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**A modified hood infiltrometer to estimate the soil hydraulic properties from the  
transient water flow measurements**

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1 **Abstract**

2 In-situ measurements of soil hydraulic properties on covered soil surfaces (i.e. vegetated or  
3 residue covered surfaces) are of paramount importance in many agronomic or hydrological  
4 researches. These soil parameters are commonly estimated with the tension infiltrometry  
5 technique. This paper presents a portable and modified design of the hood infiltrometer (MHI)  
6 that, unlike to the original hood infiltrometer, allows estimating the soil hydraulic properties  
7 from the transient cumulative infiltration curve. The MHI consists of a water-supply reservoir  
8 attaches to a hat-shaped base placed on the soil surface. The base of the hat is closed by a  
9 system of sticks and a malleable material ring. To test the viability of this new design, the  
10 hydraulic conductivity ( $K_s$ ) estimated with MHI in a loam soil using the multiple head  
11 approach was compared to the corresponding values calculated from the transient infiltration  
12 curve analysis. Next, the MHI was tested on three different soils at saturated conditions, and  
13 the sorptivity ( $S$ ) and  $K_s$  estimated by the transient infiltration curve analysis were compared  
14 to the corresponding values obtained with a disc infiltrometer (DI). An additional field  
15 experiment was performed to compare the hydraulic properties measured with MHI on a bare  
16 soil and a soil covered with plants. Results demonstrated that this design allows hermetically  
17 closing the base of the hat without disturbing the soil surface. The  $K_s$  estimated with the  
18 multiple head approach was not statistically different ( $p = 0.61$ ) to that obtained with the  
19 transient infiltration curve analysis. No significant differences between the  $K_s$  ( $p = 0.66$ ) and  $S$   
20 ( $p = 0.50$ ) values estimated with DI and MHI were observed. The  $S$  values measured with  
21 MHI on the covered soil surface were significantly higher than that measured on the adjacent  
22 bare soil. These results indicate that MHI can be a viable alternative to estimate the hydraulic  
23 properties of covered soils from the measured transient infiltration curve.

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1 **Keywords:** Disc infiltrometer; Hydraulic conductivity; Sorptivity; Nonstationary water flow  
2 state.

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## 5 **1. Introduction**

6 Measurements of the sorptivity ( $S$ ) and hydraulic conductivity ( $K$ ) on undisturbed soil  
7 surface is crucial to solve many hydrological engineering and environmental issues linked to  
8 soil water storage and transport in the vadose zone. The tension disc infiltrometer (Perroux  
9 and White, 1988) and later the hood infiltrometer (Schwärzel and Punzel, 2007) have become  
10 popular infiltration methods because of the relatively rapid and portable nature of this  
11 technique, its easy in-situ applicability, and its ability to measure at unsaturated soil  
12 conditions (Angulo-Jaramillo et al., 2000). The disc infiltrometer consists of a disc base  
13 attached to a graduated water-supply reservoir and a bubbling tower to impose a negative  
14 pressure head ( $h$ ) at the disc base (Perroux and White, 1988). The hood infiltrometer, which  
15 can measure on covered soil surface, is a similar instrument where the disc base has been  
16 replaced by a side down hood.

17 The soil hydraulic properties are commonly calculated from the cumulative water-  
18 infiltration curves, which are measured from the drop in the water level of the reservoir tower.  
19 So far, several methods to estimate the soil hydraulic properties have been proposed: (i)  
20 methods based on the Wooding (1968) equation, which uses steady-state data (Ankeny et al.,  
21 1991; Reynolds and Elrick, 1991), (ii) methods based on the transient state data (Latorre et al.,  
22 2015) and (iii) methods which combine both transient and steady states, like BEST methods  
23 (Lassabatere et al., 2006; Yilmaz et al., 2010) The main advantages of the transient water  
24 flow procedure is that it requires shorter experiments, which involves smaller sampled soil  
25 volumes and more homogeneous initial water distribution (Angulo-Jaramillo et al., 2000).

1        The disc base of the infiltrometer, whose diameter ranges from 25 cm (Perroux and White,  
2        1988) to 3.2 cm (Madsen and Chandler, 2007), is usually covered with a tightened nylon cloth  
3        of very small mesh. The base should be completely in contact with the soil surface to  
4        accomplish correct infiltration measurements. To achieve this contact, Perroux and White  
5        (1988) recommended trimming any vegetation within the sample to ground level and cover  
6        the soil with a material (i.e. sand layer) of high hydraulic conductivity. However, according to  
7        Reynolds (2006), the contact sand layer introduces an offset between the pressure head set on  
8        the bubbling tower and the pressure head applied to the soil surface. This offset should be  
9        corrected to prevent the introduction of systematic biases in infiltration results (Reynolds,  
10       2006). These limitations were partially solved by Moret-Fernández et al. (2013), who  
11       developed an alternative disc base with a malleable membrane (MDB) that allowed excellent  
12       soil surface contact without using a contact sand layer. However, this membrane base  
13       prevents infiltration measurements on abrupt surfaces or soils covered with plants or crop  
14       residues. This problem was solved by UGT (Müncheberg, Germany), who developed a  
15       tension infiltrometer where the disc base was replaced by an acrylic (12.4 cm in diameter)  
16       hood. The hood is placed open side down onto the soil, within a retaining ring inserted into  
17       the soil, and the water-filled hood is directly in contact to the soil surface. Schwärzel and  
18       Punzel (2007) compared this new design with the conventional disc infiltrometer that uses a  
19       contact sand layer and observed that the hydraulic conductivities measured with this  
20       alternative design were 10 times higher than that measured with the disc infiltrometer.  
21       Although this system allows infiltration measurements on covered soils, the slow filling of the  
22       hood during the first infiltration times prevents estimating the soil hydraulic properties by  
23       using the transient water flow method. In this case, the multiple head approach, which is more  
24       time consuming, should be used. On the other hand, the retaining ring used to close the hood,

1 which is slightly inserted into the soil, may create preferential infiltration channels that distort  
2 the infiltration measurements.

