 2012). This study focuses on the top 6 m of core U1388, which was logged in detail and sampled at the MARUM Core Repository, University of Bremen. A detailed morpho-sedimentary study of the area where the cores were collected has been carried out by Hernández-Molina et al. (2014). The defined erosional and depositional features described in that study are used to describe the location and depositional environment of the cores studied here.

Core logging

Visual description was carried out on all cores (Table 1). This included conventional sediment logging, core photography and computed tomography (CT) imaging where available. The top 6 m of IODP core U1388 was completed using visual core description. The three piston cores and their respective trigger cores collected on the CONTOURIBER-1 cruise (PC04, PC06 and PC08) were selected for high-resolution photography immediately after slabbing using an Avaatech XRF Core Scanner (Avaatech BV, Alkmaar, The Netherlands) located at the Universitat de Barcelona Faculty of Geology. In addition, CT-scanning was carried out at the Facultade de Veteranaria at the Universidad de Santiago de Compostela, Lugo (Spain) (Fig. 3) following the procedure published by Mena et al. (2015). Each core was scanned using a high-resolution Hitachi ECLOS medical scanner (Hitachi, Tokyo, Japan). The data was displayed in the freeware MRIcro (Rorden & Brett, 2005) where it can be manipulated in terms of colour, contrast and slice location in order to enhance any density contrasts within the sediment. Figure 3 shows the entire CT-scanned core using the 'Rainramp' colour bar with a contrast of 800 and a brightness of 1130 in order to optimize the density contrasts within the cores. The resulting image has a resolution of $0.2 \times 0.2 \times 0.625 \text{ mm}^3$. This hugely improves the ability to log broad grain-size trends, erosional boundaries and sedimentary features. In particular, burrows are more clearly seen, an important diagnostic criterion of contourite facies. Visual core description was backed up by smear slide analysis (sampling locations indicated in Fig. 3).

Textural analysis

The CONTOURIBER piston and trigger cores were sampled at a regular 2 cm resolution. An additional 33 samples were collected from IODP core U1388 (Table 1; Fig. 3). A total of 675 samples were analysed for grain-size at the University of Bordeaux using laser counting methods on a Malvern Particle Size Analyser machine (Malvern Instruments Limited, Malvern, UK) using standard procedures (Lewis & McConchie, 1994). Laser counters measure grain sizes from $0.05 \mu\text{m}$ to $700 \mu\text{m}$ – or clay up to medium/coarse sand. Since BC05 showed coarse to very coarse sediment content, the 23 samples from the box core were analysed using sieving methods at Heriot-Watt University.

The standard procedure for grain-size analysis in laser counters was used as outlined by McCave et al. (1986), Cooper (1998) and Martins (2003). Additional grain-size sieving methods were used on samples too coarse for the laser counters (Lewis & McConchie, 1994). Distributions are given in a geometric (volume) scaling rather than arithmetic (number) scale to ensure that there is equal emphasis on changes in clay, silt and sand content in the histogram (Blott & Pye, 2001). All analysis was carried out in the software GRADISTAT (Blott & Pye, 2001) using the geometric graphical method as laid out by Folk & Ward (1957) to ensure that each set of analysis can be directly compared to others in the study. Key statistical measures include: mean, sorting (standard deviation), skewness and kurtosis. The formulae used for these measurements and the cut-off values for defining the sorting, skewness and kurtosis are outlined in Blott & Pye (2001). Cumulative frequency curves were created to clearly visualize differences in distribution. After calculation of the essential statistics, data were compared using cross-plots (Folk, 1964; Martins, 2003). The scale and sediment classes as defined by Wentworth (1922) were used.

Sediment dating

In total, 13 samples have been dated using radiocarbon methods: three each from PC06, PC08 and U1388, and four from PC04 (Table 2; Fig. 3). Coarse-grained BC05 was thought to consist of large amounts of reworked material, therefore being unsuitable to allow for accurate dating and was excluded from the analysis. Planktonic foraminiferal species of *Globigerina bulloides* and *Globigerinoides ruber* were targeted for the dating. This was carried out at the Scottish Universities Environmental Research Centre (SUERC) Radiocarbon Dating Laboratory in the UK and the Laboratoire de Mesure du Carbone 14 (France) using accelerator mass spectrometry (AMS) methods. All radiocarbon ages were calibrated using the University of Oxford Radiocarbon Accelerator Unit Calibration Program OxCal4 (using the Marine13 calibration curve).

RESULTS

Sediment facies and facies sequences

Core logging

Logging of the seven cores demonstrates varying facies characteristics across each core site (Fig. 3). Most proximally to the Strait of Gibraltar, BC05 is located within the Southern Channel that contains the Mediterranean Outflow as it sweeps around the Gulf of Cadiz continental slope (Fig. 2). The

sediments collected from this box core are the coarsest in the study, ranging from medium to very coarse sand. Smear slide analysis reveals a mixed composition of predominantly quartz with bioclastic material. The largest grain sizes are made up of shelly debris. There is evidence of lamination, but bioturbation is not seen.

Adjacent to the Southern Channel, PC04 was collected from a mounded contourite drift at 658 m water depth. Facies range from silt to fine sand, and seven units are identified separated by both abrupt and gradational boundaries (Fig. 3). A thin 30 cm tan-brown unit tops the succession (unit 1). On analysis of the associated trigger core PC04TG, this unit reaches 35 cm in thickness (Fig. 3). Unit 2 is the thickest at 2.53 m thick. It is composed of very poorly sorted sandy mud. Within this unit there are multiple gradational changes towards sandy silts and occasional fine sand lenses. Subtle lamination is seen in the upper 20 cm, but in general the unit is highly bioturbated and mottled in appearance. The sediment is of mixed composition with a significant foraminifera portion, in addition to quartz and other terrigenous clasts. The underlying units 3, 4 and 5 are much thinner: 0.25 m, 0.48 m and 1.1 m respectively. Unit 3 appears to down cut into the underlying unit and is made up of highly bioturbated fine sand. The underlying unit shows a subtle fining upward sequence of bioturbated sandy silts and muds. An abrupt surface separates this unit from the underlying unit 5 which has a broadly coarsening upward trend consisting of fine sandy muds with sections of higher portions of sand. There is subtle cross-lamination in this unit with larger scale bioturbation in the sandier sections (Fig. 3). Unit 6 is separated from the overlying unit by a sharp surface and consists of silty mud grading upward to a very poorly sorted silty fine sand. Burrows are seen penetrating downwards from the upper boundary of this unit and large (centimetre-scale) burrows are observed in the sandier section (Fig. 3). There is poor core recovery within this sandy section, but it appears to be over 1 m in thickness with reverse grading. Muddy silt at the base of the unit is heavily bioturbated and mottled in appearance. The lowermost unit 7 is a bioturbated silty sand and has a similar composition to the overlying units with significant quartz and nannofossil portions. In the sandy sections, foraminifera are relatively lacking, and in the lower sandy section, quartz grains are iron-stained.

Collected from the same mounded drift, PC08 is located more distally to the north-west, and deeper on the continental slope than PC04, at 961 m water depth (Fig. 2). Despite being the finest-grained of the sediment cores, PC08, still shows six distinct sedimentary units separated by sharp boundaries (Fig. 3). Unit 1 is only 6 cm thick in PC08, but reaches 43 cm in the trigger core PC08TG

(Fig. 3). It is composed of muddy silt with a dominantly bioclastic composition (nannofossils, foraminifera, siliceous sponge fragments and bryzoa). As with unit 1 in PC04, it has a distinct tan colour in core and is highly bioturbated. This is separated from the underlying unit 2 by an abrupt (but with no evidence for being erosive) boundary. Unit 2 is approximately 2.25 m thick and is a poorly sorted muddy silt with localized cleaner silts. There is some minor coarsening upward in the top 20 cm of the unit. It is of mixed silici-bioclastic composition with notable quartz, foraminifera and nannofossil (coccolith) portions. Units 3 and 4 show coarser sediment accumulation with very poorly to poorly sorted muddy sands. The boundary between units 3 and 4 appears abrupt and down cutting, with unit 4 being topped by muddy sand, and grading down the core back to muddy silt. Quartz tends to make up the coarsest portion of the sediment and is commonly iron-stained. At the base of unit 4 at 5.1 m below sea floor (bsf), another abrupt grain-size change is identified, with the underlying unit 5 comprising a coarsening upward succession from muddy silt to muddy fine sand. The lowermost unit 6 is also the coarsest in PC08. It is a very poorly sorted muddy fine sand of mixed composition separated from the overlying unit 5 by a sharp boundary. The entire core is very heavily bioturbated with small-scale burrows (Fig. 3).

Core PC06 is located on the upper continental slope at 490 m water depth. It is adjacent to the northern channel defined by Hernández-Molina et al. (2014) and collected from a plastered contourite drift (Fig. 2). It consists of three sedimentary units that are separated by sharp contacts (Fig. 3). Unit 1 is 40 cm thick and consists of moderately sorted fine sand. Compositionally, this unit is dominated by terrigenous sediment (quartz and feldspars), with a minor component of shell fragments and foraminifera. The unit is devoid of sedimentary structures and only rare large-scale bioturbation is observed. There is some evidence of fluidized deformation. This unit differs greatly from the underlying unit 2 – a very poorly-sorted muddy sand. Unit 2 maintains a mixed composition, but has a much larger bioclastic component, mainly siliceous sponge fragments and foraminifera. It is heavily bioturbated and shows complex grading patterns. The underlying unit 3 shows an increase in sand portion. The unit does not show any evidence for fluidization as is observed in unit 1 and large-scale bioturbation can be seen. However, the unit is largely devoid of primary structures.

The final core to be logged, the upper *ca* 6 m of Core U1388, is located downslope of PC06 at 663 m water depth, in the northern channel (Hernández-Molina et al. 2014, fig. 2C) (Fig. 2). The upper 3.5 m of this core consists of moderately sorted medium sands (Fig. 3). This uppermost section is divided into two amalgamated sands units separated by a significant surface at 1.4 mbsf. The upper

unit is structureless with only rare large-scale bioturbation observed. There is no evidence for fluid deformation. The underlying sand shows more common large-scale burrows and coarser sand lenses. Poor recovery rates have led to a missing section directly below the sandy facies, below which a further three units have been identified (Fig. 3). These units have very different characteristics from the upper section of the core. The grain size is much more varied, with muddy, silty and sandy layers. Muddier sections are heavily bioturbated with small (millimetre-scale) bioturbation and mottled textures. Siltier and sandier layers show weak lamination. Boundaries are sharp or gradational.

