



Nitrous oxide in the North Atlantic Meridional Overturning Circulation

N₂O Framework & Objectives

- Nitrous oxide (N₂O) is a gaseous compound responsible for two key mechanisms within the Earth's climate: acts as a long-lived and powerful greenhouse and secondly, is currently the most important ozone depleting substance.
- The extent to which water mass mixing and ocean ventilation contribute to N₂O distribution at the scale of oceanic basins is poorly constrained. We used novel N₂O and chlorofluorocarbon (CFC) measurements along with multiparameter water mass analysis to evaluate the impact of water mass mixing and Atlantic meridional overturning circulation (AMOC) on N₂O distribution along the OVIDE section, extending from Portugal to Greenland.

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- Concluding remark from this study: The most important features of the North Atlantic are mixing and southward advection of the atmospheric N₂O imprint in the lower limb of the AMOC, filled by cold, recently ventilated waters. In contrast, the upper limb of the AMOC is dominated by the production of N₂O in situ in the Iceland and West European basins, along with northward transport of mode and central waters enriched with N₂O due to accumulated remineralization in tropical and subtropical latitudes.

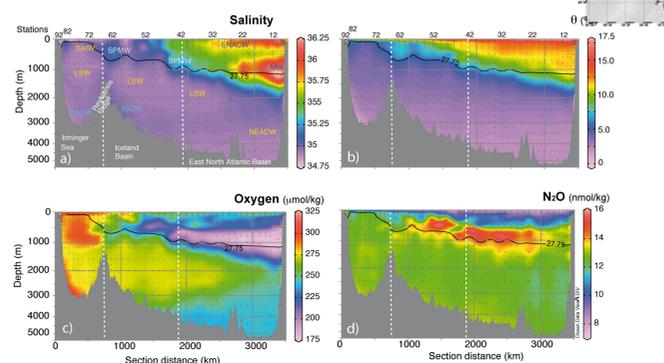
1 Atlantic Meridional Overturning Circulation: OVIDE experiment

The OVIDE section was conceived as an experiment for evaluation of the changes in the Atlantic meridional overturning circulation (AMOC) and in ocean ventilation.

The **upper limb of the AMOC** is driven by the North Atlantic current, which carries warm and saline waters from the subtropics toward the north-east Atlantic, being the main contributor the ENACW. The lower limb of the AMOC is driven by the water mass formation by deep convection in the Subpolar North Atlantic, being the LSW the most abundant in the OVIDE section.

The OVIDE has been occupied once every two years since 2002 as part of the World Ocean Circulation Experiment and CO₂/Climate Variability and Predictability Research Repeat Hydrographic programs. Since 2012, N₂O measurements were added to the set of properties studied in the OVIDE section.

3 N₂O distribution in the OVIDE section



- The distributions of θ and salinity are linked to the upper and lower limbs of the AMOC: warmer and saltier upper limb moving northwards, and fresher and cold waters moving southwards in the lower limbs
- The O₂ distribution shows a marked northwest-southeast gradient that is highly correlated with ventilation and overturning processes along the section. The combined effects of winter deep convection and the cold temperatures of the water masses when they are formed results in high O₂ values and in a homogeneous vertical O₂ pattern in the water column of the Irminger Sea.
- The N₂O lower values are observed in the eastern North Atlantic Basin in the warm ENACW, which increase northwards as the water cools, being the maximum N₂O surface concentrations in the Irminger Sea. The higher N₂O concentrations occur between 600 and 1200 m in depth in the eastern North Atlantic Basin, between the upper and lower limb of the AMOC.
- Below that maximum, the N₂O concentration remains relatively homogeneous down to the bottom and close to N₂O saturation, filling the lower limb of the AMOC.
- Why does the N₂O not match the O₂ distribution, as usually observed in oxic oceanic environments? What is the origin of the N₂O concentrations observed in the North Atlantic?**

4 Estimating the production of N₂O in ventilated regions

The AOU is widely used to infer respiration in the ocean in the basis of the assumption that the surface oxygen concentration is close to saturation with the overlying atmosphere. Nevertheless, significant oxygen disequilibrium has been observed in the high-latitude surface oceans where deep waters are formed, the case of the North Atlantic subpolar gyre.