3 The objective of this paper is to present a modified hood infiltrometer, MHI, to estimate the  
4 soil sorptivity and hydraulic conductivity on covered soil surfaces. Unlike to the original hood  
5 infiltrometer, which employs the multiple head approach, this new prototype allows recording  
6 and using the transient infiltration curve to estimate the soil hydraulic properties. Because of  
7 the multiple head approach has been satisfactorily applied on the hood infiltrometer, this work  
8 will be mainly focused on the transient infiltration flow method. To validate this new  
9 prototype, the soil hydraulic properties estimated with the MHI in three different uncovered  
10 soil surfaces were compared with those estimated with a conventional disc infiltrometer,  
11 which, similarly to MHI, also allows estimating the soil hydraulic properties from transient  
12 water flow measurements. In a second step, the MHI was tested on a covered soil. To this end,  
13 the soil hydraulic properties measured on covered surfaces were compared with those  
14 obtained in the adjacent uncovered soil..

15

## 16 **2. Material and methods**

### 17 *2.1. Infiltrometer design*

18 Similar to the disc infiltrometer (DI), the modified hood infiltrometer (MHI) consists of a  
19 hat-shaped base attached at the top to a water-supply reservoir and a bubbling tower that  
20 imposes a negative pressure head ( $h$ ) at the hat base (Fig. 1a and b). The hat base is a  
21 cylindrical acrylic tube (10 cm internal diameter –i.d-; 10 cm height) attached at the base to a  
22 metallic flat ring (3 mm thickness and 10 and 15 cm internal and external diameter,  
23 respectively), and closed at top by an acrylic lid. Three 1.5-mm deep holes are equidistantly  
24 made on the metallic ring, at 2.5 cm from the external diameter. The water reservoir consists  
25 of a 5 cm i.d. and 55 cm long acrylic tube. A vertical acrylic tube, that vertically traverses the

1 hat, is connected to the water reservoir through a ball valve (Fig. 1). This vertical tube, which  
2 is placed at 1.5 cm from the soil surface, contains a 3 mm i.d. plastic pipe (*air inlet tube*) that  
3 is connected to a bubbling tower. A 8 mm i.d. silicone tube (*air flow tube*) connects the top of  
4 the water reservoir tube to the top of the hat (Fig. 1). This tube is closed by an air flow plastic  
5 stopcock. To check the pressure head on the soil surface a water manometer is inserted at the  
6 top of the hat. Finally, a  $\pm 0.5$  psi differential pressure transducer (PT) (Microswitch,  
7 Honeywell), connected to a datalogger (CR1000, Campbell Scientist Inc.), is installed at the  
8 bottom of the water-supply reservoir (Casey and Derby, 2002). Previous laboratory  
9 experiments demonstrated the accuracy of PT for water level measurements was  $\pm 0.27$  mm.  
10 The base of the hat is closed by compressing the MHI base against the soil surface. To this  
11 end, a three detachable sticks system is used (Fig. 1b).

12

## 13 2.2. *Infiltrometer setup*

14 Installation of modified hood infiltrometer needed the following steps. Firstly, a 10 cm  
15 diameter piece of cloth is placed on the soil surface to be measured. This prevents the soil  
16 surface disturbance during the hat water-filling. A malleable material (Plasticine®) ring (11-  
17 cm i.d and 1.5 cm thickness) is pasted under the hat, and the hat plus the malleable material  
18 ring are placed on the soil surface, making sure the 10-cm diameter cloth rested within the hat.  
19 Three arm-sticks (30-cm length and 2-cm thickness) (Fig. 1b), which are welded to a  
20 perforated iron metallic head, are equidistantly and perpendicularly placed against the metallic  
21 ring perimeter. A 25 cm long screw is screwed at the end of the arm, resting the ends of the  
22 screws on the corresponding metallic ring holes. Three sticks (40-cm length and 2-cm thick)  
23 are introduced in the corresponding perforated iron heads (Fig. 1) and subsequently are driven  
24 into the soil down to 30-cm depth. The iron heads are blocked and the hat base plus the  
25 malleable material is compressed against the soil surface by screwing the arm-stick screws

1 against the hat base (Fig. 1). The strong pressure on the metallic ring, which squashes the  
2 malleable material against the soil surface, hermetically closes the base of the hat. In order to  
3 obtain sealing of the hat, the screws of the arm-sticks should be progressively and  
4 alternatively screwed.

5 Once the MHI base is installed, the bubbling tower is connected to the MHI *air inlet tube*  
6 (Fig. 1) and the water-supply reservoir is assembled on the hat. Next, the *air flow tube* is  
7 connected and the corresponding stopcock opened. The ball valve for water flow is turned off  
8 and the water reservoir is filled with water. Finally, the pressure transducer is connected to the  
9 data logger. Saturated infiltration measurements require that the pressure head inside the  
10 bubbling tower is equal to the distance between the soil surface and the end of the *air outlet*  
11 *tube* (Fig. 1). Thus, the pressure head measured with the water manometer ( $h_M$ ) is

$$12 \quad h_M = h_{BT} + h_{WL} \quad (1)$$

13 where  $h_{BT}$  is the pressure head supplied by bubble tower and  $h_{WL}$  the water level inside the hat  
14 (Fig. 1). To start the infiltration measurements, the ball valve for water flow is turned on and  
15 the plastic stopcock for air flow is kept opened until the water level inside the hat reaches 2 to  
16 4 cm height. This mechanism allows the air flows from the hat to the water reservoir, as the  
17 hat is filled with water. Once the plastic stopcock is closed, the air for water infiltration is  
18 immediately supplied from the bubble tower.