Sediment texture

Grain-size statistics

The mean grain-size and sorting parameters for each core are plotted against depth in Fig. 4. The results show that cores PC04 and PC08, together with their respective trigger cores, are characterized by mainly clay and silt-sized sediment, punctuated by thinner intervals of coarser silts and fine sands. Cores BC05, PC06 and U1388 by contrast, are much more sand rich, with some samples having mean grain sizes in the coarse to very coarse sand range. Sorting statistics are varied, with standard deviations ranging from 1.8 to 5.4 phi units, which corresponds to between moderately sorted and very poorly sorted sediment. The topmost bed of PC06 (fine sand) and the upper parts of U1388 (fine-medium sand) show the best degree of sorting – with σ values of 1.2 to 1.5, or moderately to moderately-well sorted. BC05 also shows better sorting than most of the other cores. The remainder of PC06 rarely shows mean grain sizes smaller than coarse silt ($<31.0\ \mu\text{m}$) or medium silt ($<15.6\ \mu\text{m}$). The same can be said for the lowermost metre of PC04. Therefore, the fine silt ($<15.6\ \mu\text{m}$) and muddy ($<3.9\ \mu\text{m}$) sections show a very different pattern of sorting compared to the coarser sections in PC06/U1388 and the lowermost metre of PC04. The link between grain size and sorting is evidently not straightforward, and appears to differ between the more mud-rich and more sand-rich facies (Fig. 4).

The cross-plot of mean grain size and sorting (Fig. 5A) is quite revealing. Samples taken from PC08, the most mud-rich of the piston cores, show a *decrease* in sediment sorting with increasing mean grain size. The same is true for the mud-rich upper part of PC04. However, for PC06 and the lower more mixed and silt/sand-rich part of PC04, there is a clear trend that shows an *increase* in sediment sorting (i.e. towards better sorting) with increasing grain size. The coarsest of the sediment cores, U1388 and BC05 show a return towards poorer sorting (Fig. 5A).

The cross-plot of mean grain size against skewness data is shown in Fig. 5B. Medium to coarse sands show a coarse to very coarse skew. Fine to very fine sands and silts are generally fine and very fine skewed, reaching a maximum fine skew (-0.7) for a grain size of 50 to 60 μm . The finest silts and clays are more normally distributed. The cross-plot of skewness against sorting, shown in Fig. 6A is coded for grain size rather than core site. This further emphasizes the separation in characteristic trend (and hence behaviour) of the medium to coarse sands, from the fine to very fine sands and from the silts to clays.

Figure 5C shows the relationship between mean grain size and kurtosis. The kurtosis values measured from the sediment distributions range from 0.7 (platykurtic) to 2.6 (very leptokurtic). There is a distinctive trend from platykurtic to very leptokurtic (i.e. increasing peakedness of the distribution) as the grain size increases from coarse silt to fine sand. Fine sands show the highest concentration of data around the mean whereas both the finer and coarser sediments show a broader (platykurtic) spread. Kurtosis has also been plotted against skewness and shows several interesting but complex relationships, this time plotted for grain size rather than core site (Fig. 6B). The different grain-size classes show somewhat different trends. Specifically, a concentration of fine sand data points plots quite separately from the rest the data. The majority of the data plots along two or three straight to curved trend lines (Fig. 6B).

Cumulative frequency curves

Figure 7 shows cumulative frequency curves for all the samples separated into different groups according to their mean grain size for convenience of observation. The distribution curves are significantly different depending on the mean grain size. The fine silts and medium silts show smooth S-shaped curves with a small top and base. From coarse silt to fine sand, there is a diminution of the base curve and increase in size of the upper curve. The medium to coarse sands show straight to slightly kinked distribution with base and top curves mostly absent.

Sediment dating

Table 2 shows the results from the radiocarbon dating. The range of error bars on these dates for ages less than 35 000 years BP are low (71 to 142 years), whereas those up to 45 000 years BP range up to 1916 years (Table 2). The results show that all sediments are late Quaternary to Holocene in age and

reach a maximum calibrated age of $47,983 \pm 1400$ BP, which is dated at latest Pleistocene (Tarantian). Figure 8 illustrates that depositional rates are very different between the core sites. Muddy facies show the slowest rates of accumulation (<15 cm/ka) and sandy facies show the highest rates (>30 cm/ka). Especially high rates of accumulation are observed in the upper section of core U1388 (>100 cm/ka) and in the lower section of PC04 (>198 cm/ka).

INTERPRETATION

Sediment facies

A variety of facies from fine muds through to coarse-grained sands have been interpreted in the logged cores (Figs 3 and 9). Based on the CT-scanning and visual logging, the different facies have been grouped into five principal types (Figs 9 and 10): (i) Facies A – bioturbated muddy facies; (ii) Facies B – mottled silty facies; (iii) Facies C – very fine mottled to fine-grained bioturbated sandy facies; (iv) Facies D – massive and laminated medium-grained sandy facies; and (v) Facies E – coarse-grained sandy and gravel facies. Key features of each of the facies are outlined below.

Facies A – bioturbated muddy facies. Muddy deposits are typically 100% reworked by bioturbation (Fig. 10A). The core surface appears massive with no sedimentary structures seen to have been preserved. Very small-scale burrows such as *Chondrites* can be distinguished in CT-images on the order of some mm in length and diameter. They are chaotically orientated and cross-cutting. Some longer (2 to 10 cm) vertical burrows are also seen, particularly where sand fill causes high-density contrasts with surrounding mud. Organic material and rare shell debris are also seen in the core. Bioturbated muddy facies become thicker and more dominant distally from the Strait of Gibraltar and laterally away from the main MOW core (Fig. 9; PC08).

Facies B – mottled silty facies. The silty facies occur as indistinct beds (or units), typically grading into muds or fine sands. They also form thin discrete layers and lenses within an otherwise mud-rich facies. Highly bioturbated, fine silts commonly show a mottled appearance at the core surface and an absence of primary structures (Fig. 10B). Small-scale burrows are common, but CT-imaging also shows examples of probable *Thalassinoides* and *Planolites* trace fossils that form spherical and elongate mud-filled burrows up to 10 cm in length. There are also various large-scale (up to centimetre diameter) trace fossils that run horizontally across the sediment core (*Zoophycos?*). Where

interbedded with muds or sandy facies, larger burrows can be observed clearly due to infill from adjacent facies.

Facies C – very fine mottled to fine-grained bioturbated sandy facies. Fine-grained sands are found as discrete millimetre-scale layers and lenses within other facies, and as thicker units (reaching >2 m in core U1388). Very fine-grained sands are commonly highly bioturbated, and can therefore appear mottled at the core surface (Fig. 10C). They are also characterized by larger-scale burrows (up to 1 cm in diameter) that form spherical and elongate mud-filled trace fossils such as *Thalassinoides* and *Planolites* up to 10 cm in length. These are particularly visible in CT images (Fig. 10C).

Facies D – massive and laminated medium-grained sandy facies. Where the sand content is high (greater than 70%) there is a marked reduction in evidence of bioturbation. Very large-scale (greater than 1 cm in diameter) burrows may be preserved, together with some rare hints of lamination (Fig. 10D). Most commonly, however, the medium-grained sand facies are massive (Fig. 10D). In some recovered sections, unconsolidated sands show fluid deformation related to the acquisition process (for example, PC06 unit 1). Other units show no evidence of such deformation (for example, PC06 unit 1; U1388 unit 1) and the structureless nature of the sands is interpreted to be depositional. Coarser sandy and gravel lags and lenses are also common. They tend to show gradational bed boundaries, grading to fine sand. There are also abrupt erosive contacts, particularly at the upper boundary.

Facies E – coarse-grained sandy and gravel facies. Coarse-grained sand facies are most commonly found in thin lag layers with erosive basal boundaries. Where thicker (>10 cm) accumulations are recovered, there is bi-gradational grading and evidence of lamination. The coarse sands have a large bioclastic component consisting of broken shelly material (Fig. 10E).

Facies interpretation. Following the three-scale diagnostic scheme proposed for the recognition of contourites by Lovell & Stow (1981), the seismic characteristics of contourites (Faugères et al., 1999; Nielsen et al., 2008; Rebesco et al., 2014) and the most generally accepted criteria for facies identification (Stow et al., 2002; Stow and Faugères, 2008), it can be confidently concluded that all of the sediments studied here are contourites. For the most part, Facies A to E can be associated with the standard contourite C1 to C5 terminology of Stow et al. (2002). Facies A – bioturbated muddy facies represents the lowest energy depositional environment and can be assigned to C1 or C5 facies divisions. Facies B – mottled silty facies represent divisions C2 or C4. Facies C – mottled and

bioturbated fine-grained sandy facies can be assigned to the C3 division and represent an increase in bottom current velocity and/or an increase in sandy sediment supply. However, the massive and laminated medium-grained sand (Facies D), and coarse sandy and gravel facies (Facies E) are not easily represented by the C3 division. The current study's findings suggest that a new sandy contourite facies model is required (see *Discussion*).

Facies sequences

Diverse combinations of Facies A to E are observed in the studied cores, which give a variety of different sedimentary sequences including: bi-gradational sequences; fining upward sequences; coarsening upward sequences; and massive sandy sequences. These can be matched to the standard contourite sequence definitions as defined by Stow et al. (2002) and Stow & Faugères (2008); (i) complete bi-gradational contourite sequence C1 to C5; (ii) mid-only sequence C2 to C4; (iii) base-cut sequence C3 to C5; (iv) top-cut sequence C1 to C3; and (v) C3-only contourite sands (Fig. 9). In some cases, the sands of Facies D and E do not easily fit into the standard bi-gradational sequence.

Bi-gradational sequences

Bi-gradational sequences can consist of the full range of facies from mud to sand (reverse grading), and back to mud (normal grading). There are also a number of examples that lack in C1 and C5 muddy divisions. These reflect a dominantly silty–sandy contourite facies succession that is made up of divisions C2 to C4 only. These are most commonly found in cores collected from channel localities where bottom currents are amplified (Hernández-Molina et al., 2014). They are common in cores PC06 and U1388 collected from a contourite channel and adjacent plastered drift on the Upper Slope (Figs 2 and 9). Other bi-gradational sequences excluding the C3 sandy division are seen in PC04 (Fig. 9). These facies sequences are most prevalent on contourite drifts where the transport capacity and erosional capability of the Mediterranean Outflow Water is diminished.

Base-cut sequence C3 to C5

Base-cut contourite sequences are essentially normally graded, and made up of divisions C3 to C5. Examples are seen in core U1388, located in the northern contourite channel (Figs 2 and 9). There is a relatively abrupt or erosional boundary at the base of the sequence, separating the underlying finer-grained sediment from the overlying sandy facies. This may be disrupted by large-scale bioturbation.

The overlying sand-dominated or silt-dominated divisions are generally structureless and heavily bioturbated.

