

1       **Effects of heterogeneity, connectivity and density variations on**  
2       **mixing and chemical reactions under temporally fluctuating flow**  
3       **conditions and the formation of reaction patterns**

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7       **Key Points:**

- 8       • The coupling of transient forcing and heterogeneity leads to strong local enhance-  
9       ment of mixing and reactivity.
- 10      • Buoyancy forces cause a compression of the interface which in turn increases the  
11      mixing and reaction efficiency.
- 12      • These mechanism provide a plausible explanation for the formation of dissolution  
13      patterns.

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## Abstract

Solute mixing, spreading and fast chemical reactions in aquifers are strongly influenced by spatial variability of the hydraulic properties, temporal flow fluctuations and fluid density differences. We study the coupling of heterogeneity, transient forcing and density-driven flow on mixing and chemical reactions between two fluids of different density under a stable stratification. We consider the reaction of the fast dissolution of calcite. We find that temporal fluctuations and heterogeneity cause strong local enhancement of the mixing and reaction rates and this impact increases with the degree of connectivity of hydraulic conductivity. The global mixing and reactivity, however, are on the order of or smaller than their homogeneous counterparts due to heterogeneity-induced fluid segregation. The local maxima of the mixing and reaction rates are found to be located around strongly stretched regions corresponding to high velocity zones where dispersive mass transfer mechanisms are increased by dispersion. We also find that density variations compress the interface, which in turn emphasizes local maxima in mixing and reaction rates. Numerical results provide evidence that the stretching of the interface induced by spatial heterogeneity and transient effects coupled with density variations leads to the formation of complex patterns of reactive hotspots, zones of enhanced reaction efficiency, and that its distribution is directly linked to the deformation properties and topology of the flow field. These results provide new insights into the role of spatial and temporal variability on the mixing and reaction efficiency as well as the formation of reactive geochemical patterns in actual environmental systems.

## 1 Introduction

Mixing and dispersion of solutes and induced reactive processes occurring in flow through porous media at different scales have recently received increasing attention because they play a key role in a number of important applications, such as groundwater management, aquifer remediation, artificial recharge and risk assessment in geological hazardous waste storage. While substantial progress has been made to characterize flow and solute transport processes, addressing conservative mixing and reactive transport in real-world environments remains a largely open question due to the inherent complexity of natural flow and transport systems.

Solute spreading and mixing are strongly affected by spatially nonuniform variations of the flow field in response to spatial heterogeneity of the hydraulic properties of

46 the subsurface. Significant research has been devoted to quantify the impact of hetero-  
47 geneity on the spatial extent of transported solute plumes [Gelhar, 1993; Dagan, 1987;  
48 Rubin, 2003], non-Fickian large scale transport [Berkowitz et al., 2006; Neuman and Tar-  
49 takovsky, 2009], and the impact of heterogeneity on solute mixing and chemical reaction  
50 [Dentz et al., 2011]. Recently, the mechanisms of mixing in heterogenous porous media  
51 have been related to the deformation properties of the flow field through the stretching and  
52 coalescence of material elements [Villiermaux, 2012; Le Borgne et al., 2015]. Heteroge-  
53 neous flow fields are characterized by the emergence of connected high velocity channels  
54 which form braided networks of preferential flow paths. The resulting connectivity struc-  
55 ture, which mainly depends on the log-conductivity variance and the dimensionality of  
56 the system [Fiori and Jankovic, 2012], dramatically influences flow and solute transport  
57 in aquifers [Fogg et al., 1998, 2000; Le Borgne et al., 2010; Bianchi et al., 2011; Pedretti  
58 et al., 2013; Willmann et al., 2008]. Thus, accurate characterization of channeling is essen-  
59 tial to analyze flow and transport processes in porous media and intensive research efforts  
60 over the last decades have been devoted to define connectivity indicators [Zinn and Har-  
61 vey, 2003; Knudby and Carrera, 2005; Renard and Allard, 2013] and delineate preferential  
62 flow paths in groundwater systems [Tyukhova et al., 2015; Tyukhova and Willmann, 2016].  
63 Although these previous investigations provide invaluable insights into the role of hetero-  
64 geneity and connectivity on the dispersion and mixing dynamics, they consider steady uni-  
65 form flow and rarely incorporate the influence of transient effects.

66 Many real systems are strongly influenced by transient effects on multiple time scales,  
67 such as diurnal variations in evapotranspiration or seasonal pumping, recharge and droughts.  
68 Numerous studies have investigated the impact of temporal flow fluctuations on solute  
69 transport of contaminant plumes. However, a common approach used is to assume quasi-  
70 stationarity of the flow field, this means disregarding storativity, and constant dispersion.  
71 Under such conditions, a key finding has been that transient forcing parallel to the main  
72 flow direction has little influence on mixing, whereas transverse temporal fluctuations en-  
73 hance transverse dispersion and reduce longitudinal dispersion [Rehfeldt and Gelhar, 1992;  
74 Dagan et al., 1996; Dentz and Carrera, 2003; Cirpka and Attinger, 2003; de Dreuzy et al.,  
75 2012]. For homogeneous porous media with velocity-dependent dispersion transient forc-  
76 ing leads to an important increase in transverse dispersion whereas longitudinal dispersion  
77 is reduced [Kinzelbach and Ackerer, 1986; Goode and Konikow, 1990]. Furthermore, Pool  
78 et al. [2016] derived an effective time-average formulation to describe the coupling be-

79 tween mixing and oscillatory transient flow under nonstationary flow conditions for finite  
80 storativity. These authors demonstrated analytically, numerically and experimentally for  
81 homogeneous aquifers that temporal fluctuations of the flow conditions lead to a complex  
82 space and time-dependent flow response which induces enhanced spreading and mixing  
83 of dissolved substances. The systematic investigation of the interaction of medium hetero-  
84 geneity, temporal flow fluctuations under velocity-dependent dispersion is an outstanding  
85 question.

86 Another key property that may markedly impacts groundwater flow patterns and sig-  
87 nificantly affect transport of active and passive contaminants, their dispersion, mixing and  
88 reaction behaviour is fluid-density variations. Density-driven flow plays an important role  
89 in many environmental problems such as seawater intrusion, freshwater-brine interfaces,  
90 dense contaminant plume migration, transport of DNAPLs, salt lakes, heat in geothermal  
91 systems and CO<sub>2</sub> sequestration [*Diersch and Kolditz, 2002; Neufeld et al., 2010; Hidalgo*  
92 *et al., 2015*]. In flow systems with variable density, heterogeneity and temporal fluctua-  
93 tions may significantly influence flow and transport over many length scales. In the lit-  
94 erature, the research on the effects of heterogeneity on solute transport in variable den-  
95 sity systems has been mainly focused on the analysis of unstable convective flows, where  
96 a dense fluid overlies less dense one [*Schincariol and Schwartz, 1990; Schincariol et al.,*  
97 *1997; Schincariol, 1998; Simmons et al., 2001*]. In this case, heterogeneity induces en-  
98 hanced vertical mixing and spreading of the plume. On the other hand, for seawater intru-  
99 sion problems heterogeneity affects the interface inland migration and leads to a signifi-  
100 cant widening of the mixing zone between fresh and saltwater [*Schwarz, 1999; Held et al.,*  
101 *2005; Abarca, 2006; Kerrou and Renard, 2010; Pool et al., 2015*]. Thus, the coupling of  
102 heterogeneity and density effects is expected to increase the mixing efficiency. For verti-  
103 cal front displacement under stable density stratification this is different. Under these flow  
104 conditions, *Landman et al. [2007b]* demonstrated that density variations smoothen the ir-  
105 regularities in the interface and cause the deformation of the front and then the mixing  
106 zone between the two fluids to decrease due to the competition between viscous forces  
107 induced by heterogeneity and buoyancy forces. Therefore, proper quantification of the  
108 competition between spatial heterogeneity and the non-linear nature of density-driven flow  
109 coupled with transient forcing on solute transport is a key to understanding and predicting  
110 the degree of local mixing and its impact on chemical reactions.

111 Besides heterogeneity of the medium, temporal flow fluctuations and density varia-  
112 tions of the fluid, solute transport in the subsurface is strongly affected by chemical reac-  
113 tions between dissolved chemicals (fluid-fluid) as well as between the dissolved chemicals  
114 and the solid matrix (fluid-solid). Mixing processes in porous media are particularly im-  
115 portant for controlling chemical reactions [e.g., *Rezaei et al.*, 2005; *De Simoni et al.*, 2005;  
116 *Tartakovsky et al.*, 2008; *Cirpka et al.*, 2008; *Dentz et al.*, 2011; *de Anna et al.*, 2014a]  
117 and can be related to the deformation properties of the flow field [*de Anna et al.*, 2014b;  
118 *Le Borgne et al.*, 2014; *Hidalgo et al.*, 2015; *Bandopadhyay et al.*, 2017]. A wide vari-  
119 ety of chemical reactions occur between fluids and fluids and solids in density-dependent  
120 flow systems, including calcite dissolution and precipitation [*Rezaei et al.*, 2005; *Izgec et al.*,  
121 2008], cation exchange [*Appelo and Willemsen*, 1987; *Appelo*, 1994; *Giménez-Forcada*,  
122 2010], dissolution/precipitation of gypsum [*Magaritz and Luzier*, 1985; *Stoessel et al.*,  
123 1993; *Gomis-Yagües et al.*, 1997], dolomitization [*Müller and Fischbeck*, 1973; *Magaritz*  
124 *et al.*, 1980] and karstification [*Plummer*, 1975; *Hanshaw and Back*, 1979; *Smart et al.*,  
125 1988; *Singurindy et al.*, 2004]. To date the interplay between temporal and spatial variabil-  
126 ity and density effects on mixing and dispersion of solutes and chemical pattern formation  
127 remain largely unexplored, despite its importance in controlling reaction rates and reactive  
128 transport processes for large-scale problems.

129 The objective of this study is to quantify the complex interactions between aquifer  
130 connectivity, transient-forcing and fluid-density effects on mixing and chemical reactions  
131 in heterogeneous porous media. To this end, we perform two-dimensional constant and  
132 variable-density flow and transport Monte Carlo simulations. Log-normally distributed  
133 random permeability fields assuming a Gaussian model are generated. In order to evalu-  
134 ate the impact of aquifer connectivity, more complex heterogeneous fields characterized by  
135 connected patterns of high and low conductivity are considered. We analyze the mecha-  
136 nisms of mixing by focusing on the measures of the interface width, as well as diagnos-  
137 tics which quantify the deformation of the interface. For the reactive transport system, we  
138 consider the irreversible reaction of calcite dissolution. To this end, we map the reactive  
139 transport problem onto a conservative one through the use of mixing ratios and conser-  
140 vative components [*De Simoni et al.*, 2005, 2007]. Global and local reaction efficiency is  
141 evaluated with emphasis placed on the location and topology of hotspots, zones of en-  
142 hanced reactivity. The reaction patterns formation is related to the deformation of material  
143 fluid elements in the heterogeneous flow field and specifically its stretching behavior. In

144 this study, we provide a plausible explanation for the formation of geochemical reaction  
 145 patterns observed in coastal karst aquifers. The enhanced understanding and quantification  
 146 of the interplay between spatial and temporal variability and density-driven flow developed  
 147 in this study is expected to provide new insights into mixing and reactive transport mech-  
 148 anisms which is essential for their diagnosis and prediction in heterogeneous subsurface  
 149 environments.