19

### 20 *2.3. Field testing*

21 A first field experiment to measure the maximum tension that can be maintained inside the  
22 hood infiltrometer before the air starts to enter by the base of the hat was performed. To this  
23 end, two infiltration experiments at saturation conditions were performed in a compacted soil  
24 located in an olive tree field in the campus of the Estación Experimental de Aula Dei (EEAD-  
25 Oli). The soil is loam and selected physical and chemical properties are summarized in Table

1 1. To monitor the pressure head changes in the hat, the water manometer was replaced by a  
2  $\pm 0.5$  psi PT connected to the data logger. Ten minutes after the start of the infiltration, the  
3 bubbling tower was blocked out and the infiltration continued until the pressure head in the  
4 hat was stabilized. This indicates that the vacuum into the hat was broken. According to Eq.  
5 (1), the maximum pressure head inside the hat ( $h_{\max}$ ) measured during the hat vacuum  
6 experiment was calculated according to

$$7 \quad h_{\max} = h_{M_{Final}} - (h_{BT_{Sat}} + h_{WL}) \quad (2)$$

8  
9 where  $h_{M_{Final}}$  is the final hat pressure head vacuum into the hat was broken (Fig. 2). Although  
10 it is not the main objective of this paper, an additional experiment was performed in the same  
11 field to test the viability of the MHI to estimate the hydraulic conductivity with the multiple  
12 head approach. Three infiltration measurements at three consecutive soil tensions, 0, 2 and 5  
13 cm, for 14, 20 and 25 minutes, respectively, were concluded. The pressure head inside in the  
14 hat was monitored with a  $\pm 0.5$  psi PT connected to the datalogger. Flow readings were  
15 automatically recorded every second from the drop in water level of the water supply  
16 reservoir. The  $K_s$  estimated from transient cumulative infiltration curve was next compared to  
17 the corresponding values estimated with the multiple head approach (Ankeny et al., 1991).

18 The soil hydraulic properties estimated with this new prototype were subsequently compared  
19 with those estimated with a disc infiltrometer (DI) in three experimental fields with different  
20 soil conditions (Table 1). The first field was the EEAD-Oli above described. The infiltration  
21 measurements were randomly distributed within a 25 m<sup>2</sup> surface. The second field (EEAD-  
22 NT) was located at the dryland research farm of the Estación Experimental de Aula Dei  
23 (CSIC) in the province of Zaragoza (41°44'N, 0°46'W, altitude 270 m). Soil at the research  
24 site is a loam (fine-loamy, mixed, thermic Xerollic Calciorthid) according to the USDA soil  
25 classification (Soil Survey Staff, 1975). Selected physical and chemical properties of the soil

1 (Table 1) were given in López et al. (1996) and Blanco-Moure et al. (2012). The infiltration  
2 measurements were conducted on a rectangular plot (30 x 10 m<sup>2</sup>) under no tillage treatment  
3 (NT), set up on a nearly level area (slope 0–2%). The experimental field corresponds to a long-  
4 term conservation tillage experiment started in 1991. The field was in the fallow period of 18  
5 months-long winter barley (*Hordeum vulgare* L.)–fallow rotation. The third field (CODO) was  
6 located in the Codo municipality (NE Spain; 41°30'N, 0°15'W). The land use in the area is  
7 based on a traditional agro-pastoral system involving dry cereal croplands and extensive sheep  
8 production. Soil at the research site is loam (Calcic Petrogypsis) according to the USDA  
9 classification (Soil Survey Staff, 2010). More details of chemical analysis of the soil (Table 1)  
10 were given Moret-Fernández et al. (2011). The infiltration measurements were randomly  
11 distributed within a 20 m<sup>2</sup> surface.

12 The DI employed in the experiment was a conventional Perroux and White (1988)  
13 infiltrometer with a base radius of 50 mm. To install the DI a circular thin layer of commercial  
14 sand (80–160 µm grain size), with the same diameter as the disc base, was layered on the soil  
15 surface. This allowed a good hydraulic contact between the base of the disc (covered with a 20  
16 µm mesh nylon cloth) and the soil surface. Similarly to the HI, the water level in the water  
17 reservoir was recorded with a ±0.5 psi PT, which connected to a datalogger (CR1000,  
18 Campbell Sci.) was installed at the bottom of the water supply reservoir.

19 All infiltration measurements were done on a nearly levelled and bare soil surfaces. The  
20 infiltration sites of DI were separated about 30-50 cm from MHI measurements. In all cases,  
21 only infiltration measurements at soil saturation conditions were conducted. Flow readings,  
22 which last up to 10 min, were automatically recorded every 5 s from the drop in water level of  
23 the water supply reservoir. The final soil water content, needed to calculate the hydraulic  
24 conductivity, was sampled from the upper centimetres of the soil just after removing the disc  
25 infiltrometer from the soil surface. The soil dry bulk density ( $\rho_b$ ), also used to determine the

1 initial volumetric water content of the soil, was determined by the core method with core  
2 dimensions of 50 mm diameter and 50 mm height. The core samples were taken near the  
3 measurement locations, the same day as infiltration measurements. Ten and four  $\rho_b$  samplings  
4 were taken in CODO and EEAD (EEAD-NT and EEAD-Oli), respectively. A total of 39 soil  
5 infiltration measurements for each infiltrometer type, 20 in CODO, 10 in EEAD-NT and 9 in  
6 EEAD-Oli, were completed. The  $K_s$  and  $S$  values were calculated from the cumulative  
7 infiltration curve using the Latorre et al. (2015) procedure, which analyses the transient  
8 cumulative infiltration curve using the cuasi-analytical solution of the Richards equation for a  
9 disc water source (Haverkamp et al., 1994). This procedure automatically omits the jump in  
10 the cumulative infiltration curve produced by contact sand layer, if used, and directly estimates  
11 both  $K_s$  and  $S$ . The  $K_s$  and  $S$  values estimated with DI were compared to the corresponding  
12 values measured with MHI. To this end an ANOVA test was used.