Top-cut sequence C1 to C3

Some clear examples of top-cut contourite facies sequences are seen in PC04 and PC08 collected from a mounded drift adjacent to a contourite channel (Figs 2 and 9). Top-cut sequences show reverse grading (divisions C1 to C3) and are topped by an abrupt or erosional boundary. Where overlain by finer-grained facies, bioturbation is commonly seen penetrating downward into the underlying unit. Distinct bioturbational characteristics are evident in each division; muds are heavily bioturbated and small-scale burrows occur throughout, whereas divisions C2 and C3 show a more mottled appearance with some larger (>1 cm diameter) burrows observed. In some cases, the overlying sequence shows a thin section of normal grading at its base overlain by a reverse grading trend. Such thin (1 to 2 cm scale) bottom-cut-out contourite sequences appear to be common above top-cut-out contourite sequences (see also PC08, 2.75 mbsf; Fig. 9).

C3-only contourite sands

The cores reveal a number of sand-rich sequences that show no grading, but instead are structureless and thick-bedded. Some of these are made up of the C3 division of the contourite sequence model, whereas others do not fit into the standard model. Core sites U1388 and PC06, collected from within contourite channels proximal to the Strait of Gibraltar, show the best examples. They consist of sand-dominated beds bounded by erosional scour surfaces both to the top and bottom (Fig. 9). Such facies are commonly made up of medium to coarse-grained sands and show little evidence of bioturbation [except for rare large (>1 cm diameter) burrows]. The main 'structures' to note are localized coarse lags and lenses. In some rare examples, laminated beds are observed (for example, U1388 5.2 mbsf; BC05). Medium sands deposited in high-energy contourite channels (as seen in U1388) are generally structureless with erosional surfaces and local coarse lag lenses, whereas the coarser-grained sands may show lamination (BC05). This is in contrast to finer contourite sands deposited on contourite drifts (for example, PC04), which may rarely show indistinct lamination, but are mostly bioturbated.

Contourite sand texture

This study presents one of the largest records of sedimentological data in any modern sandy contourite depositional system to date. The grain-size statistics and distribution curves reveal information on depositional processes and sediment supply.

Mean grain-size

The studied sediment samples show mean grain sizes from 6 μm to *ca* 1000 μm . Because they are contourite sediments, transported and deposited by bottom currents, a strong link between mean grain size and current velocity (or carrying capacity) should be expected. This is confirmed with the observed coarsest grained sediments accumulating in areas of amplified bottom current velocities (for example, BC05 and U1388 located within the contourite channels) and the finest grained sediments being collected from sites more distal from the MOW flow (for example, PC08). There is not, however, a direct linear correspondence for a number of reasons (McCave, 1984; Stow et al., 2008; Stow et al., 2012; Mulder et al., 2013):

- The finest portion of the sediment (clays and fine silts) is most likely to be transported as flocculated particles that are disaggregated during the grain-size analysis. The resulting grain-size distributions and statistics do not, therefore, accurately reflect the hydrodynamic grain size during transport (McCave, 1984).
- The Cadiz contourites are of mixed siliciclastic–bioclastic composition. The coarser sediment therefore includes biogenic tests (foraminifera) and broken-up shell fragments that are hydrodynamically lighter than comparable siliciclastic grains and require lower velocities for transportation.
- Volume of sediment influx into the bottom current system from downslope supply as well as the different distances of transport for individual particles within the bottom current also play an important part in presenting a non-linear relationship between grain size and current velocity. This is observed with coarsest grained sediments being observed proximal to the Strait of Gibraltar, and across the uppermost continental slope (BC05 and PC06 respectively).

For the Gulf of Cadiz, in general, there is a well-documented trend of decreasing grain size associated with weakening bottom current velocity away from the Strait of Gibraltar (Nelson et al., 1993, 1999; Llave et al., 2007). More detailed studies, including the data presented in this paper, show

focused sand accumulation in the parts of the system where bottom water is accelerated (for example, channel core sites BC05 and U1388; Hanquiez et al., 2010; Stow et al., 2013b).

Sorting

Previous studies have shown that contourites tend to be poorly sorted on account of sediment mixing by bioturbation and due to their mixed siliciclastic–bioclastic composition (e.g. Stow et al., 2002; Wetzel et al., 2008; Mulder et al., 2013). However, the data here show some sections of moderate sorting that appear to be broadly linked to mean grain size. Considering the relationship between sorting and mean grain size (Fig. 11), a number of trends can be seen. A significant trend of increased sorting with mean grain size occurs from medium-grained silts through to fine-grained sands. Examining the cumulative frequency curves for these average grain sizes (Fig. 7) and, in particular, the changes of curve trend, suggests a decrease in deposition from the suspended load, and an increasing portion of sediment being deposited from the saltation load with increasing grain size. This shows that as the bottom current velocity and carrying capacity increase, then less of the fine-grained fraction is deposited, but rather remains in suspension and is transported in the water mass.

This distinctive trend is interpreted as a typical depositional trend for contourite deposits. At the finest end of the muddy contourite spectrum, this trend is reversed to give increased sorting with *decreasing* grain size, as displayed by the fine-grained silts and clays from PC04 and PC08, collected from a mounded drift (Fig. 11). A smooth cumulative frequency curve for fine-grained silts (Fig. 7) reflects deposition from very weak bottom currents capable of transporting only the finest silts and clays in suspension, as well as background hemipelagic deposition. The decreasing sorting trend as grain sizes increase to silts is interpreted to be due to increasing rates of bioturbation and mixing with other facies. Data points representing medium and coarse-grained sands (collected from BC05 and U1388 located within contourite channels) show better sorting than the muddy and silty contourites (Fig. 11) and a complete absence of a clay, silt or fine-grained sand fraction (Fig. 7). The cumulative frequency curves show the addition of a traction load component to the deposited sediment (Fig. 7). This reflects the action of strong winnowing (and erosional) processes at these core sites. The sediments never exceed moderate sorting, probably due to a mixed bioclastic–siliciclastic sediment composition, including large bioclastic fragments with smaller lithic grains (Fig. 10E).

Two parallel but distinct trends are noted for samples from PC04 and PC06 (Fig. 11); this is inferred to be due to differences in sediment source. The bioclast and lithic grains collected from core sites BC05, PC04 and PC08 are associated with the Southern channel (Fig. 2). Tracing this channel back towards the Strait of Gibraltar it can be said that sediments were most likely derived from erosion within the Strait and its adjacent margins. PC06, however, is from a plastered drift along the upper slope (Fig. 2). Sediment supply to the PC06 core is therefore inferred to have some component fed from downslope supply (turbidity currents and spillover processes; Armishaw et al., 2002; Nelson et al., 1993, 1999; Hernandez-Molina et al., 2014) as well as alongslope bottom currents.

In summary, the sorting-grain-size trends observed are explained by a combination of current strength, sediment supply and the rate of bioturbation. Fine muddy contourites are influenced by weak bottom currents that supply and transport only fine-grained material. The sediment is completely reworked by bioturbation, and very little coarse material can reach the system. These finest grained sediments show poor sorting. The fine mottled silt-mud contourites (coarse silts and very fine-grained sands) show the worst sorting; poor to very poor. The slightly elevated current speed allows transport and deposition of clay, silt and very fine sand, but little or no winnowing once deposited. There is also extensive bioturbation, which enables the thorough mixing of all size components. Sand-rich contourites deposited by stronger currents (fine-grained and medium-grained sands) show a trend towards improved sorting. This is inferred to be mainly due to increased winnowing. Coarse-grained sands return to poorer sorting on account of a higher bioclastic component.

Skewness and kurtosis

There are very few studies that specifically describe contourite sand skewness (Lovell & Stow, 1981; Stow et al., 1998, 2002) and these authors have suggested that skewness can be used to aid interpretation of sandy facies as contourites. Specifically, they argue that a coarse-tail skew is characteristic of sandy contourites subjected to winnowing. This is certainly seen in this very large dataset, in particular in the proximal channel sands of cores BC05 and U1388, in which the medium and coarse sands (apart from two data points) show a progressive trend towards coarse or very coarse skew (Fig. 5B). It is inferred that prolonged winnowing action has removed the silt and clay fraction into suspension such that only the tractional bedload remains. The kurtosis values show that they are mesokurtic and platykurtic, which is evidence of multiple sediment sources (bimodal or polymodal distributions) as well as moderate sorting (Folk & Ward, 1957) (Fig. 5C).

The finest muds, by contrast, show a more or less symmetrical or normal distribution with zero or low skewness values, which is characteristic of sediments deposited by settling processes, i.e. hemipelagic deposition or very fine-grained contourites (<15 μm) influenced by weak (<10 cm/s) bottom currents with a fine suspension load. These are also platykurtic to mesokurtic in character, indicating broad grain-size distribution and multiple sediment sources. Increasing grain sizes to coarse silts and fine-grained sands show a progressively more fine skewed distribution up to a grain size of 50 to 60 μm , and then a decrease. This trend is interpreted as characteristic of a contourite depositional environment in which the carrying capacity of the current is the limiting factor in transport and where floccing is prevalent. The marked change in trend, which occurs at 50 to 60 μm , coincides with an increase in bedload transport by saltation and the onset of winnowing as an important process. The depositional trend is also characterized by a progressive change from platykurtic to very leptokurtic distributions. The latter are indicative of a single dominant sediment supply expressed as a unimodal distribution and/or good sorting (Folk & Ward, 1957).

These same trends and data clusters are recognized on the cross-plot of skewness and sorting (Fig. 12). In this plot, the mean grain size is colour-coded so that its control can be imaged clearly, and summary distributions are outlined for each distinct grain size (Fig. 12). The other cross-plots in this study also show that there is a strong link between grain size and different statistical parameters. This shows that velocity (and thus transport capacity of the current) exerts a strong control on grain-size distribution. However, the varied and complex patterns observed in cross-plots indicate that there is an additional range of controls on contourite sedimentation including: (i) different transport and depositional mechanisms, such as bedload, suspended load and settling; (ii) mixed sediment sources – bottom current, turbidity current and hemipelagic supply; and (iii) post-depositional bioturbation.

DISCUSSION

This section discusses: (i) the key features of the contourite sediments recovered along the slope of the eastern Gulf of Cadiz; (ii) their textural characteristics and key insights on contourite depositional processes; and (iii) a new facies model for contourite sands, at core and sediment scale, and how this differs from other deep-water facies models.

Characterizing contourite facies

General facies characteristics for the identified sand-rich contourite depositional system along the slope of the eastern Gulf of Cadiz are: (i) indistinct bedding with sharp or gradational contacts; (ii) a typical absence or lack of preservation of primary sedimentary structures and presence of bioturbation; (iii) fine mud to medium-grained sand textures with very poor to poor sorting dominating and only rare moderate sorting; (iv) mixed terrigenous-biogenic composition, and mixed shallow to deep-water components; and (v) bi-gradational sequences, with common partial sequences of coarsening-up and fining-up grain size. There is a distinctive uniformity of appearance between all sites in the study area and with contourite facies from other parts of the Cadiz region. Key diagnostic criteria are outlined below.