## 150 2 Problem Statement and Methodology

151 We investigate the effect of heterogeneity and density variations on mixing and  
 152 chemical reactions between two fluids (fresh and saline water) for a stable stratification  
 153 in response to periodic temporal flow fluctuations. A fully saturated two-dimensional het-  
 154 erogeneous porous medium is considered with the  $z$ -axis pointing vertically upward, see  
 155 Figure 1. Fluid mass conservation is given by the continuity equation [Voss, 1984],

$$\rho S_p \frac{\partial p}{\partial t} = -\nabla \cdot \rho \mathbf{q}, \quad (1)$$

156 where  $S_p = \phi \alpha_w + (1 - \phi) \alpha_s$  [ $\text{LT}^2\text{M}^{-1}$ ] is the specific pressure storativity,  $\alpha_w$  [ $\text{LT}^2/\text{M}$ ] the  
 157 compressibility of water,  $\alpha_s$  [ $\text{LT}^2/\text{M}$ ] of the solid matrix,  $\phi$  [ $\text{L}^3\text{L}^{-3}$ ] is porosity,  $\rho$  [ $\text{ML}^{-3}$ ]  
 158 fluid density and  $\mathbf{q}$  [ $\text{LT}^{-1}$ ] Darcy velocity defined below. Using the equivalent freshwater  
 159 head  $h_f = p/(\rho_f g) + z$  with  $\rho_f$  [ $\text{ML}^{-3}$ ] freshwater density and the Boussinesq approxima-  
 160 tion, this equation reads as

$$S_f \frac{\partial h_f}{\partial t} = -\nabla \cdot \mathbf{q} \quad (2)$$

161 where  $S_f = \rho_f g S_p$  [ $\text{L}^{-1}$ ] is specific storativity. The Darcy velocity  $q$  is given by the Darcy  
 162 equation, which reads in terms of equivalent freshwater head as

$$\mathbf{q} = -k \left( \nabla h_f + \frac{\rho - \rho_f}{\rho_f} \hat{e}_z \right) \quad (3)$$

163 with  $k$  [ $\text{LT}^{-1}$ ] the hydraulic conductivity and  $\hat{e}_z$  the unit vector in  $z$ -direction. The spa-  
 164 tially variable hydraulic conductivity  $k(x, z)$  is modeled as a spatial random field, which is  
 165 generated from a multi-Gaussian random field  $f(x, z)$ . We consider both lognormally dis-

tributed  $k(x, z) = \exp[f(x, z)]$  as well as connected  $k(x, z)$  fields, which are generated by the transformation outlined in *Zinn and Harvey* [2003]. Details are given in Section 2.3.

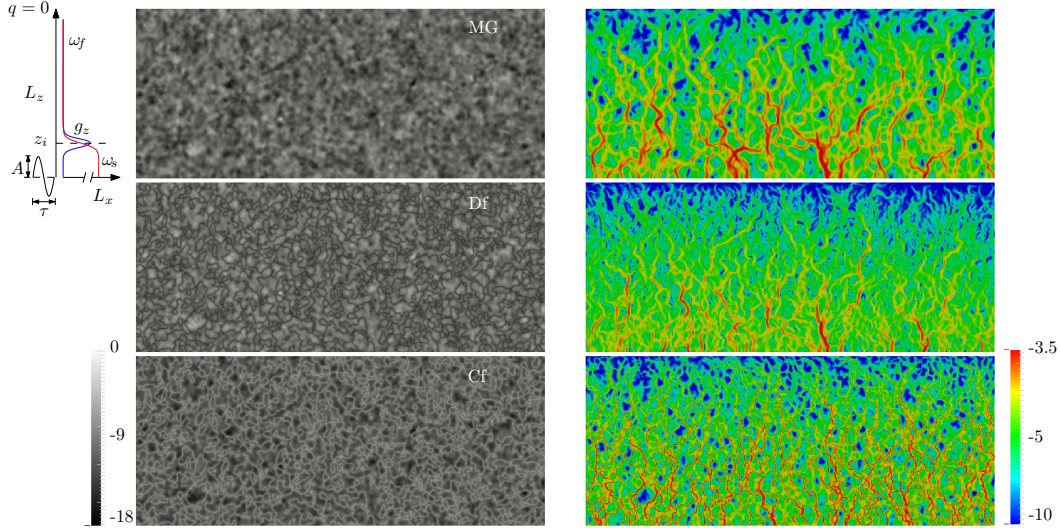
We consider flow scenarios with constant and variable fluid density. For the density-dependent flow scenarios, fluid density  $\rho$  is assumed to be a linear function of the salt mass fraction  $\omega$  (mass of dissolved salt per unit mass of fluid) given by  $\rho = \rho_f(1 + \beta c)$ , where  $\beta$  is the buoyancy factor defined as  $\beta = (\rho_s - \rho_f)/\rho_s$  with  $\rho_s$  and  $\rho_f$  the saline and freshwater densities, respectively, and  $c$  is the mixing ratio between the two fluids defined as  $c = \omega/\omega_s$ , with  $\omega_s$  the salt mass fraction of the saline water. Equation (1) is subject to a dynamic prescribed head assumed to be sinusoidal along the bottom boundary ( $z = 0$ ), characterized by a single harmonic with a constant amplitude ( $A$ ) and period ( $\tau$ ), which simulates for example tidal fluctuations; the top boundary is assumed to be located far from the prescribed harmonic flow perturbation so that the gradient of head can be considered to be zero (no-flow boundary of both salt and water) [*Pool et al.*, 2016]. This setup is characterized by a temporally mean flow of 0. Furthermore, the amplitude of the flow fluctuations decreases exponentially fast with distance from the boundary [see, e.g., *Jacob*, 1950; *Ferris*, 1951]. For an equivalent homogeneous medium, for example, the characteristic distance is given by the inverse of the wave number  $\mu = \sqrt{S_f \pi / (k_g \tau)}$ , where  $k_g$  is the geometric mean conductivity. Note that this implies, that even for a homogeneous medium, the flow field varies in space. Solute dispersion in transient flow in a homogenous medium under these conditions were studied in *Pool et al.* [2016].

The mixing ratio  $c$  evolves according to the advection-dispersion equation [e.g., *Pool et al.*, 2016]

$$\phi \frac{\partial c}{\partial t} = -\mathbf{q} \cdot \nabla c + \nabla \cdot \mathbf{D}_h \nabla c, \quad (4)$$

where local mixing is represented by a velocity-dependent dispersion coefficient  $\mathbf{D}_h$  [ $L^2 T^{-1}$ ] defined as  $\mathbf{D}_h(x, z, t) = \mathbf{D}(x, z, t) + \phi \mathbf{I} D_m$ , with  $D_m$  [ $L^2 T^{-1}$ ] the molecular diffusion coefficient,  $\mathbf{I}$  the identity matrix and  $\mathbf{D}$  [ $L^2 T^{-1}$ ] the dispersion tensor characterized by  $\alpha_l$  [L] and  $\alpha_t$  [L], the longitudinal and transverse dispersivities, respectively [*Bear*, 1972]. Equation (4) is obtained by combining fluid mass balance, eq. (1), and solute mass balance and further applying the Boussinesq approximation. The transport equation (4) is subject to a non-dispersion boundary condition at  $z=0$ . This implies that the mixing ra-

195 tio equals either that of the maximum ( $c=1$ ) for fluid inflows or that of the resident mass  
 196 fraction for fluid outflows [Voss and Provost, 2002; Frind, 1982]. For the initial condi-  
 197 tions, we consider an horizontal interface between the two fluids located at a specific dis-  
 198 tance  $z_i$  from the bottom boundary. Mixing and spreading are quantified from the gradient  
 199  $g_z(x, z, t) = -\partial c(x, z, t)/\partial z$  of the mixing ratio.



200 **Figure 1.** Example of permeability heterogeneous realization ( $\sigma_{\ln k}^2=4$ ) for the multi-Gaussian (MG),  
 201 connected (CF) and disconnected (DF) fields and boundary conditions (left) and spatial distribution of the  
 202 modulus of the velocity (m/s) for  $t = \pi/2$  (righth).

203 This setup may be compared to steady displacement of a solute front [Gramling  
 204 *et al.*, 2002], or displacement of an initial line source [Le Borgne *et al.*, 2010, 2015]. Note,  
 205 however, that a steady state flow scenario equivalent to the transient one under consider-  
 206 ation here is characterized by 0 flow velocity. An equivalent effective steady state trans-  
 207 port model is characterized by a transport velocity that decreases exponentially with dis-  
 208 tance and an effective dispersion coefficient [Pool *et al.*, 2016]. Thus, temporal fluctuations  
 209 of the flow boundary conditions are the drivers of the heterogeneity-induced dispersion,  
 210 mixing and reaction phenomena studied in the following. In the following subsection the  
 211 reactive transport problem is formulated and the expressions for the reaction rates under  
 212 mixing-limited conditions are given. Subsection 2.2 defines the measures to characterize  
 213 solute dispersion and mixing as well as flow deformation, which is are key processes for



214 the chemical reactions under consideration. Subsection 2.3 outlines the numerical method-  
 215 ology.

## 216 **2.1 Chemical reaction and reaction rate**

217 In this study, the geochemical setup of calcite dissolution/precipitation, which is very  
 218 common in variable-density flow systems such as coastal aquifers [see, e.g. *Sanford and*  
 219 *Konikow*, 1989; *Rezaei et al.*, 2005], is considered. The complete reaction system is de-  
 220 scribed in *De Simoni et al.* [2007]. This system can be characterized by two end-members,  
 221 freshwater and saline water both under chemical equilibrium conditions with calcite, and  
 222 its mixing ratio ( $c$ ). The reaction is assumed to be fast, this means equilibrium is attained  
 223 instantaneously. Within the reaction system, we focus on the rate of calcite dissolution and  
 224 precipitation. Note that we do not consider porosity and permeability changes in response  
 225 to dissolution or precipitation of calcite. Thus, no feedback between the chemical reaction  
 226 and the flow and transport properties is considered.