13 In order to check the viability of MHI on covered soils, an additional field experiment was  
14 done in the CODO field (Table 1). The hydraulic properties measured with MHI on soil  
15 covered with a plant of *Salsola Kali* were compared with the corresponding measurements  
16 obtained on the bare soil. Nine replications were performed in both covered and bare soils, and  
17 the distance between the pair of infiltration points was around 30-40 cm.

18

### 19 **3. Results and discussion**

20 Field experiments demonstrated that the system used to fix the hat of the infiltrometer on the  
21 soil surface is portable and easy to install. This also was an efficient system to hermetically  
22 close the base of the hat without disturbing the soil surface and without preventing the lateral  
23 water flow by capillarity. The time needed to install the MHI was less than 6 minutes. On  
24 average, the time to fill the hat up to 3.5 cm height sheet of water and the time to start the  
25 bubbling in the bubbling tower was about 2-3 seconds and 6 seconds, respectively. The hat

1 vacuum experiment showed that the average maximum pressure head allowed inside the hat  
2 was -12.1 cm (Fig. 2). Infiltration experiments under unsaturated infiltration conditions  
3 demonstrated that MHI can infiltrate at negative pressure heads (Fig. 3). The non-significant  
4 differences between the  $K_s$  estimated with the MHI using the multiple head approach and that  
5 using the transient cumulative infiltration analysis demonstrated that MHI can satisfactorily  
6 run with both methods (Tabla 2). However, while the multiple head approach needed about 1  
7 hour to estimate  $K_s$ , less than 15 minutes were needed with the transient infiltration method.

8 The cumulative infiltration curve obtained with MHI showed a large jump in the first  
9 seconds of the infiltration measurements (Fig. 4). This corresponds with the filling of the hat  
10 once the ball valve is opened. Despite this irregular shape, the  $K_s$  and  $S$  were successfully  
11 estimated with the Latorre et al. (2015) numerical procedure which, similarly to DI with  
12 contact sand layer, allowed correcting the infiltration jump. This method also allowed  
13 estimating the time needed to start the bubbling (Fig. 4). Overall, the deep well observed in  
14 the  $K_s$  and  $S$  error distribution (Fig. 4), calculated by the numerical optimization of the  
15 Haverkamp et al. (1994) model (Latorre et al., 2015), indicates the infiltration curves recorded  
16 with MHI allows accurate estimates of the soil hydraulic properties.

17 No significant differences between the  $K_s$  and  $S$  calculated in the three fields with the DI  
18 and the corresponding values estimated with the MHI were observed (Table 3). The standard  
19 deviation and the dispersion of the  $K_s$  and  $S$  values, due to the soil surface hydraulic properties  
20 variability, was similar in the two infiltrometers (Fig. 5). The estimated hydraulic parameters  
21 were within the same order of magnitude than those obtained by Moret and Arrúe (2007) and  
22 Moret-Fernández et al. (2013) in the same field, or Lampurlanés and Cantero-Martínez (2006)  
23 in a similar semiarid dry-land region. These results indicate that the MHI can be an alternative  
24 instrument to estimate the soil hydraulic parameters from the transient infiltration curve.

1 Comparison between the hydraulic properties measured in the CODO field on bare and  
2 covered soil showed that  $S$  under *Salsola* ( $0.303 \text{ mm s}^{-0.5}$ ) was significantly higher ( $p = 0.015$ )  
3 than the measured in bare soil ( $0.184 \text{ mm s}^{-0.5}$ ). These differences can be attributed to the  
4 higher organic matter content accumulated on the soil surface, under the *Salsola* plant, which  
5 may increase the water absorption capabilities during the first infiltration stages. No  
6 significant differences in  $K_s$  were observed between the different soil surfaces, which values  
7 were  $0.064$  and  $0.068 \text{ mm s}^{-1}$  for the bare and covered soil, respectively.

#### 9 **4. Conclusions**

10 This paper presents a modified design of the tension hood infiltrometer (MHI), that allow  
11 estimating the sorptivity ( $S$ ) and hydraulic conductivity ( $K$ ) on covered soils using the  
12 transient cumulative infiltration curve. Field tests demonstrated that MHI can estimate the soil  
13 hydraulic conductivity using both the multiple head approach and the transient infiltration  
14 curve analysis. The new prototype was validated by comparing the soil hydraulic properties  
15 estimated with this technique on three different uncovered soil surfaces with those measured  
16 with a conventional disc infiltrometer. The results demonstrate that this technique allowed  
17 accurate estimates of both sorptivity and hydraulic conductivity. Finally, this work  
18 demonstrates that this prototype also allows satisfactory estimations of the soil hydraulic  
19 properties on covered soil surfaces. Compared to the hood infiltrometer (Schwärzel and  
20 Punzel, 2007), the MHI allows using the transient cumulative infiltration curve to estimate the  
21 soil hydraulic properties, which substantially reduces the length of the experiment. These  
22 results show that the MHI can be an alternative to the DI when infiltration measurements are  
23 required on covered soils. However, caution should be taken when using this instrument, since  
24 erratic results can be obtained if: (i) the hat of the infiltrometer is not hermetically closed  
25 against the soil surface, for which lateral bubbling will be observed in the hat; or (ii) a

1 bubbling is observed from the wetted soil surface inside the hat. In these cases, similarly to  
2 the original hood infiltrometer, the experiment should be stopped and repeated in other place.