Discontinuities

Sharp boundaries are common in the contourite sands observed; they are all interpreted as representing an increase in energy, or bottom current velocity, to a point of non-deposition and/or erosion. In some cases, they have been partially homogenized due to bioturbation. Erosional discontinuities are more commonly associated with coarser sediment, and especially with massive sands and laminated sands observed in cores collected from the contourite channels. This is commensurate with higher bottom current velocity in these regions, but can also indicate downslope sediment supply and local erosion via turbidity currents.

Primary sedimentary structures

The fine-grained contourites (especially muds and silts) examined in this study show little to no primary sedimentary structures such as lamination and cross-lamination. According to most previous studies this is the norm for contourite drift sediments (e.g. Stow & Faugères, 2008), and is due to a lack of preservation of original structures as a result of extensive reworking by bioturbation (Wetzel et al., 2008). This is true also for fine-grained sands, whereas the medium and coarse sands show a significant decrease in the extent of bioturbation, and therefore should be more likely to preserve primary sedimentary structures. Some beds, generally restricted to channel localities such as BC05 show lamination. However, for most of the sandy contourites in this study there is a distinct lack of any structures, or only subtle hints of structure (Fig. 10D). One possible cause of this is the highly fluidized nature of the soft sediment contourite sand cores that results in the destruction of structures during the core acquisition process. For example, the uppermost section of PC06 has been greatly fluidized on acquisition and transportation due to the unconsolidated and coarse-grained nature of the

sediment. Therefore, the original sedimentary features have not been preserved in the top 40 cm (Fig. 3). Nevertheless, there are thick sections of sandy core at depth that are structureless and yet appear to have been undeformed by the coring process. The mechanism for forming these massive sands remains an outstanding question.

Bioturbation

Bioturbation is commonly used as a key diagnostic tool for the identification of contourite sediments, in particular for the muddy to sandy contourites of the standard sequence model (Gonthier et al., 1984; Stow & Faugères, 2008). However, the study of contourite ichnology is still very much in its infancy (Baldwin & McCave, 1999; Löwemark et al., 2004; Shanmugam, 2006; Wetzel et al., 2008; Essex & Stow, 2013) and a definitive ichnofacies model for contourite deposition is yet to be developed (Bromley, 1996). The bioturbation style and trace fossil types observed in this study are broadly comparable to those observed in contourites elsewhere, as outlined by (Wetzel et al., 2008), including a clear link between grain size and ichnofacies. The key observations regarding trace fossil style for the C1 to C5 facies divisions are listed below.

Muddy contourite facies (C1 and C5 divisions) are heavily bioturbated, but with a moderate diversity of ichnofossils. There is intense bio-deformation by small-scale (millimetre-diameter) burrowers such as *Chondrites* leading to a very subtle mottled appearance. Several generations of burrows may be evident, including larger-diameter (centimetre-scale) traces. This complete sediment reworking is attributed to sluggish (<10 cm/sec) bottom currents providing ventilation and nutrients to the sea floor that promotes benthic communities.

Mottled silt and very fine-grained sandy contourite facies (C2 and C4 divisions) are also heavily bioturbated, but with a greater diversity of burrowers than the muddy contourites. There is a very distinctive mottled appearance from intense bio-deformation and multiple generations of ichnofossils. Burrow size ranges from millimetre to centimetre. It is inferred that more energetic (10 to 15 cm/sec) bottom currents provide enhanced ventilation and nutrients to the sea floor that promotes still more active benthic communities, larger organisms and a greater biodiversity compared with muddy contourites (Wetzel et al., 2008).

The *fine-grained sandy contourite facies* (C3 division) recovered from some of the more proximal drift cores in this study show less bioturbation, with fewer, large burrows such as *Thalassinoides* and *Planolites*. Locally, hints of laminations are present, indicating less-extensive sediment reworking by bioturbation. This change of ichnofacies is thought to be due to the increasing bottom current velocities required for fine sand deposition coupled with the removal of the silt and clay fraction. More vigorous currents restrict habitation of the sea floor, as small organisms and juveniles are swept away. Therefore, it is only larger species, and deep penetrating burrowers that are able to survive (Wetzel et al., 2008).

Medium to coarse-grained sandy facies only rarely show any trace fossils indicating that bottom current velocity or accumulation rates were sufficiently high to prevent the habitation of the sea bed by fauna. Where the core reaches high sand content (greater than 70%) there is a marked reduction in bioturbation of the sediment. For clean medium-grained to coarse-grained sands to be deposited, vigorous bottom currents on the order of 50 to 100 cm/sec are expected (Stow et al., 2009) and this is generally too hostile an environment for bottom feeders. Deep burrowing organisms may survive, but the removal of juvenile organisms makes it difficult for a community to establish. The boundaries between this and other facies are often abrupt and erosional. In some cases, amalgamated sands are seen (as in the channel deposits at core site U1388), which can significantly increase the thickness of sandy facies.

Towards a new sand-rich contourites facies model

Compiling the observations and interpretations of the cores in this study, collected from a sand-rich contourite depositional system, a revised set of facies models is proposed to aid in distinguishing contourite deposits from other deep-water sandy facies.

The standard contourite facies model

Fine-grained contourite sands commonly form the C3 division of the standard contourite facies model (Fig. 13). *Muddy contourites* are fine and very fine-grained (mainly C1 and C5 divisions). They show few to no erosional boundaries, but rather gradational changes between facies and continuous sedimentation. They are extensively bioturbated with a wide range of trace fossils often cross-cutting one another. Sediment sorting is poor to very poor, there is a near symmetrical grain-size distribution (zero to very low skewness values) (Fig. 14), and a platykurtic or mesokurtic character. These features

are quite distinctive from those of fine-grained turbidites, which typically show sedimentary structures throughout, partial standard sequences, distinct erosional basal boundaries, and normal grading or graded laminated units as the result of event deposits. However, there are close similarities with hemipelagic sediments, from which they are hard to distinguish (Stow & Tabrez, 1998; Stow & Faugères, 2008).

The C2–C3–C4 divisions comprise mottled silts and fine-grained muddy sands (Fig. 13). Omission surfaces, hiatuses and minor erosive boundaries occur in any of these divisions, although the erosive contacts are more common in C3, and deposition is semi-continuous. Bi-gradational sequences are common, with mostly gradational boundaries and a sequence of ichnofacies associated with the different divisions. A range of partial sequences are observed in the study, as has also been reported in the distal part of the Gulf of Cadiz contourite depositional system (Alonso et al., 2016). Small-scale cyclicality is represented by subtle changes in grain size (Jones, 2016). The fine muddy sands of the C3 division rarely exceed 1 m in thickness. These silty to very fine sandy contourites are moderate to very poorly sorted, have a fine to very fine skew, and range from mesokurtic to very leptokurtic (Fig. 14). These features differ markedly from those of turbidites, which are characterized by partial or complete Bouma Sequences, with erosional bases, normal grading and common primary sedimentary structures. If bioturbation is present in turbidites, it occurs as a top-down feature extending downwards from the muddy E-division. The amount of bioturbation and depth of penetration within the Bouma Sequence depends largely on the time between turbidity current events (Wetzel et al., 2008).

Fine to medium-grained sandy contourite facies

Fine to medium-grained contourite sands, in general, were not seen to occur as part of a standard contourite sequence (or associated top-cut, base-cut variations), rather they tend to occur as distinct units. Two dominant fine to medium-grained sandy contourite facies types can be identified: (i) fine-grained bioturbated sandy facies; and (ii) fine-grained or medium-grained, clean (mud-free) sands that are mostly structureless sands with rare bioturbation. Rare laminated medium sands also occur. Bioturbated fine sands are typical of contourite mounded, or plastered drifts (as seen in cores PC04 and PC06), whereas massive and rare laminated facies occur in higher-energy sandy sheeted drifts and contourite channels (U1388). There is a marked reduction of bioturbation compared with medium-grained contourites, although where present it is larger scale. Erosional and non-depositional surfaces are common. They differ from coarse-grained turbidite facies by their lack of any normal grading, the rare presence of large scale trace fossils (Fig. 13), and their relative lack of any water escape

structures (except where core acquisition has caused deformation as is seen in unit 1 of PC06). This study has shown that sandy contourites range from poorly sorted to moderately sorted and have a coarse skew on account of winnowing and non-deposition of the finer sediment portion (Fig. 14).

Coarse-grained sandy and gravel lag contourite facies

Finally, *coarse-grained sand and gravel lag facies* were poorly represented in this dataset (only observed in contourite channel site BC05). This is also the case in contourite deposits worldwide, with coarse contourite accumulations being extremely rare. They have been seen in other studies in the Gulf of Cadiz (Stow et al., 2013b), and have been collected by BC05 in this study. These facies typically comprise thin gravel-rich layers within a medium-grained to coarse-grained sand and can show lamination. Unlike coarse-grained turbidites or downslope mass wasting deposits (debrites), these gravel facies tend to be laterally constricted and relatively finer grained. This is not surprising on account of the large amount of energy required to transport such grain sizes in bottom currents. As a result, they tend to be limited to contourite channels (for example, BC05). They rely on a local sediment supply of gravel and coarse sand and abnormally vigorous bottom currents (>50 cm/sec) and occur principally within contourite channel systems (Stow et al., 2013b; Hernández-Molina et al., 2014).

CONCLUSIONS

The results presented document the world's largest present-day sand-rich deep-water contourite depositional system, which extends over an area of 4000 km² west and north-west from the exit of the Strait of Gibraltar. This has accumulated at mid-slope depths (500 to 1200 m) under the influence of the Mediterranean Outflow Water (MOW). All sediments examined for this study can be reliably interpreted as contourites on the basis of their oceanographic setting, seismic and morpho-sedimentary features, and detailed facies characteristics. However, there is a wide range of contourite facies present – not only sands. These include: (i) bioturbated muddy contourites; (ii) mottled silty contourites; (iii) very fine-grained mottled and fine-grained bioturbated sandy contourites; (iv) massive and laminated medium-grained sandy contourites; and (v) coarse-grained sandy/gravel contourites. These facies tend to co-occur in partial and full bi-gradational facies sequences.