227 We follow the method proposed by *De Simoni et al.* [2005, 2007], which expresses  
 228 the equilibrium reaction rate in terms of the mixing ratio  $c$  between fresh and saline water  
 229 as

$$r_l(x, z, t) = \varphi_w \rho \frac{\partial^2 c_A}{\partial c^2} \nabla c(x, z, t) \cdot [\mathbf{D}_h(x, z, t) \nabla c(x, z, t)] \quad (5)$$

230 where  $\varphi_w$  is the mole fraction of water and  $c_A$  the concentration of secondary species  
 231 [*De Simoni et al.*, 2005]. In the freshwater-saline water mixing problem considered,  $c_A$  is  
 232 the molality of  $\text{Ca}^{2+}$  (mol/kg<sub>water</sub>) [*De Simoni et al.*, 2007]. The reaction rate depends on  
 233 chemistry through the speciation term  $\partial^2 c_A / \partial c^2$ , and on mixing as expressed by the sec-  
 234 ond term, which is identical to the mixing rate, as discussed below. It is controlled by the  
 235 gradient of the mixing ratio and the hydrodynamic diffusion-dispersion tensor. The speci-  
 236 ation term is determined by using PHREEQC [*Parkhurst*, 1995] for the end-member com-  
 237 positions given in Table 1 (fresh and saline water). The global reaction rate (moles that  
 238 precipitate/dissolve in order to keep equilibrium conditions) is defined by areal integra-  
 239 tion of the local reaction rate, ensemble average over all medium realizations and temporal  
 240 average over a period of the boundary fluctuations

$$r_g(t) = \left\langle \frac{1}{\tau} \int_0^\tau dt' \int_{\Omega} dx dz r_l(x, z, t + t') \right\rangle \quad (6)$$

241 with  $\Omega$  the aquifer domain; the angular brackets denote the ensemble average. The  
 242 variability of the reaction rate within the domain is characterized in terms of the probabil-  
 243 ity density function  $p_{r_l}(r)$  of the local reaction rate, which is obtained by spatial sampling  
 244 of  $r_l(x, z, t)$  in individual realizations and subsequent ensemble averaging of the resulting  
 245 single realization PDFs.

246 **Table 1.** Chemical composition of the end-members used in the calculations

Solution	pH	Ca	Mg	Na	K	Cl	log $P_{CO_2}$
Saline groundwater	7.21	9.64	22.43	496.53	9.28	564.13	-2.01
Freshwater	7.30	1.65	0.00	0.00	0.00	0.00	-2.00

\* Units for concentrations are in  $\text{mmol kg}_{\text{water}}^{-1}$

## 248 2.2 Dispersion and Mixing

249 We probe the impact of heterogeneity and density effects on the mixing and spread-  
 250 ing dynamics. Dispersion is characterized by the width of the mixing zone between the  
 251 two fluids, which is quantified here by the second-centered moment of the vertical compo-  
 252 nent of the gradient of the mixing ratio  $c$ . Mixing is characterized by the mixing rate [see,  
 253 e.g., *Kapoor and Kitanidis, 1998; Fiori and Dagan, 2002; Le Borgne et al., 2010; Bolster*  
 254 *et al., 2011*] and plays a key role for the dissolution/precipitation reaction under consid-  
 255 eration as outlined in the previous section. The local mixing dynamics are related qual-  
 256 itatively to flow deformation and flow topology in terms of the deformation rate tensor.  
 257 These quantities are defined in the following.

### 258 2.2.1 Interface width

259 The interface width characterizes the area between the two fluids and delineates the  
 260 region of potentially high mixing and reactivity, depending on the local dispersion coeffi-  
 261 cient and flow deformation, as discussed in the following subsections. The interface width  
 262 is quantified from the first and second moments of the gradient of the mixing ratio distri-  
 263 bution, and reads as

$$\sigma_z^2(x, t) = \{m_z^{(2)}(x, t) - m_z^{(1)}(x, t)m_z^{(1)}(x, t)\}, \quad (7)$$

264 where the first and second moments of  $g_z(x, z, t)$ ,  $m_z^{(1)}(x, t)$  and  $m_z^{(2)}(x, t)$  respec-  
265 tively, are given by

$$m_z^{(i)}(x, t) = \frac{1}{\int g_z(z, t|x) dz} \int g_z(z, t|x) z^i dz. \quad (8)$$

Due to the non-uniform flow field induced by the temporal flow fluctuations and the spatial heterogeneity, the second-centered moment of the  $z$ -component of the gradient of the mixing ratio distribution (7) varies along the  $x$ -axis. The effective width  $\sigma_e^2(t)$  is given by spatial, ensemble and temporal averaging of  $\sigma_z^2(x, t)$  as

$$\sigma_e^2(t) = \left\langle \frac{1}{\tau L_x} \int_0^\tau dt' \int_0^{L_x} dx \sigma_z^2(x, t + t') \right\rangle. \quad (9)$$

266 Note that we perform the temporal average over one period of the boundary fluctuation in  
267 order to emphasize the trend.

### 268 **2.2.2 Mixing rate**

269 The mixing rate, or scalar dissipation rate [*Tennekes and Lumley, 1972; Pope, 2000*]  
270 is defined as

$$\chi_l(x, z, t) = \nabla c(x, z, t) \cdot [\mathbf{D}_h(x, z, t) \nabla c(x, z, t)]. \quad (10)$$

271 The reaction rate (5) of the mixing-limited dissolution-precipitation reaction is directly  
272 proportional to the mixing rate. The mixing rate quantifies the local mixing mechanisms,  
273 namely the existence of concentration gradients and their dissipation by local dispersion.  
274 The variability of the mixing rate within the domain is characterized by its PDF  $p_{\chi_l}(\chi)$ ,  
275 which is obtained through spatial sampling of  $\chi_l(x, z, t)$  and subsequent ensemble aver-  
276 aging of the resulting single realization PDFs. The global mixing behavior is measured  
277 by the global mixing rate, which is obtained by areal integration, ensemble and temporal  
278 average as

$$\chi_g(t) = \left\langle \frac{1}{\tau} \int_0^\tau dt' \int_{\Omega} dx dz \chi_l(x, z, t + t') \right\rangle. \quad (11)$$

279 It is a measure of the diffusive flux across the interface.

In order to emphasize the role of (velocity-dependent) dispersion on the mixing, we contrast the mixing rate with the square of the gradient of the mixing ratio

$$\gamma_l(x, z, t) = \|\nabla c(x, z, t)\|^2, \quad (12)$$

280 where  $\|\cdot\|$  denotes the  $L^2$ -norm.  $\gamma_l(x, z, t)$  measures the strength of local gradients of  
 281  $c(x, z, t)$ . Unlike for scenarios with constant dispersion coefficients, here the occurrence  
 282 of high concentration gradients does not imply a high mixing rate. The global square gra-  
 283 dient  $\gamma_g(t)$  is defined in analogy to (11).

For the displacement of a solute front in homogeneous porous medium under con-  
 stant flow velocity and constant diffusion coefficient  $D$ , the global mixing rate is  $\chi_g(t) =$   
 $\sqrt{D}/(2\sqrt{2\pi t})$ . Thus,  $\chi_g$  decays with time due to the diffusive smoothing of the interface  
 between the displacing and the displaced fluids. The square gradient is simply  $\gamma_g(t) =$   
 $\chi_g(t)/D$ . For displacement under unsteady flow conditions in a homogeneous porous  
 medium [Pool *et al.*, 2016], the global mixing rate is

$$\chi_g(t) = \frac{\mathcal{D}(t) + D_m}{2\sqrt{\pi\sigma_0^2(t)}}, \quad (13)$$

where  $\mathcal{D}(t)$  and  $\sigma_0^2(t)$  is given by expressions (36) and (42) in Pool *et al.* [2016]. As  
 shown there, the gradient of the mixing ratio can be approximated by a Gaussian distri-  
 bution. Thus, the PDF of local mixing rates  $p_{\chi_l}(\chi)$  is given by

$$p_{\chi_l}(\chi) = \frac{1}{2\chi\sqrt{\ln(\chi_m/\chi_0)\ln(\chi_m/\chi)}}, \quad (14)$$

284 where  $\chi_m = (\mathcal{D} + D_m)/(2\pi\sigma_0^2)$  is the maximum mixing rate at a given time and  $\chi_0$  a  
 285 minimum mixing rate. The square gradient  $\gamma_l$  is related to the mixing rate by rescaling as  
 286  $\gamma_l = \chi_l/(\mathcal{D} + D_m)$ . Therefore,  $\gamma_l$  and  $\chi_l$  differ in their maxima but their spatial distribution  
 287 is the same.

### 2.2.3 Flow Deformation

288  
 289 The local deformation properties of the flow field are directly related to the mixing  
 290 and therefore reaction behavior [de Barros *et al.*, 2012; Le Borgne *et al.*, 2013, 2014, 2015;

291 *Bandopadhyay et al.*, 2017]. Stretching and compression of the local fluid support leads to  
 292 a steepening of concentration gradients, which in the presence of local dispersion give rise  
 293 to enhanced mixing and thus reaction. The deformation of the flow field is characterized  
 294 by the deformation rate tensor [*Okubo*, 1970a; *de Barros et al.*, 2012], which describes the  
 295 local strain and rotation properties of the flow field. It is defined by [*Ottino*, 1989]

$$\epsilon(x, z, t) = \begin{pmatrix} \frac{\partial v_x(x, z, t)}{\partial x} & \frac{\partial v_x(x, z, t)}{\partial z} \\ \frac{\partial v_z(x, z, t)}{\partial x} & \frac{\partial v_z(x, z, t)}{\partial z} \end{pmatrix} \quad (15)$$

296 with  $\mathbf{v}(x, z, t) = \mathbf{q}(x, z, t)/\phi$  the pore water velocity. The vorticity is defined by  
 297  $\omega = \epsilon_{21} - \epsilon_{12}$ , the normal strain and shear rates are defined by  $\alpha = 2\epsilon_{11}$  and  $\sigma = \epsilon_{21} + \epsilon_{12}$ ,  
 298 respectively. The Okubo-Weiss parameter [*Okubo*, 1970b; *Weiss*, 1991] is defined in terms  
 299 of the negative determinant of the deformation gradient tensor  $\Theta(x, z, t) = -4\det[\epsilon(x, z, t)]$ .  
 300 It can be decomposed into  $\Theta = \Theta_\zeta^2 - \omega^2$ , where  $\Theta_\zeta = \alpha^2 + \sigma^2$  is a measure for local stretch-  
 301 ing. Thus, positive values of  $\Theta$  are associated with local stretching deformation, negative  
 302 values with rotation of a fluid element. A spatial map of  $\Theta(x, z, t)$  delineates regions of  
 303 predominant stretching or rotation action of the flow field [*Okubo*, 1970b; *Weiss*, 1991].  
 304 Regions characterized by strong stretching action can be delineated by  $\Theta_\zeta(x, z, t)$ , which  
 305 is termed rate of strain in the following. Vorticity has only a minor impact on mixing and  
 306 mass transfer because it merely describes the rotation of a material element [*de Barros*  
 307 *et al.*, 2012]. Strain deformation on the other hand affects the mixing and mass transfer  
 308 properties through stretching or compression of material elements and subsequent steep-  
 309 ending or attenuation of concentration gradients. As the interface passes through a re-  
 310 gion of high  $\Theta_\zeta$ , it gets strongly stretched, which steepens the gradient of the mixing ra-  
 311 tio [*de Barros et al.*, 2012]. If local scale dispersion is sufficiently strong, stretching leads  
 312 to enhanced mixing as expressed by the mixing rate (10).