3

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## Figures captions

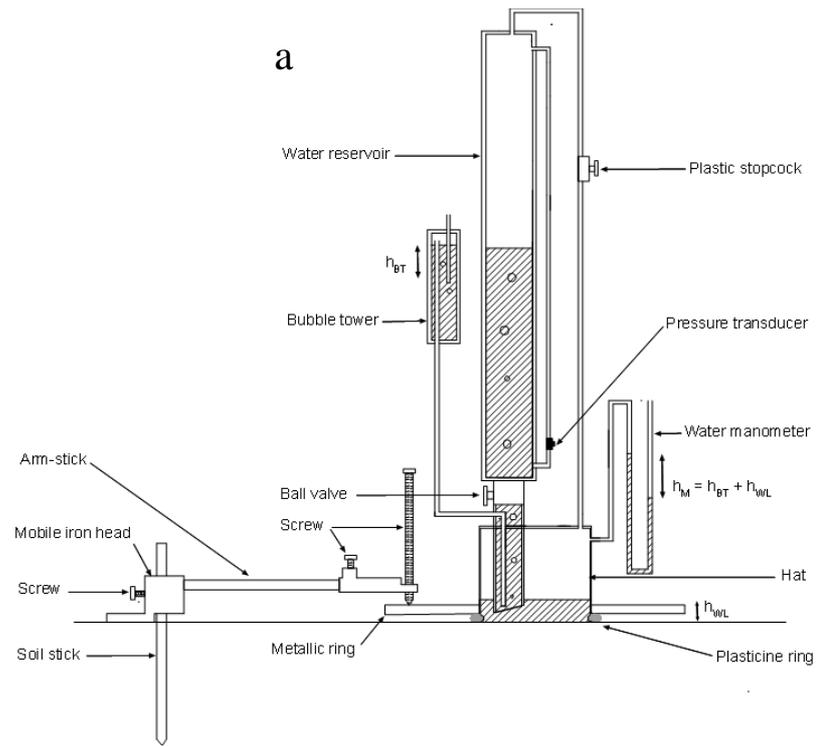
**Figure 1.** Diagram (a) and picture measuring in the NT field (b) of the modified hood infiltrometer.

**Figure 2.** Cumulative infiltration (I) and hat pressure head (h) curves measured with the modified hood infiltrometer during the hat vacuum capability experiment. White and grey colours indicate the first and second replication of infiltration measurements

**Figure 3.** First replication of the multiple head infiltration measurement (a) and the corresponding pressure head inside the infiltrometer hat (b) recorded with the MHI in the EEAD-Oli field

**Figure 4.** Measured (circles) and modelled (line) cumulative infiltration curves measured with the modified hood infiltrometer in the bare soil surface of the CODO field (a), and error distribution functions estimated for the sorptivity ,S, (b) and hydraulic conductivity ,  $K_s$  (c) according to the Latorre et al. (2015) procedure.

**Figure 5.** Soil sorptivity (S) and hydraulic conductivity (K) values measured with the disc and modified hood infiltrometers (MHI) in the CODO and EEAD (EEAD-NT and EEAD-Oli) experimental fields.