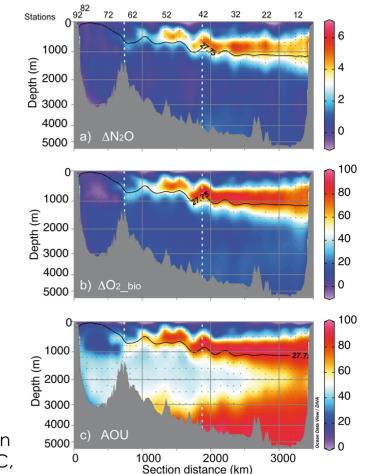
ΔN_2O known as apparent N₂O production, used as a proxy for the biotic N₂O produced since the water mass was last in contact with the atmosphere

$$\Delta N_2O = [N_2O]_{obs} - [N_2O]_{Eq}^{TTP}$$

AOU, proxy of the aerobic remineralization, do not include the disequilibrium with the atmosphere when water masses are formed.

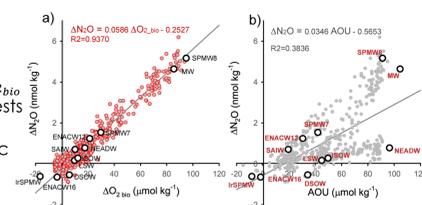
In this study, we computed the ΔO_2_{bio} that takes into consideration the preformed properties of water masses and is equivalent to true oxygen utilization (Ito *et al.*, 2004), providing a more realistic tracer of biological oxygen consumption.

The distribution of ΔN_2O and ΔO_2_{bio} in the OVIDE section, point to higher impact of biological production on the observed N₂O concentrations in the upper limb than in the lower limb of the AMOC, in which N₂O concentrations are driven mainly by temperature-solubility.



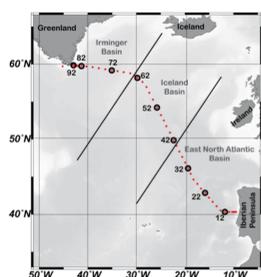
Exploring the stoichiometric N₂O:O₂ ratios

These results indicate that, when the oxygen saturation in preformed properties is considered, the stoichiometric ratio $\Delta N_2O/\Delta O_2_{bio}$ applies to the entire water column. This suggests that the double distributions in the correlation between ΔN_2O and AOU for the North Atlantic are probably due to biased use of AOU for ventilated waters and not to nitrification processes.



2 Methods

Repeated Hydrographic section: Cruises



Parameters
Salinity & Temperature
Dissolved Oxygen
Nitrate+ phosphate+ silicate
CFCs_{11,12}
+ N₂O

Years
2012
2014
2016

Extended Optimum Multiparameter Water Mass Analysis:

Compute the source water type (SWT) proportions that contribute to each water sample for each cruise (X_j) and Oxygen anomaly corrected by disequilibrium of water masses when formed: ΔO_2_{bio}

- Year 2012: García-Ibañez *et al.*, 2015
- Year 2014: García-Ibañez *et al.*, 2017
- Year 2016: In progress

Water mass averaged concentration of N₂O and O₂:

$$A_j = \sum_{i=1}^n X_{i,j} A_i^{MT} + \xi_j \quad j = 1 \text{ to } 475 \text{ samples}$$

A_j is the measured concentration of N₂O or O₂ in sample j, A^{MT}_i is the "mixing type" property = averaged concentration of A in the SWT i, X_{i,j} is the water mass fraction