### 313 **2.3 Numerical Methodology**

314 The impact of heterogeneity on mixing and chemical reactions is investigated by per-  
 315 forming several sets of Monte Carlo simulations. Spatial heterogeneity in the hydraulic  
 316 conductivity is characterized by multi-Gaussian lognormally distributed random fields.  
 317 Medium realizations are generated by using the GCOSIM3D code [*Gomez-Hernandez and*  
 318 *Journal*, 1992]. The geometric mean of hydraulic conductivity ( $k_g$ ) for all the simulations

319 is set to  $1.15 \cdot 10^{-4}$  m/s. Two different values for the log-conductivity variance are consid-  
 320 ered ( $\sigma_{\ln k}^2 = 1$  and 4) and directional correlation in the two spatial dimensions is assumed  
 321 to be isotropic with correlation length  $\lambda = 10$  m. A model domain of extent  $L_x \times L_z$  is  
 322 considered, with  $L_x = 10^2 \lambda$  and  $L_z = 40 \lambda$ , discretized in grid cells of size  $d_x \times d_z$ , with  
 323  $d_x = \lambda/10$  and  $d_z = \lambda/10$ . Moreover, in order to study the effect of channelling patterns,  
 324 commonly found in sedimentary systems, on mixing and chemical reactions more complex  
 325 heterogeneous media are considered by following the transformation of the multi-Gaussian  
 326 fields proposed by *Zinn and Harvey* [2003]. From this transformation [see also *Knudby*  
 327 *and Carrera*, 2005], we obtain heterogeneous fields characterized by low-conductivity  
 328 paths (Df, disconnected fields) or high-conductivity paths (Cf, connected fields), see Fig-  
 329 ure 1. As a criterion for the size of the numerical ensembles, we have chosen the stability  
 330 of the average of the spatial variance of the front width. The minimum number of realiza-  
 331 tions to achieve this turned out to be 30 for the multi-Gaussian fields and 40 for the con-  
 332 nected fields. The effective values of the previously described observables are computed  
 333 by performing the average over these realizations.

334 Furthermore, in order to examine the effect of variable fluid density on mixing and  
 335 reactions, both constant-density and variable-density saturated flow and transport simula-  
 336 tions were performed. Coupled flow and transport equations are solved by using a 2nd-  
 337 order finite volume and 6th-order compact finite differences code [see details in *Hidalgo*  
 338 *et al.*, 2013]. Values for  $\alpha_l$  and  $\alpha_t$  of 2 m and 0.2 m are considered, which are consis-  
 339 tent with those estimated for real systems of this scale [see, e.g. *Gelhar et al.*, 1992; *Zech*  
 340 *et al.*, 2015]. Time is scaled with the characteristic advection time defined by  $\tau_a = A/v_z$   
 341 with  $v_z$  the maximum velocity at the boundary for an equivalent homogeneous system  
 342 considering  $k_g$  given by  $v_z = \sqrt{2} A k_g \mu / \phi$  [see, e.g., *Jacob*, 1950; *Ferris*, 1951], where  
 343  $\mu = \sqrt{S_f \pi / (k_g \tau)}$  represents the inverse of the characteristic distance for the attenuation of  
 344 the temporal head fluctuations in the domain. In other words,  $\gg \mu^{-1}$ , denotes the range of  
 345 the periodic head fluctuations.

346 Table 2 provides the employed values for the flow and transport simulations and  
 347 the statistical properties of the heterogeneous fields. Figure 1 shows the model geome-  
 348 try, boundary conditions and (left) the hydraulic conductivity distribution for one random  
 349 heterogeneous multi-Gaussian realization (MG) and the corresponding transformed discon-  
 350 nected (Df) and connected (Cf) fields after the procedure used by *Zinn and Harvey* [2003]

351 and *Knudby and Carrera* [2005] and (right) the spatial distribution of the modulus of the  
 352 velocity (m/s) for  $t = \pi/2$ .

353 **Table 2.** Parameters used in numerical simulations.

Parameter	Value	Description
$L_x$ [m]	1000	Domain x length
$L_z$ [m]	400	Domain z length
$z_i$ [m]	100	Interface location
$k_g$ [m/s]	$1.15 \cdot 10^{-4}$	geometric mean permeability
$S_f$ [ $m^{-1}$ ]	$3 \cdot 10^{-5}$	Specific storage coefficient
$\sigma_{\ln k}^2$	1,4	variance of the log-hydraulic conductivity
$\lambda$ [m]	10	Correlation length
$\phi$ [-]	0.25	Porosity
$\alpha_L$ [m]	2	Longitudinal dispersivity
$\alpha_T$ [m]	0.2	Transverse dispersivity
$D_m$ [ $m^2/s$ ]	$1 \cdot 10^{-9}$	Molecular diffusion coefficient
$A$ [m]	1	Oscillation amplitude
$\tau$ [s]	43200	Oscillation period

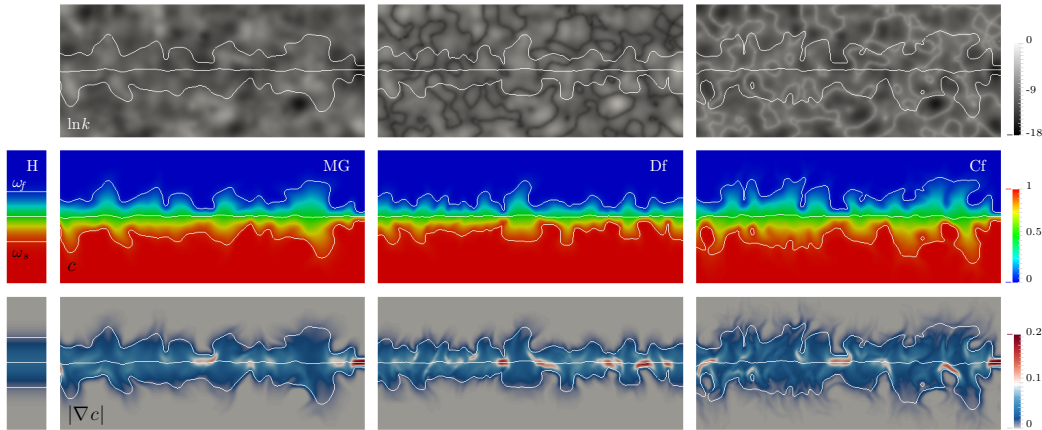
### 354 **3 Results and Discussion**

355 In the following, we discuss the impact of spatial heterogeneity and connectivity of  
 356 hydraulic conductivity, and the role of buoyancy on the dispersion mixing and reaction  
 357 behaviors. To this end, we first focus on the impact of heterogeneity and connectivity for  
 358 constant fluid density. The obtained behaviors are then compared to the behaviors in the  
 359 presence of fluid density variations.

#### 360 **3.1 Spatial Heterogeneity and Connectivity**

361 We study the impact of spatial heterogeneity and connectivity on the dispersion of  
 362 the interface between the displacing and displaced fluids, the mixing of the two fluids,  
 363 and chemical reactivity. Figure 2 illustrates the conductivity distribution for the different  
 364 heterogeneity scenarios under consideration as well as the distribution of the mixing ratio

365 and magnitude of its gradient, which delineates the interface. We consider global behavior  
 366 of dispersion, mixing and reaction measures defined in Section 2. The spatial variability  
 367 of mixing and reaction is quantified by the PDFs of the local mixing and reaction rates.  
 368 The occurrence of mixing and reaction patterns is studied by spatial maps of the local  
 369 reaction and mixing rates and related to the medium and flow structure through the map  
 370 of local strain rates.



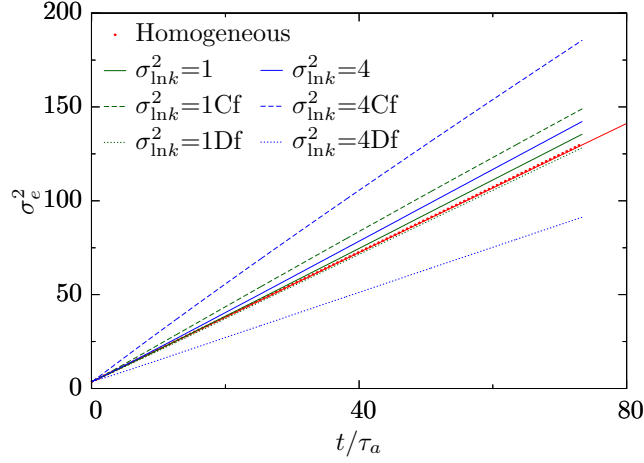
371 **Figure 2.** Example of a zoomed-in portion of the hydraulic conductivity distribution ( $\ln k$ ), snapshot of  
 372 the mixing ratio distribution ( $c$ ) and its gradient for the homogeneous case (H) and one heterogeneous multi-  
 373 Gaussian (MG), disconnected (Df) and connected (Cf) realization ( $\sigma_{\ln k}^2=4$ ) at time  $t' = t/\tau_a = 74$ .

### 374 **3.1.1 Dispersion**

375 Periodic temporal forcing leads to a noninstantaneous flow field, i.e. space and time-  
 376 dependent flow response, in which the oscillation amplitude decays exponentially with  
 377 distance from the boundary whereas the phase lag increases linearly [see, e.g., *Jacob*,  
 378 1950; *Van Der Kamp*, 1972; *Townley*, 1995]. The resulting time-dependent velocity field  
 379 enhances solute dispersion even if the medium is homogeneous [see, e.g. *Kinzelbach and*  
 380 *Ackerer*, 1986; *Goode and Konikow*, 1990; *Pool et al.*, 2016].

384 Figure 3 shows the temporal evolution of the second-centered moment ( $\sigma_c^2(t)$ ) of  
 385 the gradient of the mixing ratio for different values of the log-conductivity variance ( $\sigma_{\ln k}^2$ )  
 386 and different degrees of connectivity. Results for the equivalent homogeneous case [*Pool*  
 387 *et al.*, 2016], which is characterized by the same dispersivities, storativity and porosity  
 388 as the heterogeneous model and with a unique hydraulic conductivity equal to  $k_g$ , are in-





381 **Figure 3.** Temporal behavior of  $\sigma_e^2(t)$  for different values of  $\sigma_{\ln k}^2$  for the multi-Gaussian, connected (Cf)  
 382 and disconnected (Df) fields. The red line represents the results obtained from the analytical solution derived  
 383 in *Pool et al.* [2016] for illustration.