**Fig. 1.**

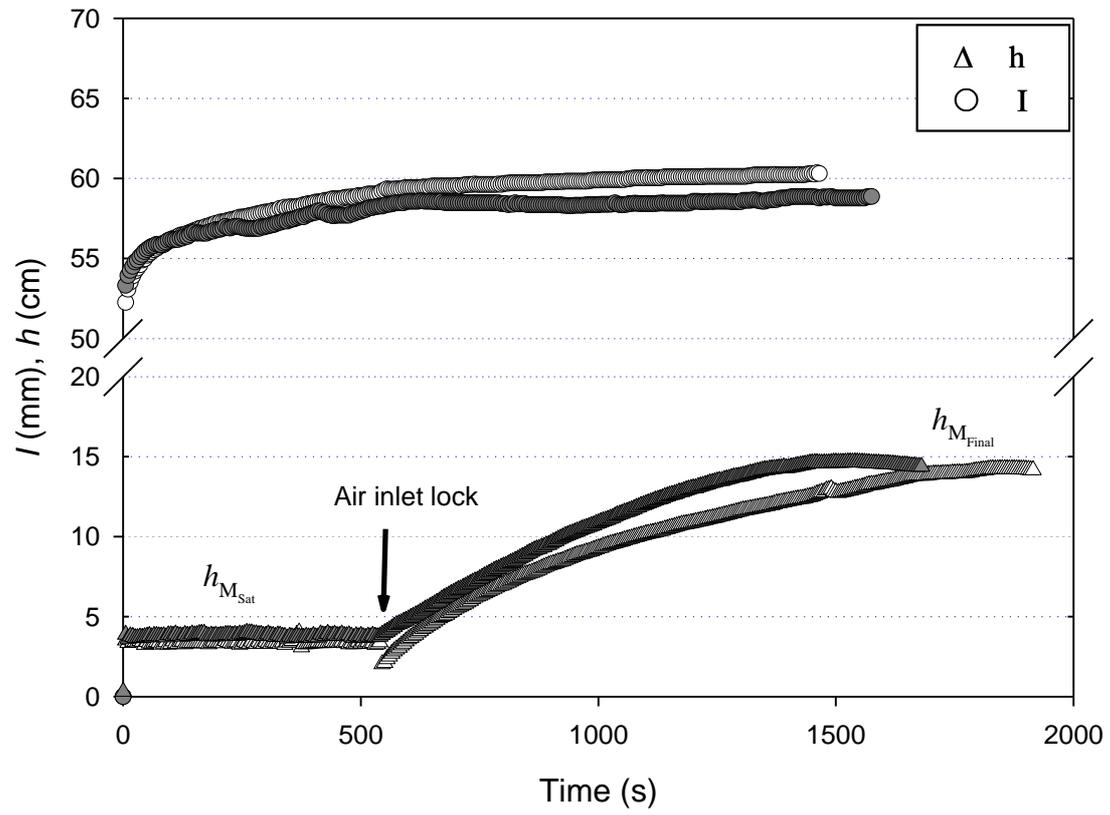


Fig. 2.

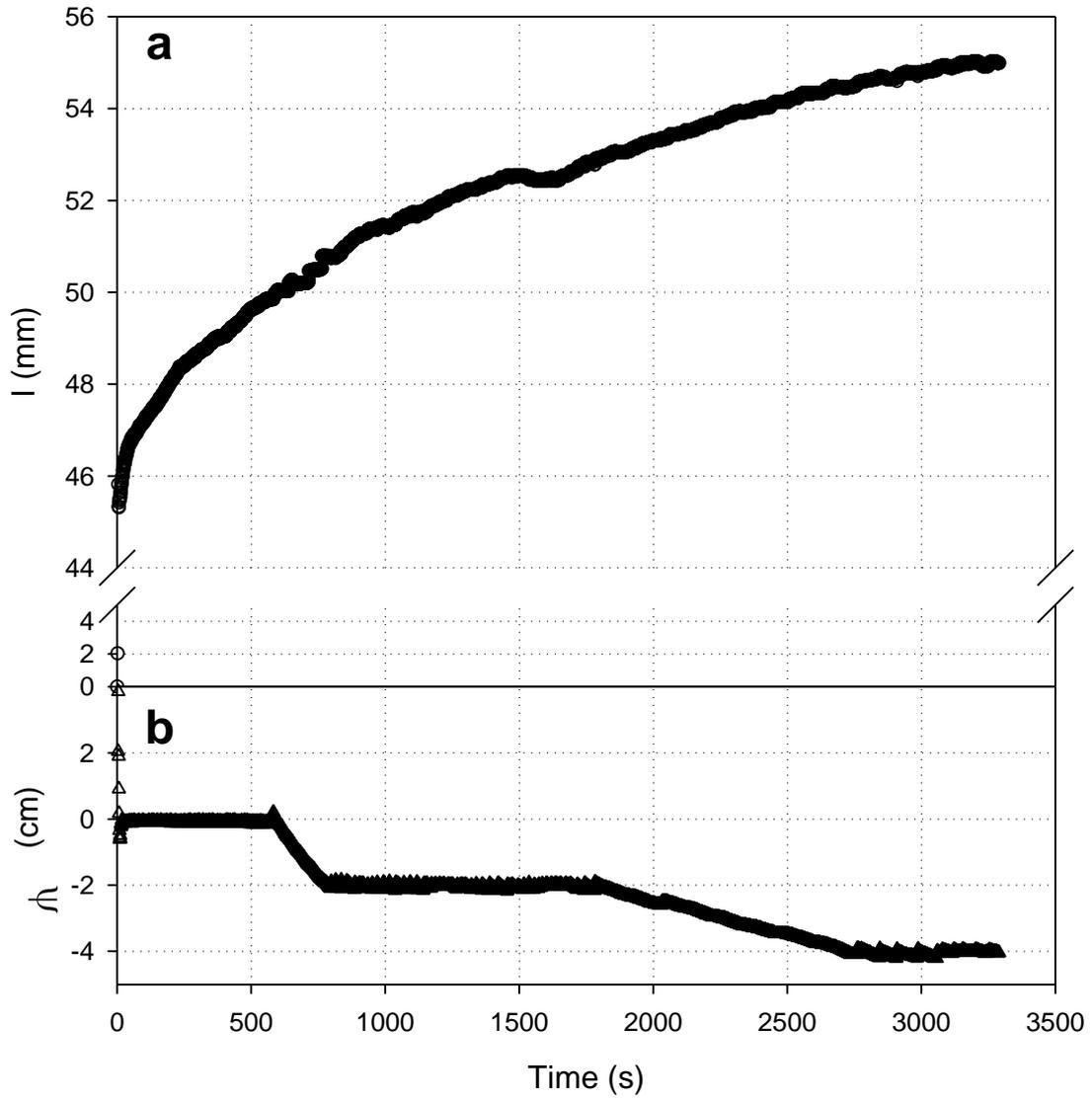


Fig. 3.