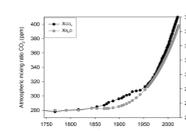
	θ (°C)	Salinity ^a	N ₂ O ^b (nmol kg ⁻¹)	O ₂ ^b (μmol kg ⁻¹)	ΔN ₂ O ^c (nmol kg ⁻¹)	ΔO ₂ ^{bio} (μmol kg ⁻¹)
ENACW ₁₆	16.10±0.13	36.20±0.02	7.88±0.34	244.73±5.20	-0.65	-3.48
ENACW ₁₂	12.30±0.18	35.66±0.03	10.39±0.10	229.13±1.61	1.23	21.31
MW	11.7±0.2	36.500±0.011	13.28±0.25	155.95±3.84	4.64	85.43
SAIW	6.0±0.2	34.70±0.03	12.84±0.25	278.57±3.91	0.71	10.17
SPMW ₁	8.00±0.11	35.23±0.016	15.52±0.21	195.04±3.18	5.17	94.82
SPMW ₂	7.07±0.07	35.16±0.006	12.64±0.21	250.84±3.27	1.54	29.82
ISPMW	5.0±0.02	35.03±0.013	12.60±0.29	316.53±4.53	-0.62	-16.46
LSW	3.00±0.19	34.87±0.02	12.64±0.10	277.14±1.53	0.17	9.75
ISOW	2.60±0.08	34.98±0.003	12.84±0.20	276.18±3.09	0.28	12.37
DSOW	1.30±0.06	34.90±0.006	13.28±0.15	303.65±10.38	-0.54	5.87
NEADW	1.98±0.03	34.89±0.003	12.33±0.14	234.39±2.11	0.78	17.68
SD			0.704	28.52	0.73	10.73

ENACW₁₆ and ENACW₁₂: Eastern North Atlantic Central Waters of 16 and 12 °C, respectively.
MW, Mediterranean Waters;
SAIW, Subarctic Intermediate Waters;
SPMW₁ and SPMW₂, Sub-polar Mode Waters of 7 and 8 °C, respectively
ISPMW, Irminger Sub-polar Mode Water;
LSW, Labrador Sea Water;
ISOW, Iceland-Scotland Overflow Water; DSOW, Denmark Strait Overflow Water; NEADW₁, lower North-East Atlantic Deep Water

Transient time distribution method (TTD):

Computation of mean age of water mass:

- N₂O equilibrium: Time-dependent atmospheric X_{N₂O} when water masses are formed

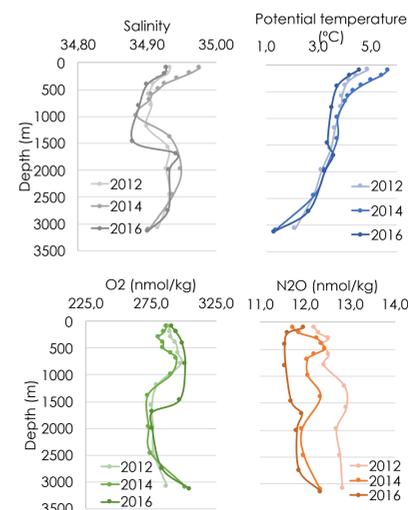


Increases in atmospheric N₂O must be considered in order to distinguish the fraction of N₂O derived from biotic processes from that of non-biotic origin related to equilibrium with atmospheric N₂O when water masses are formed

5 Work in progress

Exploring the temporal variation in the OVIDE:

2012, 2014, 2016



Data from central St in Irminger Sea

The N₂O samples were collected in the OVIDE section during the summers of 2012, 2014 and 2016. Coinciding with our study period, during 2016 recent investigations recorded the largest intensification of deep convection of the Labrador Sea Water since 1994. The imprint of this deep convection is appreciate in the vertical profile of the thermohaline properties, oxygen and N₂O. This enhance the uptake capacity of greenhouse gases in the Work in progress try to decipher the uptake capacity north Atlantic subpolar gyre of the anthropogenic N₂O into the deep AMOC.

Our study shows how including the N₂O measurements on hydrographic repeated sections allow to evaluate the oceanic response to large scale changes in biogeochemical processes and in oceanic circulation.

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