389 cluded for comparison. Numerical results show that, as expected, for the multi-Gaussian  
 390 fields the second-centered moment of the gradient of the mixing ratio distribution  $\sigma_e^2(t)$   
 391 slightly increases with the log-conductivity variance. Temporal oscillations combined with  
 392 high degree of connectivity substantially widen the mixing zone between the fresh and  
 393 saline water, an effect that is significantly enhanced for increasing  $\sigma_{\ln k}^2$ . Thus, for moder-  
 394 ately heterogeneous connected permeability fields ( $\sigma_{\ln k}^2=1$ ),  $\sigma_e^2(t)$  is much larger than that  
 395 for highly heterogeneous multi-Gaussian permeability fields ( $\sigma_{\ln k}^2=4$ ). On the other hand,  
 396 for the disconnected fields,  $\sigma_e^2(t)$  is smaller than the one for the equivalent homogeneous  
 397 case, and decreases when the log-conductivity variance increases. These results may be  
 398 attributed to the impact of connectivity on the effective conductivity. While for the multi-  
 399 Gaussian fields the effective conductivity corresponds to the geometric mean [see, e.g.  
 400 *Matheron, 1967; Dagan, 1989*], for the connected fields it is higher than  $k_g$  and increases  
 401 with  $\sigma_{\ln k}^2$  whereas for the disconnected fields it is lower and decreases with  $\sigma_{\ln k}^2$  [*Zinn and*  
 402 *Harvey, 2003*]. For the disconnected fields, an increase in  $\sigma_{\ln k}^2$  promotes the emergence  
 403 of regions of disconnected high velocity embedded in low-velocity channels, which leads  
 404 to a decrease in the effective conductivity and then in the transient-induced velocity. As  
 405 a result, the effect of the temporal fluctuations on solute dispersion decreases leading to a  
 406 smaller width of the mixing zone. However, for the connected fields the increase in  $\sigma_{\ln k}^2$   
 407 increases the effective conductivity and leads to flow fields characterized by a continuous

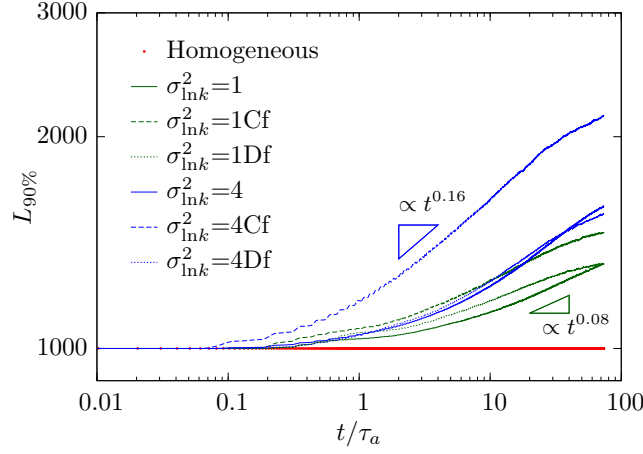
408 network of high velocity paths. The strong velocity variations induced by these channeling  
409 patterns and the periodic temporal fluctuations lead to a significant deformation and elon-  
410 gation of the interface in flow direction such that the width of the mixing zone increases.

411 The deformation of the interface can be quantified in terms of the temporal change  
412 of the length of a particular concentration isoline. The interface length is a measure of  
413 deformation on one hand and the length of the contact line between the displacing and  
414 displaced fluids. Thus, it is both an alternative measure of dispersion and serves for the  
415 quantification of the mixing efficiency. Unlike in fluids, however, where line length, or  
416 its rate of increase (Lyapunov exponents) is directly used as a measure for mixing, here  
417 the local values of dispersion along the interface play a key role in the quantification of  
418 mixing as discussed below.

419 Figure 4 shows the temporal evolution of the 90% isoconcentration contour line  
420 for the multi-Gaussian, connected and disconnected fields for different values of the log-  
421 conductivity variance. The elongation of this specific isoline evolves as a power law of  
422 time with the exponent increasing with the log-conductivity variance [*Le Borgne et al.*,  
423 2013]. Note that for uniform steady flow in heterogeneous media, one would expect an  
424 initially linear growth of the interface length caused by persistent flow velocities. For the  
425 transient flow scenario under consideration, however, the velocity magnitude decreases in  
426 the direction of the flow [*Pool et al.*, 2016], which then leads to a slower increase of the  
427 interface length. At distances from the inlet boundary that are smaller than the correla-  
428 tion length, segments of the interface move as in 1-dimensional homogeneous media. The  
429 evolution of each segment can be described by the solution for a homogeneous medium as  
430  $z_0(t) = \ln(1+t/\tau_v)/2\mu$ , where  $\tau_v = 2\pi/v^2\mu^2\tau$ , see Eq. (30) in [*Pool et al.*, 2016]. Thus, the  
431 initial elongation, the result of persistent velocity contrasts along the interface, i.e., differ-  
432 ent  $v$ , grows approximately as  $z_0(t)$ . At later times, velocities decorrelate and the interface  
433 evolves sublinearly. Note that, although the deformation of the interface is significantly en-  
434 hanced by the degree of connectivity and grows faster, similar exponents are obtained for  
435 the multi-Gaussian and for the connected fields.

### 438 3.1.2 Mixing rate

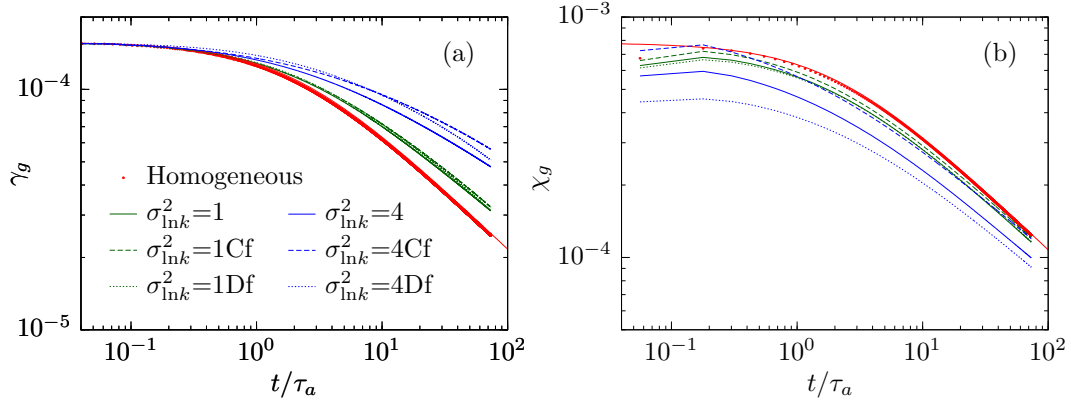
439 The temporal evolution of the square of the mixing ratio gradient and the global  
440 mixing rate obtained by averaging over the period is illustrated in Figure 5. Results reveal



436 **Figure 4.** Temporal evolution of the length of the 90% isoconcentration contour line for the multi-Gaussian,  
 437 connected (Cf) and disconnected (Df) fields considering different values of the log-conductivity variance.

441 that heterogeneity significantly impacts on the concentration gradients. Spatial variability  
 442 of the hydraulic conductivity leads to local velocity variations which promotes segrega-  
 443 tion and the steepening of concentration gradients [Le Borgne *et al.*, 2013]. This effect,  
 444 which is magnified by the degree of heterogeneity and connectivity, causes the destruction  
 445 of the gradients to significantly slow down compared to the homogeneous case, see Fig-  
 446 ure 5a. For constant dispersion coefficient, this result would suggest a drastic increase in  
 447 the mixing rate with heterogeneity as it increases with the square of the local concentra-  
 448 tion gradients. However, for the velocity-dependent dispersion tensor under consideration  
 449 here, we find that the global mixing rate is (i) generally smaller than the one for the ho-  
 450 mogeneous scenario, and (ii) decreases for increasing heterogeneity in all cases, see Figure  
 451 5b. This behavior is due to the velocity dependence of dispersion. The largest gradients  
 452 are located at low conductivity zones where velocities and thus dispersion are very small,  
 453 see Figure 2. Thus, while heterogeneity-induced segregation leads to a strong increase of  
 454 concentration gradients, it lowers the global mixing efficiency due to the lack of dispersion  
 455 in the regions of the highest gradients. As shown in Figure 2, concentration gradients are  
 456 also enhanced in the vicinity of high conductivity zones due to strong velocity variations,  
 457 which can be clearly seen for the connected fields. As discussed below this increase of the  
 458 mixing ratio gradients is linked to increased strain deformation. Thus, as local dispersion  
 459 is significantly enhanced in high conductivity zones this leads to high local mixing rates.  
 460 However, due to the strong spatial segregation, spatial heterogeneity has less of an impact  
 461 on the global mixing rate, but may have a significant impact on the spatial distribution and

462 variability of local mixing rates, and leads to localization in regions of high mixing rate as  
 463 discussed below.



464 **Figure 5.** Temporal behavior of the square of the mixing ratio gradient and the mixing rate for the multi-  
 465 Gaussian, connected (Cf) and disconnected (Df) fields for different values of the log-conductivity variance.

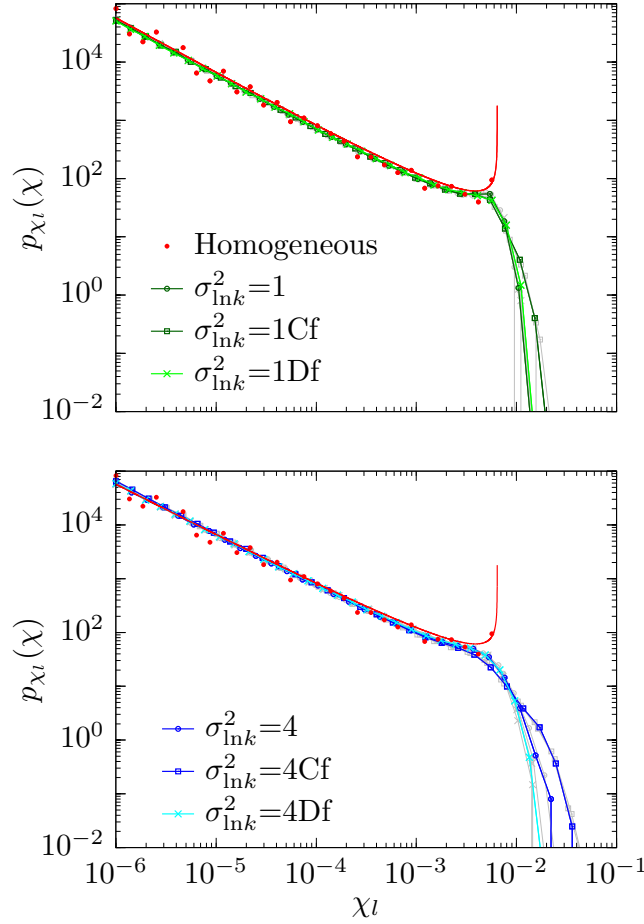
466 The spatial variability of the mixing rate induced by structural heterogeneity can be  
 467 characterized by the PDF  $p_{\chi_l}(\chi)$  of its local values. As described in Section 2.2.2, the  
 468 PDF is obtained by spatially sampling all values in the model domain that are above or  
 469 equal to a specific threshold ( $\chi_l \geq 10^{-6}$ ). Figure 6 shows  $p_{\chi_l}(\chi)$  at time  $t = 74\tau_a$ , which  
 470 coincides with a maximum of the head signal at the boundary. Note that the behaviors of  
 471 the PDF of the temporally averaged reaction rates (not shown) are qualitatively similar to  
 472 ones at the maximum time. Also, at this time the global mixing rates for the connected  
 473 and homogeneous scenarios have similar values (see Figure 5b). The obtained distribution  
 474 of local mixing rates for the heterogeneous realizations, however, is much broader for the  
 475 heterogeneous than the for the homogeneous case, which reflects the spatial heterogeneity  
 476 in flow velocity and dispersion. The PDF for the homogeneous case, Eq. (14), behaves as  
 477  $\chi^{-1}$  at small values of the mixing rate, which is characteristic of a Gaussian shaped spa-  
 478 tial distribution. At higher values there is a cut-off at the maximum concentration, where  
 479  $p_{\chi_l}(\chi)$  has an integrable singularity. The PDFs for the heterogeneous cases behave as  $\chi^{-1}$   
 480 at small mixing rates, which indicate a quasi-Gaussian behavior of the spatial distribu-  
 481 tion of  $\chi_l$  at least for small values of  $\chi_l$ . However, the range of  $\chi$  is much larger for the  
 482 heterogeneous scenarios and the behavior at large  $\chi$  is very different. Instead of a singu-  
 483 lar behavior like in the homogeneous case,  $p_{\chi_l}(\chi)$  decreases smoothly to zero. This type  
 484 of behavior can be explained by a broad distribution of local maxima of the mixing rate