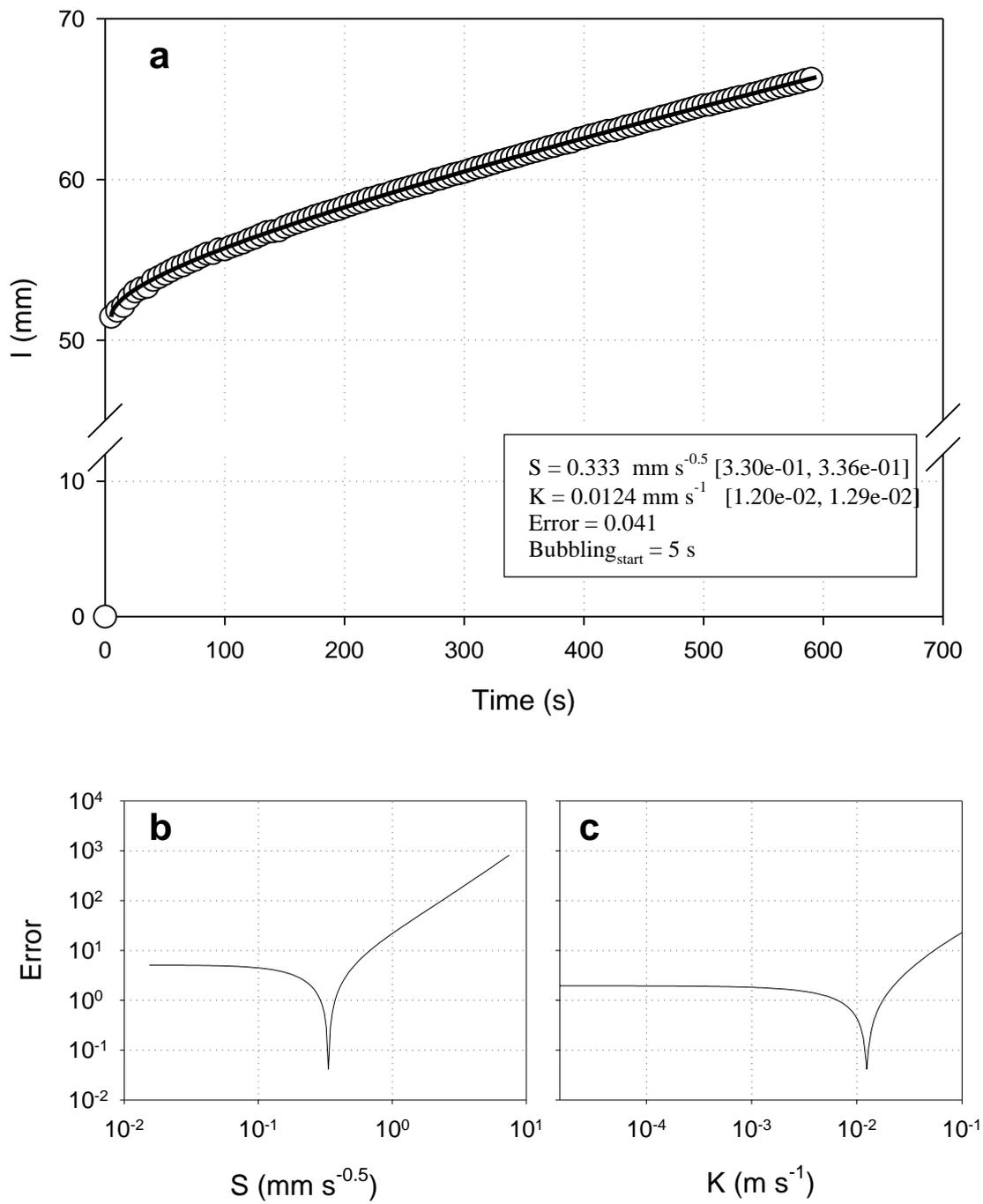


Fig. 4.