Grain-size analysis of 675 samples from the eastern Gulf of Cadiz reveals distinctive trends in textural properties linked to depositional process under the action of bottom currents (Fig. 14). The finest muddy contourites (<20 μm) show normal grain-size distributions, poor to very poor sorting, zero or low skewness, and platykurtic to mesokurtic kurtosis. These are deposited by settling from weak (<10 cm/sec) bottom currents with a fine suspension load. Muddy to fine sandy contourites (20 to 200 μm) show a uniform trend towards better sorting and a more leptokurtic distribution. They show a progressively finer skewed distribution up to a grain size of 50 to 60 μm , and then a decrease. These are the typical depositional trends for contourites – as velocity and carrying capacity increase (from 10 to 50 cm/sec), then more of the finest fraction remains in suspension. From 50 to 60 μm , bedload transport (principally saltation load) becomes more important than suspension transport and there is a marked increase in winnowing. Clean sandy contourites (>200 μm) from the study area show a range of textural properties. They are better sorted than the fine muddy contourites and have no clay content, but display varied skewness and kurtosis values. These contourites result from the action of dominant bedload transport (saltation and traction), strong winnowing and local erosion at current speeds in excess of around 50 to 70 cm/sec.

Important controls on the nature and distribution of contourites; their textural, ichnofacies and facies models, include: (i) transport and depositional processes from bottom currents and their velocity variation; (ii) sediment supply from mixed sources; and (iii) the post-depositional action of burrowing organisms.

The results presented from the Gulf of Cadiz contourite depositional system provide an important modern analogue for contourite sands in the subsurface. Research into contourite sands and other sandy deposits generated or affected by bottom currents is in its infancy. A better understanding of their sedimentary facies associations is required because these deposits are expected to become important for hydrocarbon exploration in next decades. The study of contourites and related sands is a novel and rapidly expanding area of research and a collective effort towards their accurate interpretation and understanding is required. Therefore, further studies along other margins should clarify the real distinction between sandy contourites and reworked deep-water sands derived from turbiditic and other processes. Such studies will test the facies model proposed in the Gulf of Cadiz against other margins with various geological and oceanographic configurations.

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FIGURE CAPTIONS AND TABLES

Table 1: The sediment cores used in the study of sediment facies sequences in the Gulf of Cadiz and their depositional environment. *In this study, logging was restricted to the uppermost 10 m of the Integrated Ocean Drilling Program (IODP) cores.

Table 2: Calibrated ^{14}C dates and errors acquired from four cores across the eastern Gulf of Cadiz. Dating completed at: (i) the SUERC Radiocarbon Facility, (University of Glasgow); and (ii) the Laboratoire de Mesure du Carbone 14 (Université Paris-Saclay).

Fig. 1. The Gulf of Cadiz and the main study area (red outline box). Key oceanographic components marked are the eastward-flowing Atlantic Inflow Waters (AIW) and the Mediterranean Outflow Water (MOW). Numbers relate to the Hernández-Molina et al. (2003) contourite depositional sectors: (1) proximal scour and sand ribbons sector; (2) overflow-sedimentary lobe sector; (3) channels and ridges sector; (4) active contourite drift sector; and (5) submarine canyons sector. Modified from Hernández-Molina et al. (2003) and Zitellini et al. (2009).

Fig. 2. Location of the recovered sediment cores in this study and their position in relation to morpho-sedimentary features defined by Hernández-Molina et al. (2014). Modified from Habgood et al. (2003) and Hernández-Molina et al. (2006, 2014).

Fig. 3. Lithological logs for the seven cores used in this study. Core photographs and CT images shown alongside the lithological logs where available. Samples taken for smear slide analysis, grain size measurement and ^{14}C dating are indicated.

Fig. 4. Sorting plot with average grain-size curve. Measurements were made at a 2 cm resolution except for core U1388 where samples are indicated as points. Some sections show increasing sorting with increasing grain size, whereas others show decreasing sorting with increasing grain size.

Figure 5: (A) Cross-plot of sorting against mean grain-size reveals that mud and fine silt has a negative relationship with sorting and medium silt to fine sand has a positive association with sorting. Medium to coarse sands return to poorer sorting. (B) Mean grain size versus skewness. Finest grain sizes show a minor coarse tail or near-normal distribution. Sands show a coarse skew. (C) Mean grain

size versus kurtosis. Fine sands show a higher concentration of grain sizes around the mean whereas the finer silts and coarser sands show a broader (platykurtic) spread.

Figure 6: (A) Sorting versus skewness cross-plot showing multiple trends. Data points are colour-coded with respect to mean grain size, and there appears to be a direct link between cross-plot trends and mean grain size. (B) Kurtosis versus skewness cross-plot showing multiple trends. Data points are colour coded with respect to mean grain size.

Fig. 7. Cumulative frequency curves for each mean grain-size division. The curves show an important change in the dominant transport process between fine silt and very coarse sand, as represented by a change in the trend of the curve.

Fig. 8. Age–depth plots and accumulation rates for the dated sediment cores. The anomalous date at about 3 m depth in the very sand-rich core U1388 is probably due to an especially high level of sediment re-working at the sea floor.

Fig. 9. Facies observed across the seven sediment cores in the Gulf of Cadiz with assigned contourite divisions C1 to C5 after Stow & Faugères (2008). Localities are indicated on Fig. 2. The standard contourite sequence is also shown (modified from Gonthier et al., 1984; Stow & Faugères, 2008). Locations of core photographs in Fig. 10 are indicated.

Fig. 10: Contourite facies observed across the eastern Gulf of Cadiz. (A) to (C) Displayed as core photographs (left) and CT scan (right). (A) Facies A – bioturbated muddy facies; PC08, 0.5 mbsf. (B) Facies B – mottled silty contourite facies; PC04, 2.7 mbsf. (C) Facies C – very fine sandy mottled to fine-grained sandy bioturbated facies; PC06, 0.4 mbsf. (D) Facies D – massive (left) or laminated (right, boxed) medium-grained sandy contourite facies; U1388, 0.94 mbsf; U1388, 5.2 mbsf. (E) Facies E – coarse-grained sandy and gravel facies; BC05, 0 cm.

Fig. 11. A cross-plot of sorting against mean grain size reveals that the finest sediments have a negative relationship with sorting and medium silt to fine-grained sand has a positive association with sorting. Medium to coarse sands return to poorer sorting. Sorting definitions given by Folk (1964).

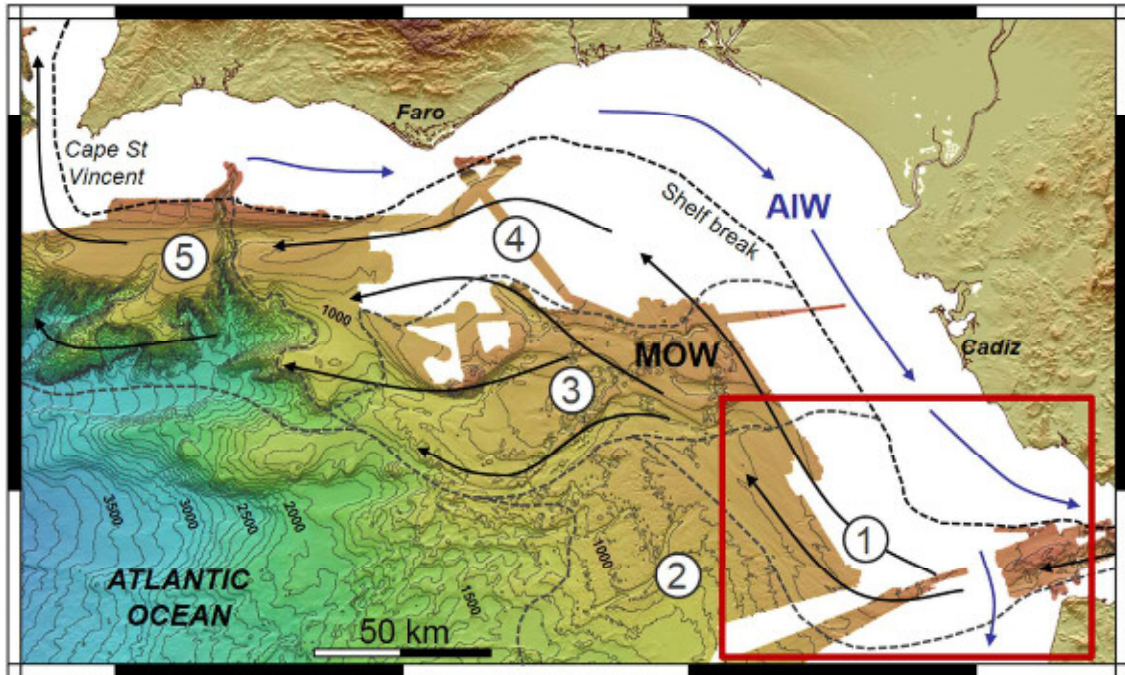
Figure 12: Sorting versus skewness cross-plot showing multiple trends. Data points are colour coded with respect to mean grain size, and there appears to be a link between cross-plot trends and mean grain size. Grain-size distribution curves for each of the highlighted areas of the graph. Skew changes in response to grain size. See text for explanation.

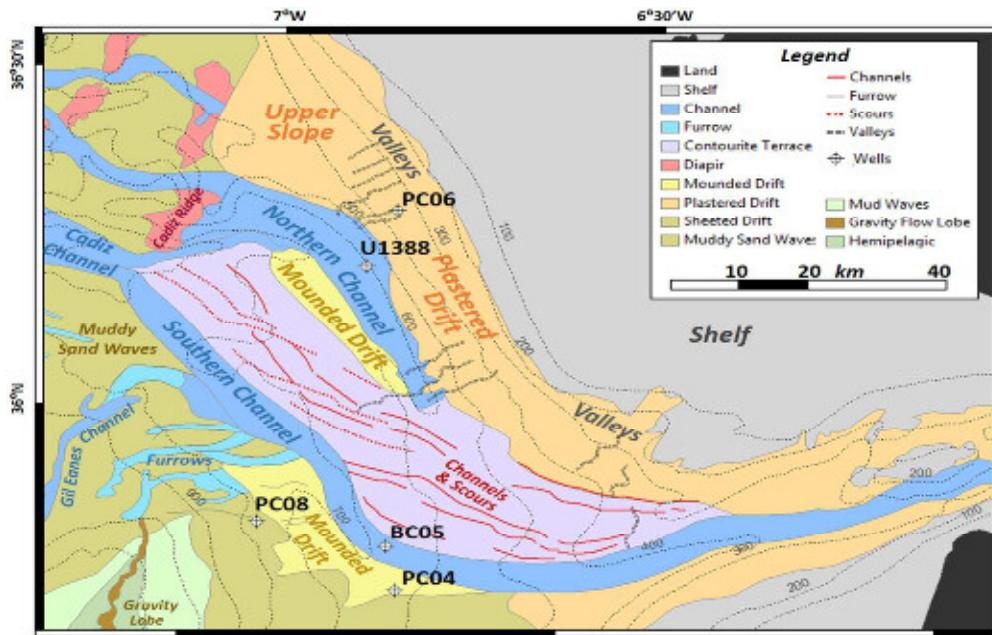
Fig. 13. Contourite sand facies and key characteristics.