485 as illustrated in Figure 9. Thus, the PDF of mixing ratios  $p_{\chi_l}(\chi)$  may be understood as  
486 a superposition of local PDFs corresponding to (14) for different (local) maximum mix-  
487 ing ratios  $\chi_m$ . This argument will be explored in future work. Note that as the degree of  
488 connectivity increases, the occurrence of high values of the mixing ratio increases. This  
489 increase, which reflects high dispersive mass transfer processes, demonstrates that spatial  
490 variability of hydraulic conductivity and especially channelling patterns lead to locally en-  
491 hanced mixing. Therefore, while the global behavior of the mixing rate is not informative  
492 on the mixing behavior in the different heterogeneity scenarios, the PDF of local mixing  
493 ratios, which play a key role in the effective reaction efficiency, reflects the local mixing  
494 mechanisms which are controlled by the local spatial structure of the velocity field. The  
495 global mixing rate for the heterogeneous cases may be smaller than the one for the ho-  
496 mogeneous case due to segregation of the spatial distribution of the mixing ratio, but the  
497 occurrence of high local values increases, which is an indicator for the creation of mixing  
498 and reaction hotspots as discussed below.

### 504 **3.1.3 Reaction rate**

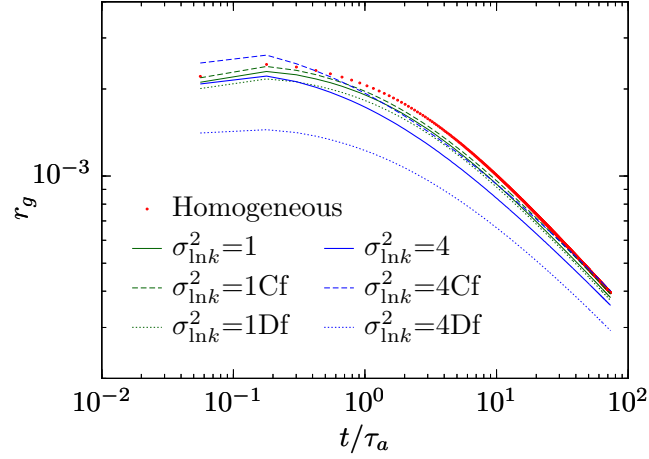
505 In order to quantify the effect of heterogeneity and connectivity on chemical reac-  
506 tions, we first consider the global reaction rate given by equation (6). In order to compute  
507 the local reaction rates, the second derivative of  $c_A$  with respect to the mixing ratio  $c$  is  
508 evaluated as described in Section 2.1.

511 The behavior of the reaction rate for the different heterogeneity scenarios is very  
512 similar to those observed for the mixing rate in the previous section, which is expected  
513 due to the strong relation between mixing and reaction rates expressed in Eq. (6). Figure  
514 7 shows the temporal evolution of the global reaction rate for different values of  $\sigma_{\ln k}^2$  and  
515 different degrees of connectivity. Similar to the global mixing rate, Figure 5b, numerical  
516 results show that the global reaction efficiency is higher for the homogeneous than those  
517 for all of the heterogeneous scenarios. Moreover, the global reaction efficiency decreases  
518 with increasing  $\sigma_{\ln k}^2$ , which again can be traced back to the increase of segregation and  
519 the lack of dispersion in low conductivity and thus velocity regions. As for the mixing  
520 rate, this segregation on one hand leads to a reduction of the global reaction rate, on the  
521 other it leads to a broadening of the distribution of local reaction rates with localization of  
522 of high reaction rates in regions of high dispersive mass transfer.



499 **Figure 6.** PDF of the local mixing rate for the multi-Gaussian, connected (Cf) and disconnected (Df) fields  
 500 considering different values of the log-conductivity variance at time  $t' = t/\tau_a = 74$ , which corresponds to  
 501 the maximum of the head signal at the boundary. The grey symbols denote corresponding PDFs in individual  
 502 realizations, which differ only slightly from the ensemble PDF. The red solid line represents predictions from  
 503 the equation (14).

523 The PDF  $p_{r_l}(r)$  provides a measure for the variability of local reaction rates  $r_l(x, z, t)$ .  
 524 Figure 8 shows  $p_{r_l}(r)$  computed at time  $t = 74\tau_a$  for the multi-Gaussian, connected and  
 525 disconnected fields for different values of  $\sigma_{\ln k}^2$ . This time has been chosen because it co-  
 526 incides with a maximum of the head signal at the boundary, at which the impact of fluc-  
 527 tuations is maximum. The behaviors of the PDF of the temporally averaged reaction rates  
 528 are qualitatively similar to ones at the maximum time. The results suggest that, although  
 529 at this specific time the global reaction rate for the homogeneous case is found to be sim-  
 530 ilar as those for the connected scenarios, see Figure 7, the range of local reaction rates  
 531 is significantly enhanced by the degree of heterogeneity and connectivity. The behavior



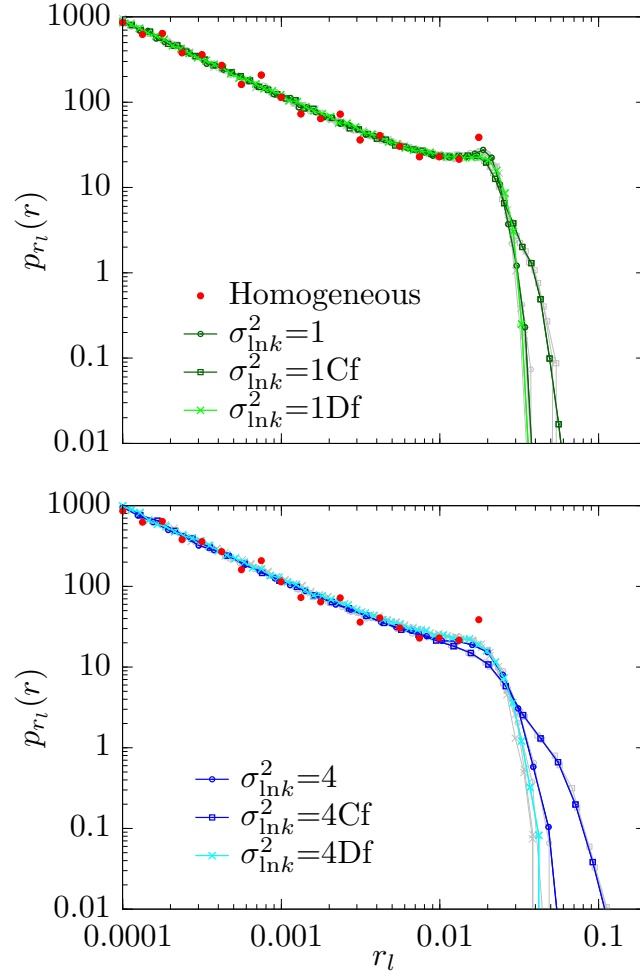
509 **Figure 7.** Temporal behavior of the global reactivity for the multi-Gaussian, connected (Cf) and discon-  
 510 nected (Df) fields considering different of the log-conductivity variance.

532 is similar to the one observed for the PDF of the mixing rates in the previous section.  
 533 At small reaction rates, we observe a similar  $1/r$  decay that can be attributed to a Gaus-  
 534 sian shape of the spatial distribution of  $\chi_l$  and thus of  $r_l$ . The behavior at large  $r$  is again  
 535 due to a broad distribution of maximum reaction rates in the heterogeneous medium. The  
 536 maximum reactions rates for all the heterogeneous realizations are significantly higher than  
 537 those for the homogeneous case, up to over one order of magnitude larger for the strongly  
 538 heterogeneous connected fields. Therefore, the numerical results demonstrate that the ef-  
 539 fects of heterogeneity and connectivity on mixing lead to a strongly heterogeneous distri-  
 540 bution of local mixing rates, which tends to increase locally the potential mixing and as a  
 541 consequence reactivity of the system. In the following, we consider the spatial organiza-  
 542 tion of the local mixing and reaction rates.

#### 547 **3.1.4 Mixing, reaction and deformation patterns**

548 In order to identify and delineate regions of locally increased mixing reaction po-  
 549 tential, we analyze and quantify the deformation of the flow field as described in Sec-  
 550 tion 2.2.3 in terms of the rate of strain  $\Theta_\zeta(x, z, t)$ . The strain rate is expected to increase  
 551 in areas of high velocity, which together with enhanced dispersion leads to higher mixing  
 552 and thus reaction rates.