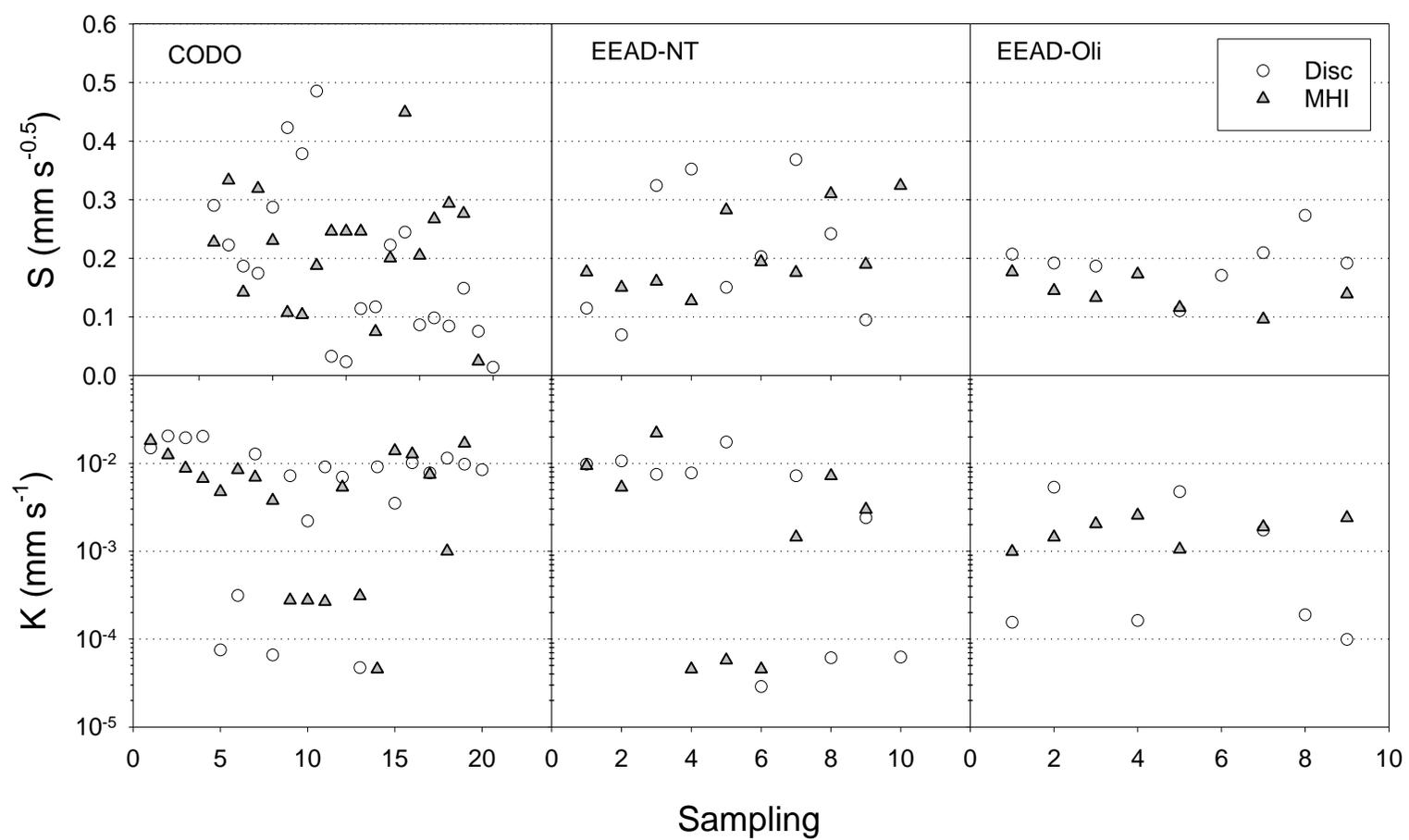


Figure 5.

**Table 1.** Altitude, average annual, precipitation (P), temperature (T), and average soil dry bulk density ( $\rho_b$ ) of the experimental plots located in the no-tillage (NT) and olive tree (Oli) farms of the Estación Experimental de Aula Dei (EEAD) and in the Codo municipality (CODO)

| Soil     | Altitude<br>(m) | P<br>(mm) | T<br>(°C) | $\rho_b$<br>(Mg m <sup>-3</sup> ) | pH<br>(H <sub>2</sub> O, 1:2.5) | EC (1:5) <sup>a</sup><br>dS m <sup>-1</sup> | Sand | Silt | Clay | g kg <sup>-1</sup> |        |                |
|----------|-----------------|-----------|-----------|-----------------------------------|---------------------------------|---|------|------|------|--------------------|--------|----------------|
|          |                 |           |           |                                   |                                 |   |      |      |      | CaCO <sub>3</sub>  | Gypsum | Organic carbon |
| EEAD-Oli | 260             | 390       | 14.5      | 1.61                              | 7.7                             | 0.24  | 366  | 404  | 230  | 353                | -      | 15.3           |
| EEAD-NT  | 270             | 390       | 14.5      | 1.52                              | 8.3                             | 0.31  | 313  | 451  | 236  | 473                | -      | 13.3           |
| CODO     | 400             | 313.7     | 14.5      | 1.34                              | 7.5                             | 1.96  | 422  | 409  | 169  | 923                | 402    | 2.5            |

<sup>a</sup>EC, electrical conductivity.

**Table 2.** Average and standard deviation (parenthesis) values for the K and S parameters estimated in the EEAD-Oli field with the modified hood infiltrometer using the transient flow (TF) and multiple head (MH) methods.  $K_s$ ,  $K_2$  and  $K_4$  are the soil hydraulic conductivity at saturation and -2 and -4 cm of pressure head, respectively.

|                  | $K_s$   | $K_2$<br>mm s <sup>-1</sup>                       | $K_4$   |
|------------------|---|---|---|
| TF               | 2.09 x 10 <sup>-3</sup> (9.0 x 10 <sup>-4</sup> ) | -   | -   |
| MH               | 1.85 x 10 <sup>-3</sup> (6.0 x 10 <sup>-4</sup> ) | 9.00 x 10 <sup>-4</sup> (4.0 x 10 <sup>-4</sup> ) | 5.00 x 10 <sup>-4</sup> (1.0 x 10 <sup>-4</sup> ) |
| Sig <sup>1</sup> | 0.68  | -   | -   |

<sup>1</sup> denotes the significance value for the ANOVA analysis

**Table 3.** Average and standard deviation (parenthesis) values for the  $K$  estimated with the disc and modified hood infiltrrometer in the CODO and EEAD experimental fields.

|                  | CODO        | EEAD-NT<br>S (mm s <sup>-0.5</sup> ) | EEAD-Oli    | CODO  | EEAD-NT<br>K (mm s <sup>-1</sup> )            | EEAD-Oli                                      |
|------------------|-------------|--------------------------------------|-------------|---|---|---|
| Disc             | 0.18 (0.13) | 0.21 (0.07)                          | 0.19 (0.06) | $6.2 \times 10^{-3}$ ( $5.0 \times 10^{-3}$ ) | $8.7 \times 10^{-3}$ ( $6.0 \times 10^{-3}$ ) | $1.7 \times 10^{-3}$ ( $2.0 \times 10^{-3}$ ) |
| Hood             | 0.22 (0.09) | 0.21 (0.11)                          | 0.14 (0.11) | $5.3 \times 10^{-3}$ ( $7.0 \times 10^{-3}$ ) | $6.7 \times 10^{-3}$ ( $6.0 \times 10^{-3}$ ) | $1.7 \times 10^{-3}$ ( $1.0 \times 10^{-3}$ ) |
| Sig <sup>1</sup> | 0.37        | 0.93                                 | 0.19        | 0.34  | 0.76  | 0.87  |

<sup>1</sup> denotes the significance value for the ANOVA analysis