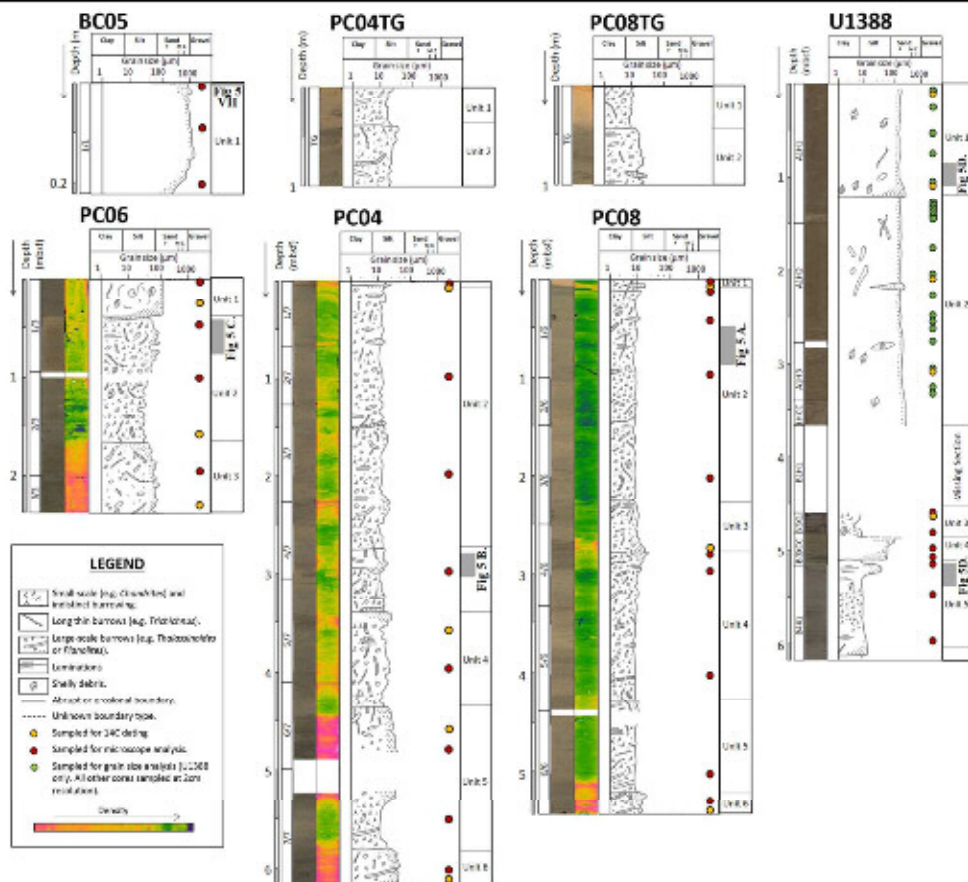
Fig. 14. End-member models of grain-size distribution as expected for each depositional process acting in deep water. In reality, there is a continuum between processes; and mixed systems may occur. Distribution is also strongly controlled by sediment provenance.

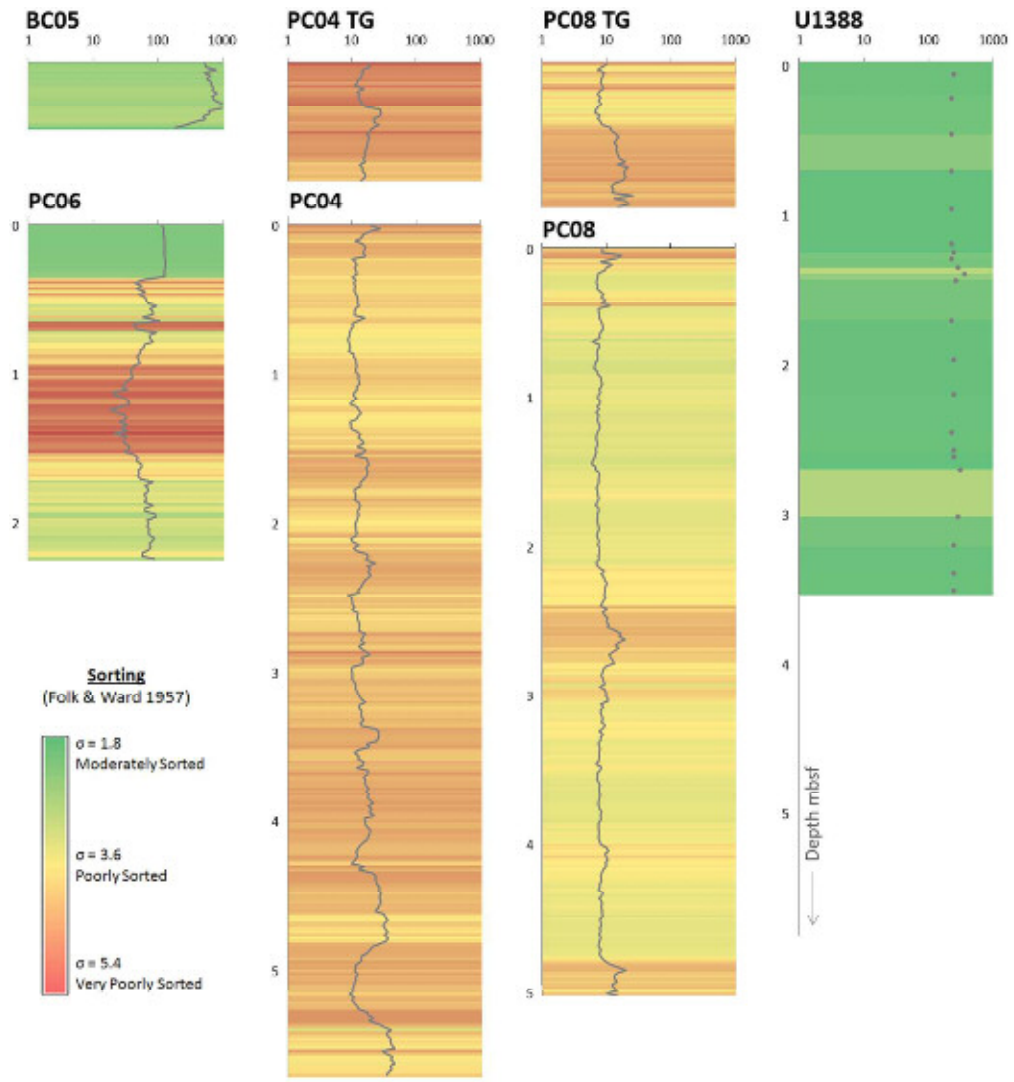
Cruise	Core	Latitude (°N)	Longitude (°E)	Water depth (m)	Core length* (m)	Contourite deposit type (from Hernández-Molina et al (2014))	# Grain-size samples
CONTOURIBER	BC05	35.845	-6.780	738	0.23	Contourite channel	23
CONTOURIBER	PC04	35.736	-6.732	658	6.22	Mounded drift	286
CONTOURIBER	PC04 TG	35.736	-6.732	658	1.00	Mounded drift	39
CONTOURIBER	PC06	36.305	-6.764	490	2.40	Plastered drift	17
CONTOURIBER	PC08	35.821	-6.968	961	5.36	Mounded drift	252
CONTOURIBER	PC08 TG	35.821	-6.968	961	1.00	Mounded drift	48
IODP 339	U1388	36.161	-6.476	663	121.00*	Contourite channel	33

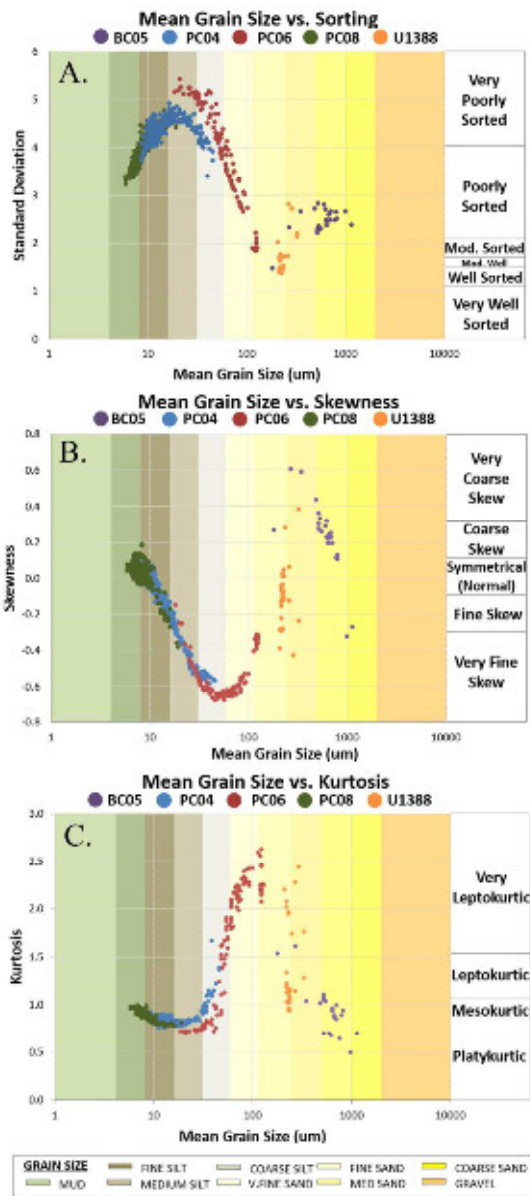
Core	Sample mbsf	¹⁴ C age (BP)	Error ±	Age cal BP	Error ±	Laboratory	Material
PC06	0.36	5525	28	5917	71	SUERC	Mixed planktonic
	1.52	9585	29	10455	87	SUERC	Mixed planktonic
	2.24	10645	35	11968	142	SUERC	Mixed planktonic
U1388	0.08	8390	40	8983	138	Lab. Mesure du ¹⁴ C	Mixed planktonic
	3.08	7990	45	8455	97	Lab. Mesure du ¹⁴ C	Mixed planktonic
	4.76	42220	950	45449	1916	Lab. Mesure du ¹⁴ C	Mixed planktonic
PC04	0.04	11292	30	12768	134	SUERC	Mixed planktonic
	3.56	34279	183	38910	514	SUERC	Mixed planktonic
	4.38	35212	196	39836	765	SUERC	Mixed planktonic
	5.64	35750	470	39960	1075	Lab. Mesure du ¹⁴ C	Mixed planktonic
PC08	0.04	4973	28	5346	85	SUERC	Mixed planktonic
	2.60	39032	305	42949	555	SUERC	Mixed planktonic
	5.04	45900	1400	47983	1400	Lab. Mesure du ¹⁴ C	Mixed planktonic