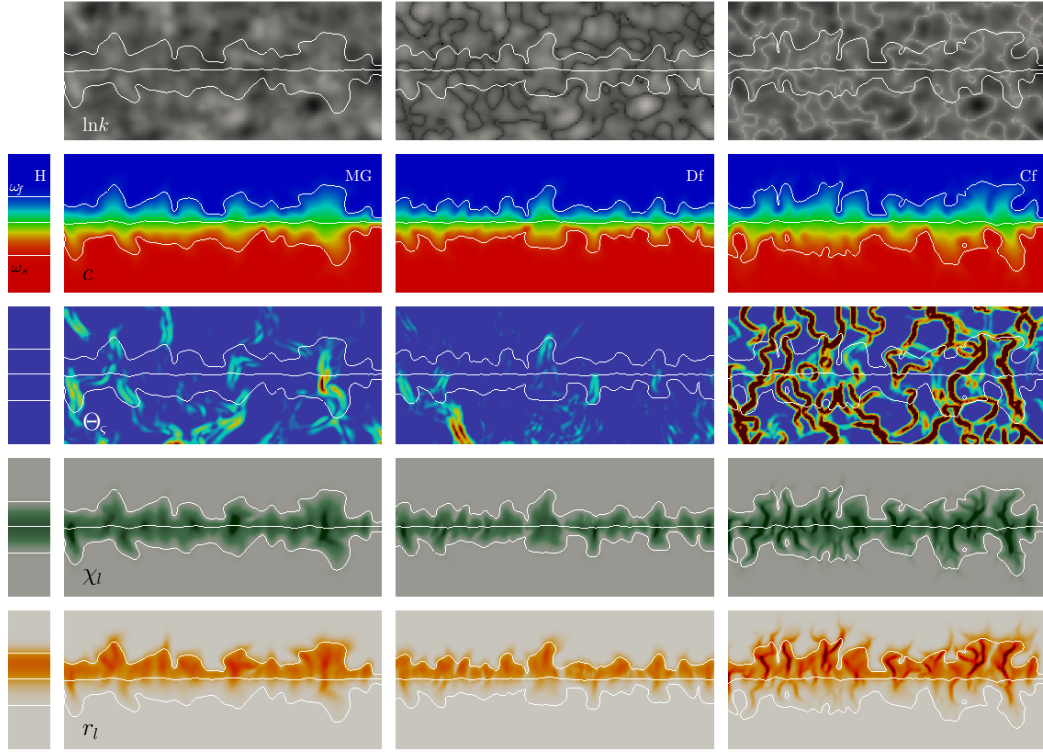
558 Figure 9 displays a zoomed-in example of, from top to bottom, the hydraulic con-  
 559 ductivity and mixing ratio distributions, the strain rate ( $\Theta_\zeta$ ), and the local scalar dissipa-



543 **Figure 8.** Reaction rate probability density function (PDF) for the multi-Gaussian, connected (Cf) and  
 544 disconnected (Df) fields considering different of the log-conductivity variance at time  $t' = t/\tau_d = 74$ , which  
 545 corresponds to the maximum of the head signal at the boundary. The grey symbols denote corresponding  
 546 PDFs in individual realizations, which differ only slightly from the ensemble PDF.

560 tion ( $\chi_l$ ) and reaction rate distributions ( $r_l$ ) for the homogeneous case and multi-Gaussian,  
 561 connected and disconnected fields ( $\sigma_{\ln k}^2=4$ ) at time  $t = 74\tau_d$ . Note that the strain deformation  
 562 for the homogeneous field is approximately zero here. This is because, although the  
 563 storativity is non-zero and then a spatially nonuniform time-dependent velocity field in re-  
 564 sponse to the fluctuations of the lower boundary condition is expected to occur, it is small  
 565 enough to lead to a practically instantaneous and therefore spatially uniform flow response.  
 566 Figure 9 reveals that the strain deformation is enhanced by heterogeneity and significantly  
 567 increases with the degree of connectivity. Connectivity leads to spatial velocity gradients,  
 568 which strongly promote the emergence of connected patterns of high deformation. This





553 **Figure 9.** From top to bottom, example of a zoomed-in portion of the  $\ln k$  for one conductivity hetero-  
 554 geneous realization for the homogeneous (H), multi-Gaussian (MG), disconnected (Df) and connected (Cf)  
 555 fields, snapshot of the mixing ratio distribution, local values for the  $\Theta_S$  parameter and the local scalar dissipa-  
 556 tion  $\chi_l$  and reaction rates  $r_l$  distribution at time  $t' = t/\tau_a = 74$ . The white lines represent normalized salinity  
 557 contours (5, 50 and 95%)

569 leads to an elongation of the contact surface between the two fluids, and steepening of  
 570 mixing ratio gradients in the vicinity of the regions of high conductivity where disper-  
 571 sion is high. As a result, the local mixing and reaction rates are significantly enhanced.  
 572 Thus, the maximum reaction efficiency and thus dissolution occurs in the regions of high-  
 573 est strain, which are aligned with the connected channels of high conductivity.

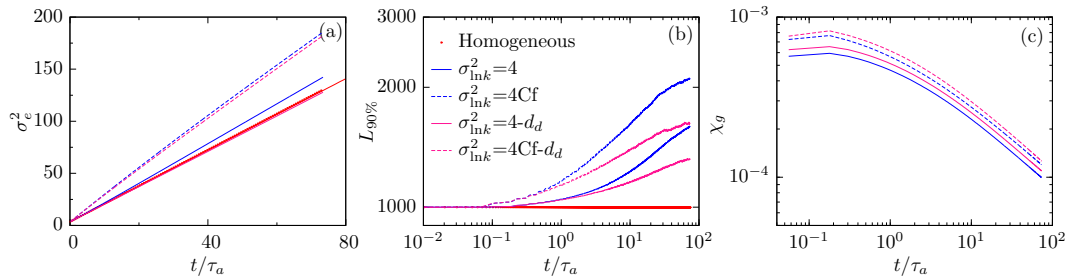
574 Hotspots of geochemical reactivity are clearly related to the deformation of the flow  
 575 field induced by spatial heterogeneity and the periodic temporal fluctuations. Therefore,  
 576 evaluation of the deformation properties and topology of the flow field allow to localize  
 577 and predict regions of enhanced geochemical reaction efficiency. Our results also show  
 578 that reaction rates are maximal in the freshwater end of the interface [see also *Rezaei*  
 579 *et al.*, 2005], particularly between the 5 and 50% salinity contour lines. This behavior is

580 due to the speciation term  $\partial^2 c_A / \partial c^2$  in Eq. (6) which is maximum for small mixing ra-  
 581 tio ( $c$ ) between the two end-members, and decreases asymptotically with the proportion  
 582 of saline water [see Figure 5b in *De Simoni et al.*, 2007]. Therefore, a mixing ratio with  
 583 a small fraction of saline water increases significantly the geochemical reactivity of the  
 584 system.

585 Note also that for the homogeneous case a perfectly horizontal reactive front is ob-  
 586 tained, whereas the combined effect of spatial variability and temporal fluctuations allow  
 587 mixing and geochemical reaction patterns to emerge. Thus, irregular reaction conduit net-  
 588 works are obtained, which are characterized by vertical structures, localized where flow  
 589 velocities are relatively high, and connected to a main horizontal bedding plane.

### 590 3.2 Impact of density variations

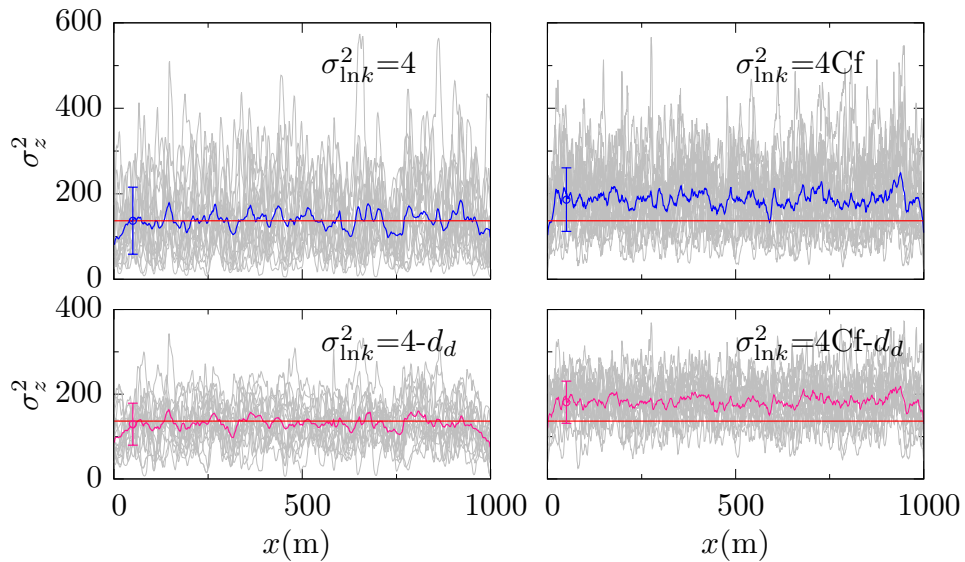
591 The dispersion of non-reactive tracers in density-dependent groundwater flow sys-  
 592 tems has been investigated experimentally, theoretically and numerically in a series of  
 593 studies [see, e.g. *Hassanizadeh and Leijnse*, 1988; *Schotting et al.*, 1999; *Johannsen et al.*,  
 594 2002; *Landman et al.*, 2007a]. These works found that density variations may lead to an  
 595 increase or a reduction of dispersion, which is controlled by the ratio of viscous to buoy-  
 596 ancy forces.



597 **Figure 10.** Density-constant and -dependent ( $d_d$ ) results for the temporal evolution of (a) the second-  
 598 centered moment of the mixing ratio gradient in  $z$ -direction, (b) the length of the 90% isoconcentration  
 599 contour line and (c) the scalar dissipation rate for the homogeneous, multi-Gaussian (MG) and connected (Cf)  
 600 fields ( $\sigma_{lnk}^2 = 4$ ).

601 The impact of density variations on the temporal evolutions of  $\sigma_\ell^2(t)$ , the interface  
 602 length, and the global mixing rate  $\chi_g(t)$  for the multi-Gaussian and connected fields is il-  
 603 lustrated in Figure 10. Although the observed behaviors are qualitatively similar to those

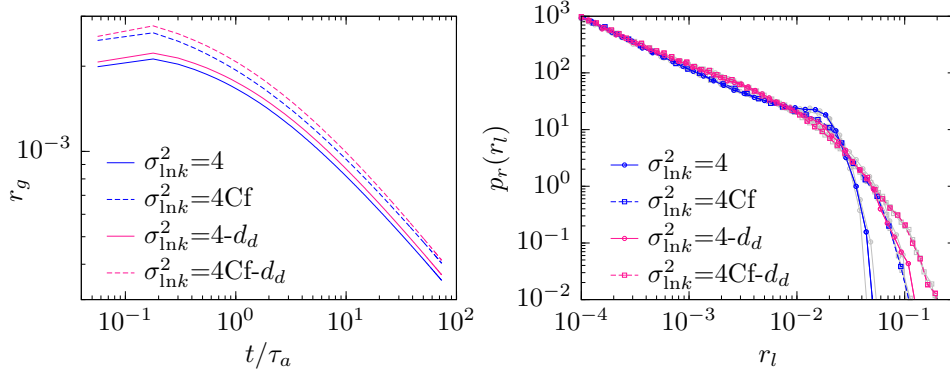
604 obtained for constant density, the data reveal the impact of buoyancy forces on the disper-  
 605 sion and mixing dynamics. For the strongly heterogeneous multi-Gaussian fields ( $\sigma_{\ln k}^2=4$ )  
 606 the variance is dramatically reduced compared to the scenario with constant density, and  
 607 follows a similar trend as the one obtained for the equivalent homogeneous case. For the  
 608 connected fields, the variance is slightly reduced in the presence of density variations. The  
 609 decrease of the width of the mixing zone with density variations is due to the competition  
 610 of heterogeneity-induced velocity variations (viscous forces) and buoyancy forces. Local  
 611 density gradients promote downward rotations in the flow field, which in turn lead to a  
 612 decrease in the emergence of vertical elongations in the interface. As illustrated in Figure  
 613 11, the amplitude of the induced vertical interface elongations, particularly in the freshwa-  
 614 ter end of the interface, is reduced by buoyancy effects. As a consequence, the deforma-  
 615 tion of the front and the width of the mixing zone decrease, see Figure 10(a,b).