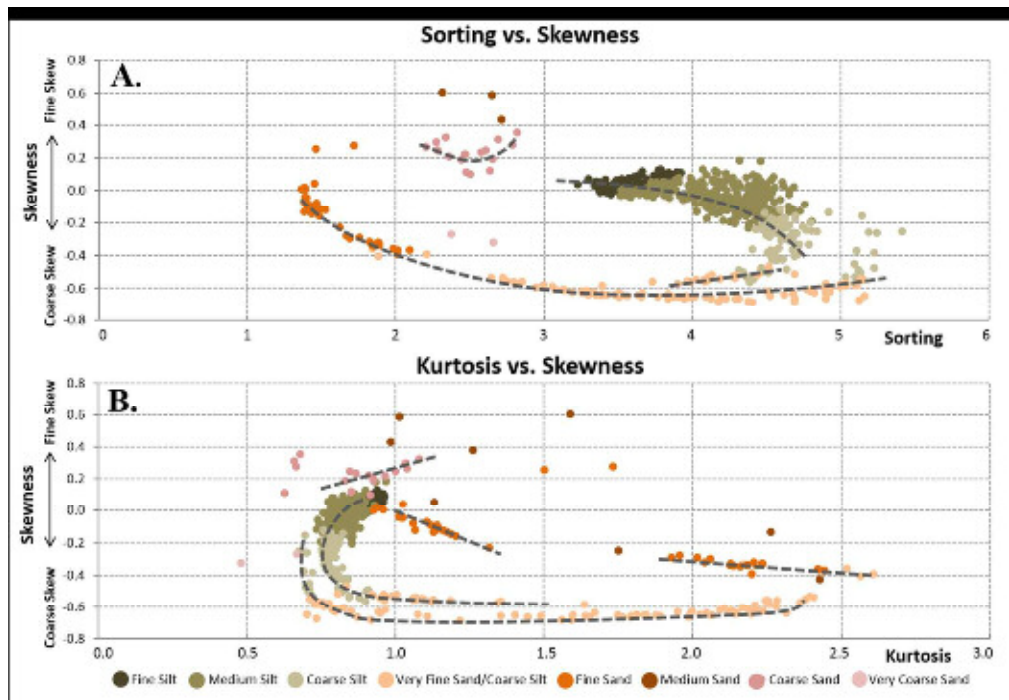


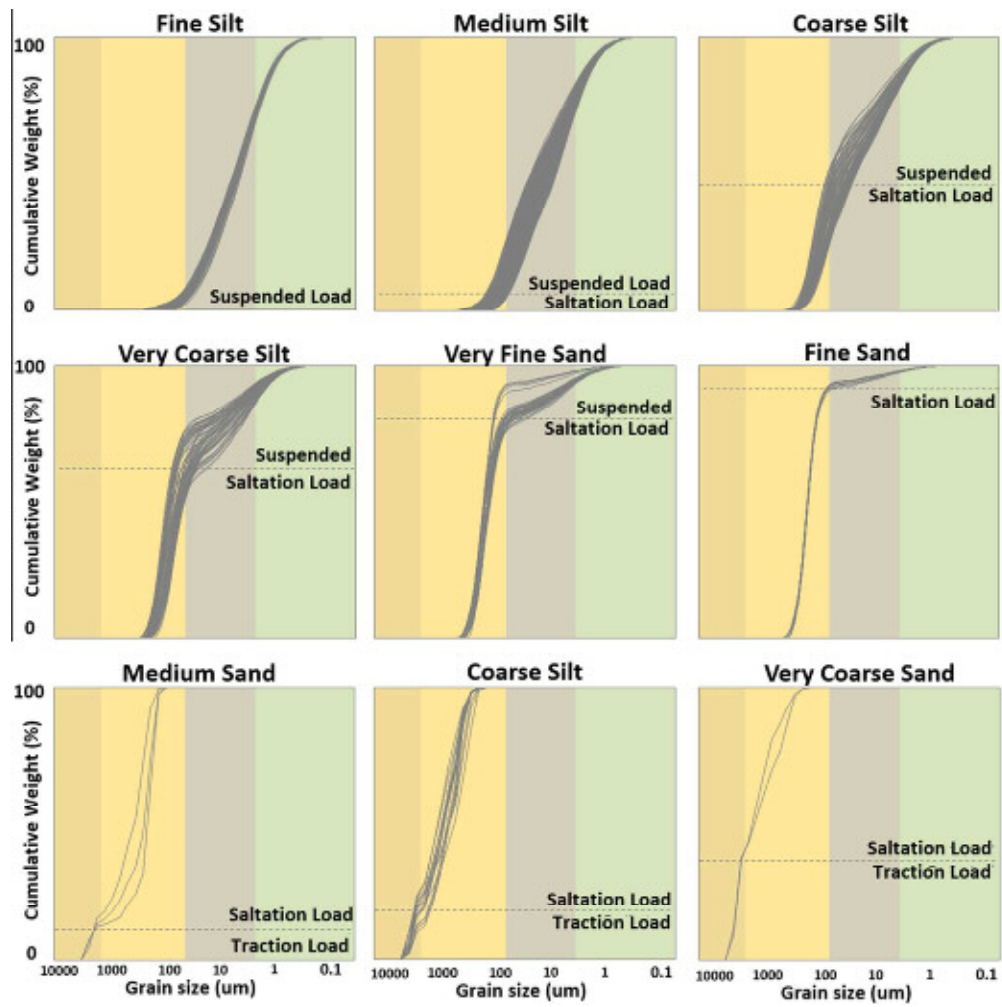


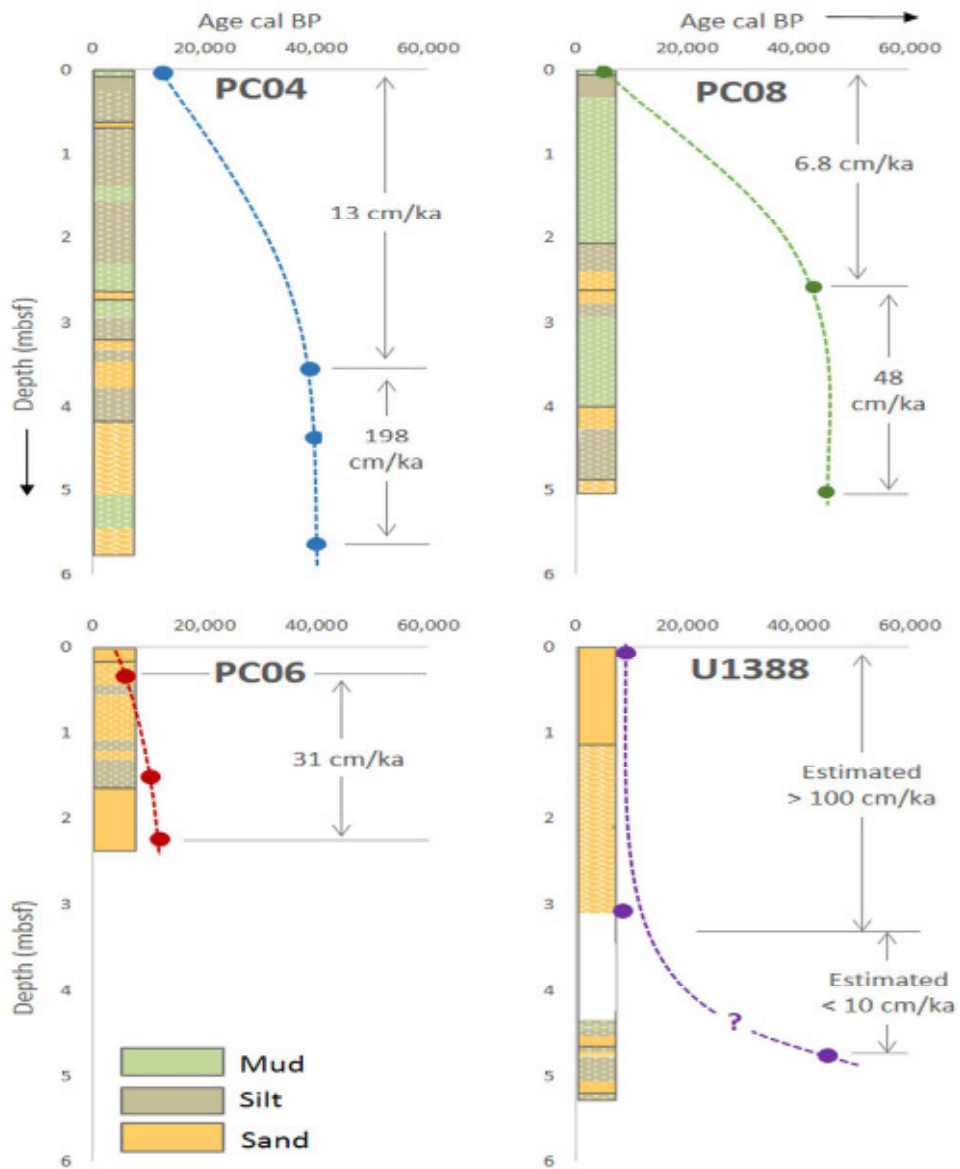


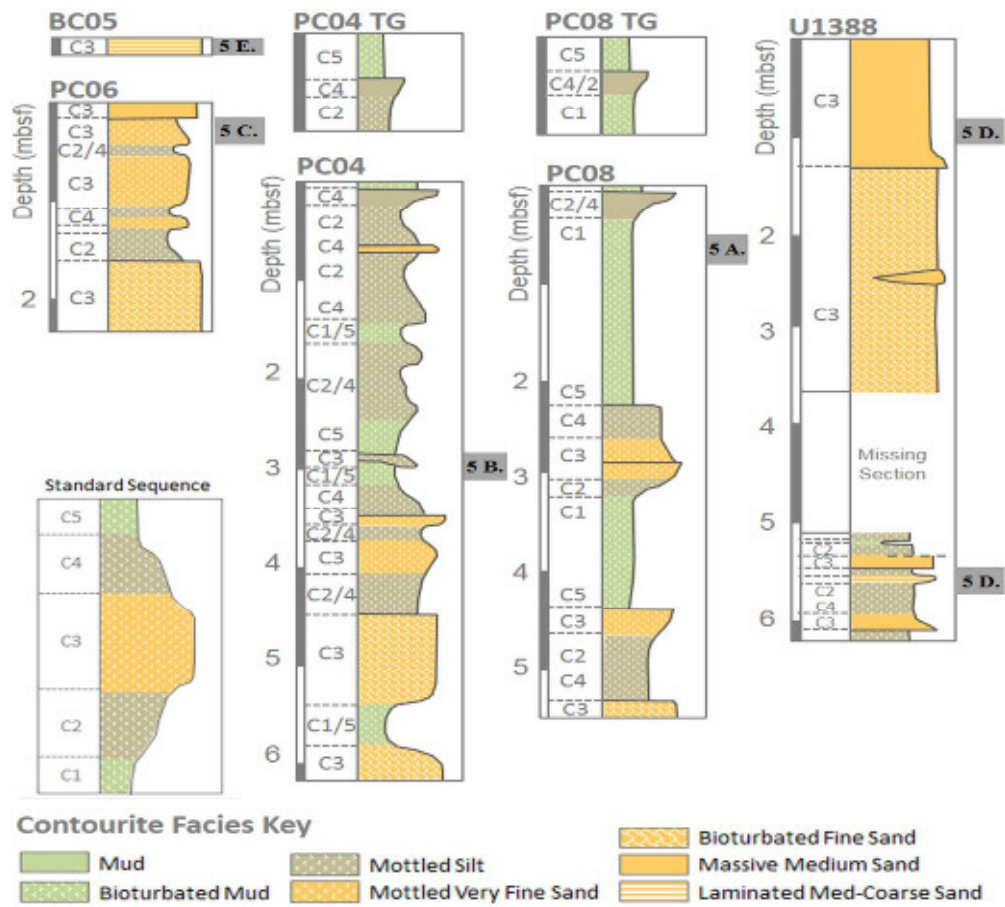


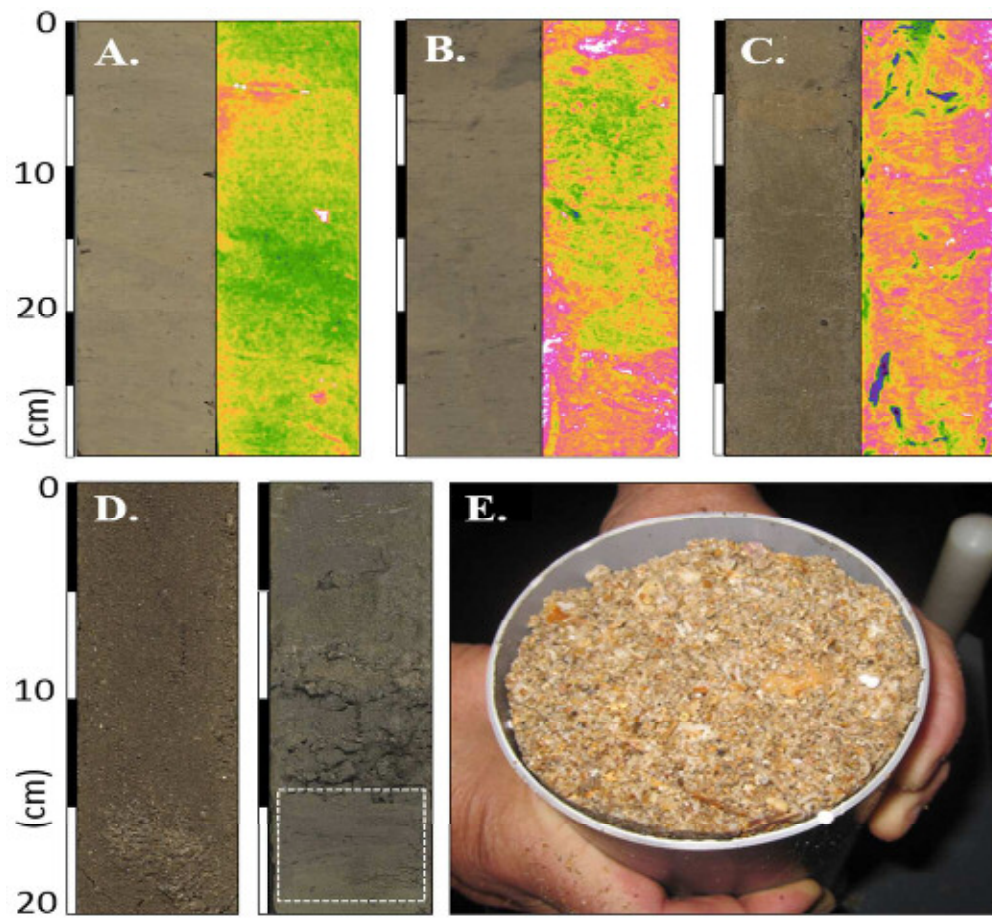




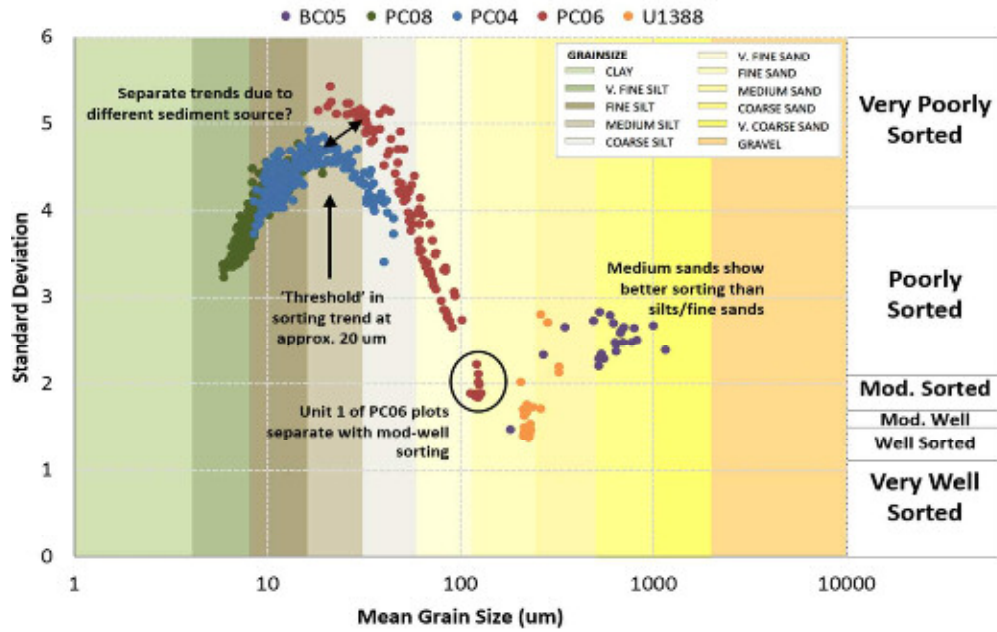


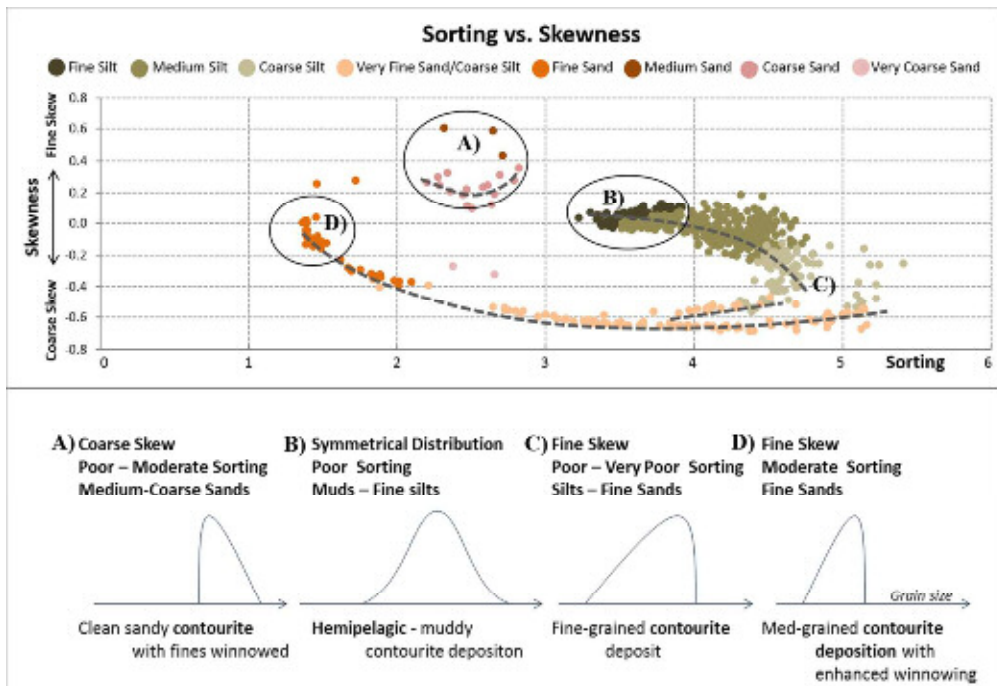









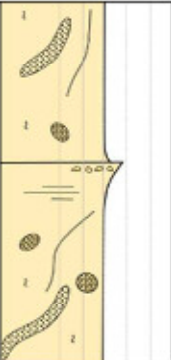
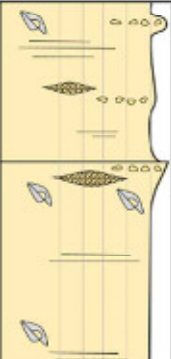


Mean Grain Size vs. Sorting





c s f m c g

<p>C1</p>  <p>C2</p>  <p>C3</p>  <p>C4</p>  <p>C5</p> 	<p style="text-align: center;">Standard Contourite Facies Model</p> <p>Facies: Very fine grained sand, heavily bioturbated and mottled in appearance. Bi-gradational grading to extensively bioturbated silts and muds.</p> <p>Ichnofacies: Very highly bioturbated with a mottled appearance. Larger (cm-scale) burrows also common.</p> <p>Sorting: Very poor</p> <p>Skew: Fine</p> <p>GDE: Contourite mounded drifts</p>
<p>C3</p> 	<p style="text-align: center;">Fine to Medium Grained Sandy Contourites</p> <p>Facies: Fine grained bioturbated to massive medium grained sands. Omission or erosional surfaces common. Some indications of grading, often associated with erosional surfaces.</p> <p>Ichnofacies: Reducing levels of bioturbation with increasing grainsize. Large-scale burrows.</p> <p>Sorting: Moderate</p> <p>Skew: Fine to coarse</p> <p>GDE: Contourite plastered drifts</p>
<p>C3</p> 	<p style="text-align: center;">Coarse Grained Sandy and Gravel Lag Contourites</p> <p>Facies: Medium to very coarse sands with gravel lags. Can be laminated. Erosional scours and surfaces common. Large bioclastic component common.</p> <p>Ichnofacies: No evidence of bioturbation</p> <p>Sorting: Moderate to poor</p> <p>Skew: Coarse to very coarse</p> <p>GDE: Contourite channels</p>