616 **Figure 11.** Spatial distribution of the variance of the gradient of the mixing ratio distribution,  $\sigma_z(x, t)$   
 617 equation (7), for the equivalent homogeneous case, red line, and all heterogeneous realizations, gray lines,  
 618 and ensemble results considering density-constant and density-dependent ( $d_d$ ) flow, blue and pink lines,  
 619 respectively, for the multi-Gaussian and connected fields at time  $t' = t/\tau_d = 74$ .

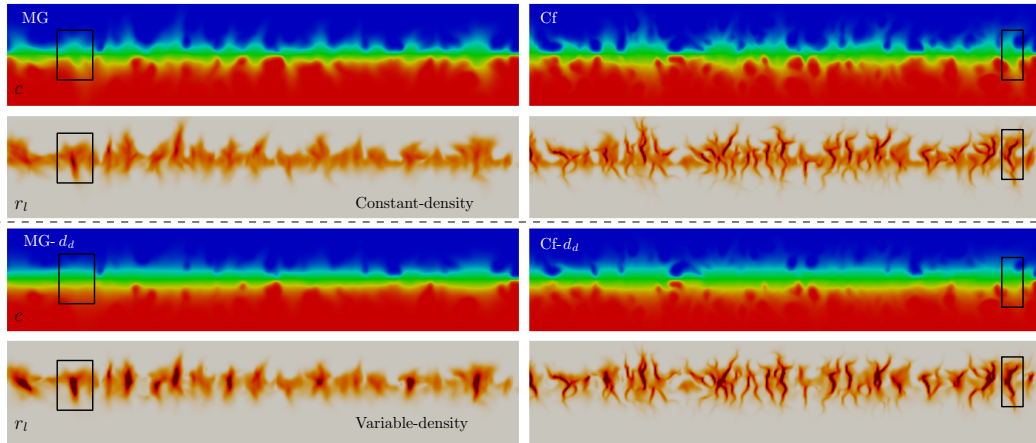
620 Thus one may conclude that the mixing rate also decreases due to buoyancy. How-  
 621 ever, the contrary is the case. The global mixing rate increases with density variations  
 622 as shown in Figure 10(c). This can be traced back to enhanced compression at the in-  
 623 terface. Buoyancy forces lead to a straightening out of the interface, which as discussed

624 above manifests in the decreases of its elongation and width  $\sigma_e(t)$ . This straightening out  
 625 leads to a compression of the interface, which in turn increases the gradients of the mix-  
 626 ing ratio in regions of high conductivity and thus high dispersion, see Figure 13. As a  
 627 result, the global mixing rate increases compared to the scenarios of constant density. This  
 628 at first counter-intuitive behavior stresses the importance of interfacial compression on the  
 629 mixing and reaction behavior.



630 **Figure 12.** Density-constant and -dependent ( $d_d$ ) results for the temporal behaviour of the global reactivity  
 631 (left) and reaction rate probability density function (right) for the multi-Gaussian (MG) and connected (Cf)  
 632 fields ( $\sigma_{\ln k}^2=4$ ) at at time  $t' = t/\tau_a = 74$ , which corresponds to the maximum of the head signal at the bound-  
 633 ary. The grey symbols denote corresponding PDFs in individual realizations, which differ only slightly from  
 634 the ensemble PDF.

635 In short, buoyancy effects cause concentration gradients to increase in the most per-  
 636 meable regions, which enhances the mixing rate. The increase of the mixing efficiency  
 637 leads directly to an increase of the reactivity of the system. Figure 12 illustrates the effect  
 638 of density variations on the evolution of the global reactivity and the PDF of the local re-  
 639 action rate. Note that, while the global reaction rate slightly increases with density varia-  
 640 tions compared to the case of constant density, the range of local reaction rates is strongly  
 641 enhanced by buoyancy effects in both the multi-Gaussian and connected fields. Mixing  
 642 and reaction are controlled by two primary mechanisms, (i) the combined effect of struc-  
 643 tural heterogeneity and transient forcing, which leads to enhanced strain in regions of high  
 644 conductivity and thus dispersion, and (ii) density effects which lead to compaction of the  
 645 interface, and thus an additional steepening of the concentration gradients.



646 **Figure 13.** Constant-density and variable-density results for the concentration and reaction rate distribution  
 647 for one multi-Gaussian field ( $\sigma_{\ln k}^2=4$ ).

648 Figure 13 compares the distributions of the mixing ratio and local reaction rates under  
 649 constant and variable density. The patterns are qualitatively similar. However, as pre-  
 650 viously discussed, density variations induce an enhancement of the concentration gradients  
 651 in high-velocity zones leading to a significant increase in the local reaction rates (black  
 652 squares). This results in a widening of the conduit section. It can be expected that this  
 653 effect is amplified when accounting for porosity and permeability changes. Moreover, al-  
 654 though the deformation of the mixing zone decreases with buoyancy effects and the shape  
 655 of the interface becomes more horizontal, complex reaction patterns emerge with a ver-  
 656 tical development of conduits which end sharply and are connected to a central horizon-  
 657 tal structure. This configuration is very common in coastal karst aquifers where there are  
 658 large main chambers, which can provide a measure of the glacio-eustatic sea level posi-  
 659 tions [Mylroie and Carew, 1988], with interconnected passages and vertical conduits [see,  
 660 e.g., Williams, 1977; Mylroie and Carew, 1995; Stoessell et al., 1989; Stoessell, 1995].  
 661 Therefore, the coupling of structural heterogeneity, transient forcing and density-driven  
 662 flow can be a very likely mechanism which would provide maze reaction network patterns  
 663 where the induced dispersive mass transfers mechanisms would be the dominant control  
 664 on the configuration of the conduits and cave formation.

#### 4 Summary and Conclusions

We study the interplay between temporal fluctuations, heterogeneity and density effects on mixing and chemical reactions between two fluids of different density under a stable stratification. In a two-dimensional stochastic modeling framework, dispersion and mixing are quantified by the interface width, the global mixing rate and the PDF of local mixing rates. We consider the mixing-induced dissolution of calcite. The reaction rate is determined from a mixing ratio-based formulation which decouples the solute transport and chemical speciation problem [De Simoni *et al.*, 2005].

We find that temporal fluctuations and heterogeneity slightly increase the width of the mixing zone between the two fluids, and that this effect is significantly enhanced by the degree of connectivity. Channeling patterns of high hydraulic conductivity promote significant velocity variations which leads to a strong deformation of the interface between the two fluids. This deformation, characterized by vertical interface elongations, evolves as a power law and stabilizes at asymptotic times. We also find that heterogeneity and transient forcing lead to a spatially nonuniform distribution of dispersion and concentration gradients. Hence, in high conductivity zones dispersive mass transfer is considerably enhanced. On the other hand, in low conductivity zones velocities and dispersion are very low which induces local maxima of the concentration gradients. This significant spatial variability leads to a wide distribution of the local mixing rates, whose maxima increase with the degree of heterogeneity and connectivity. However, this effect is not reflected in the behavior of the global mixing rate as a result of strong segregation, this means the regions of high mixing rate occupy only a small fraction of the domain. This suggests that the global mixing rate is not very informative of the mixing efficiency. Similar behaviors are observed for the global and local reaction rates. While the global reaction rate behaves in a similar way for all the heterogeneous fields, a strong spatial variability is observed for the local reaction rate distribution where the local potential reactivity is strongly increased by heterogeneity and connectivity. The variability in the local mixing and reaction rates is illustrated by their probability density functions, which are obtained through spatial sampling and contrasted to the corresponding PDFs for an equivalent homogeneous medium. We observe a clear increase in the range of mixing and reaction rates towards high values due to heterogeneity, particularly in the connected fields. Note that features such as connectivity depend on the spatial dimensions [Fiori and Jankovic, 2012]. Thus, while the

697 basic mechanisms are expected to be qualitatively the same in 3 dimensions, they may be  
698 quantitatively different.

699 Numerical results provide evidence that hotspots of chemical reactivity, this means  
700 zones of locally enhanced reaction rates, are clearly related to the deformation of the flow  
701 field induced by the coupling of spatial heterogeneity and transient forcing. Thus, hotspots  
702 are found to be localized in strongly stretched regions corresponding to high conductivity  
703 zones. Therefore, the evaluation of the strain tensor and topology of the flow field allows  
704 to delineate reaction patterns and predict regions high reactivity.

705 When density variations are included, there is a competition between viscous and  
706 buoyancy forces. While viscous forces induced by heterogeneity and temporal fluctuations  
707 promote the deformation of the interface, density variations tend to stabilize the shape  
708 of the mixing zone and favor an horizontal interface. As a result, the width of the mix-  
709 ing zone decreases. This interface compression leads to an increase of the concentration  
710 gradients in high conductivity zones which in turn enhanced mixing and reaction rates.  
711 Therefore, the interplay between spatial heterogeneity, density-driven flow and transient  
712 forcing magnify geochemical processes and the dissolution efficiency. Moreover, our re-  
713 sults provide a potential explanation for the formation of complex geochemical patterns  
714 observed karst systems, with vertical conduit developments connected to main horizontal  
715 bedding caves and chambers. Further analysis considering larger amplitude sea-level fluc-  
716 tuations such as regressions and transgressions or tectonic up- or down-lift displacements  
717 is required to explain more complex network patterns observed, for example, in coastal or  
718 island karst aquifers. In addition to that, other factors, such as the increase in dimension-  
719 ality of the media as well as the coupling of changes in permeability due to the increase  
720 of porosity, may affect transport and flow patterns and then enhance mixing and reactivity,  
721 which should be examined in future investigations.

## 722 **Acknowledgments**

723 The data used in this paper can be obtained upon request from the corresponding au-  
724 thor. Maria Pool acknowledges the support of the Juan de la Cierva Incorporacion grant  
725 (MINECO, Spain). Juan J. Hidalgo is gratefully acknowledged for the valuable assistance  
726 in the use of his code. The support of the European Research Council (ERC) through the  
727 project MHetScale (617511) is gratefully acknowledged. The authors also acknowledge the  
728 support of the Ministry of Economy, Industry and Competitiveness through the projects

729 MEDISTRAES and MEDISTRAES II (CGL2013-48869-C2-1 and CGL2013-48869-C2-2-  
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