

1 **Skillful seasonal predictions of boreal winter accumulated heating**  
2 **degree days and relevance for the weather derivative market**

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## ABSTRACT

5 It is demonstrated that boreal winter accumulated heating degree days, a weather derivative  
6 product frequently demanded by energy suppliers (among others), can be skillfully predicted  
7 with a lead time of one month, i.e. at the beginning of the previous November, for many  
8 regions of the Northern Hemisphere extratropcis. This finding contradicts the assumption  
9 of poor seasonal predictability for this variable. The present short note is meant to properly  
10 inform the participants of the weather derivative market in order to assure market trans-  
11 parency and to foster the scientific discussion on how to disseminate with this formerly  
12 unknown expert knowledge.

# 1. Introduction

The seasonal predictability of the boreal winter (DJF) climate in the Northern Hemisphere (NH) extratropics is generally assumed to be poor (Kushnir et al. 2006; Folland et al. 2012; Kim et al. 2012). However, it was recently demonstrated that DJF-mean 2m air temperatures over a large fraction of the NH can be predicted from Eurasian snow cover increase during the previous October, as described by the Snow Advance Index (SAI) (Cohen and Jones 2011; Cohen et al. 2012). In a follow-up study, it was shown that this index can be equally used for predicting DJF-precipitation totals in Europe, obtaining significant forecast skill over the Iberian Peninsula and southern Norway (Brands et al. 2012).

The physical basis for these empirical findings has been provided by a large number of observational and idealized general circulation model (GCM) studies (Cohen et al. 2001; Gong et al. 2003; Cohen et al. 2007; Hardiman et al. 2008; Fletcher et al. 2009; Kolstad and Charlton-Perez 2011; Smith et al. 2011). Following the conceptual model in Cohen et al. (2007), an above-normal Eurasian snow cover extend in October enhances the upward wave activity flux, which, in turn, weakens the stratospheric polar vortex. Due to the relatively long decorrelation time (or memory) of the stratosphere, this weakening tends to persist and propagate downward into the troposphere (Baldwin et al. 2003), thereby favoring the negative phase of the Arctic Oscillation during the following winter season. Although some aspects of this teleconnection pathway are not yet fully understood, it is increasingly recognized that the predictability of the boreal winter climate might be much larger than previously thought and that replicating the above mentioned lagged teleconnection in operative dynamical seasonal forecasting models could be key for improving their forecast skill (Cohen and Jones 2011; Peings et al. 2012).

The present study assesses the relevance of this formerly unknown expert knowledge for the weather derivative market and discusses some implications. This is done for the example of accumulated heating degree days (HDD), a common weather derivative product (Zeng 2000) which, for example, is traded at the Chicago Mercantile Exchange (CME) GLOBEX

40 electronic platform (<http://www.cmegroup.com/trading/weather/>). Before the start of  
41 the winter season, energy suppliers (among others) usually make contracts based on HDD  
42 in order to insure against a possible loss of sales that would be caused by an anomalously  
43 warm winter (Quayle and Diaz 1980; Sailor and Munoz 1997; Timmer and Lamb 2007). In  
44 this study, it will be shown that winter accumulated heating degree days (hereafter ‘DJF-  
45 HDD’) can be skillfully forecasted in many regions of the Northern Hemisphere extratropics,  
46 among them the southern United States, by using a statistical forecasting method based  
47 on simple linear regression. Since the standard deviation with respect to the climatological  
48 mean is usually taken as reference for defining the ‘strike’, i.e. the pre-negotiated threshold  
49 value above/below which the energy supplier/buyer will be paid out or not (Zeng 2000;  
50 Jewson and Brix 2005; Considine 2012), it constitutes the forecast variable in the present  
51 study. The added value/skill of the proposed forecasting method will be measured in terms  
52 of the percentage with which the purely climatological forecast (i.e. always predicting the  
53 climatological mean/standard deviation = 0) is outperformed. Having available a skillful  
54 prediction for the DJF-HDD on the 1st of November, i.e. one month ahead, is a kind of  
55 expert knowledge which should be made available in the field of weather derivative trading  
56 in order to assure market transparency.

## 57 2. Data and Methods

58 Accumulated heating degree days are commonly defined as follows:

$$HDD = \sum_{i=1}^n (T_{base} - T_{mean})^+ \quad (1)$$

59 where  $n$  is the number of days the HDD is calculated for, i.e.  $n=90$  for the DJF-season fo-  
60 cussed on in this study (note that the 29th of February is ignored in case of leap years),  $T_{base}$  is  
61 the baseline temperature which is set at the commonly used value of  $18^\circ C$  or  $65^\circ Fahrenheit$ ,  
62 valid for a typical uninsulated building, and  $T_{mean}$  is the daily-mean 2m air temperature at a

63 given location. The plus sign indicates that only positive values are accumulated (Buyukalaca  
64 et al. 2001).

65  $T_{mean}$  are calculated upon 6-hourly reanalysis data from ERA-Interim (Dee et al. 2011),  
66 which are publicly available at the European Center of Medium-Range Weather Forecasts  
67 ([http://data-portal.ecmwf.int/data/d/interim\\_daily/](http://data-portal.ecmwf.int/data/d/interim_daily/)). To confirm the results ob-  
68 tained with this gridded dataset for selected cities of the United States, in-situ station data  
69 from the United States National Climatic Data Center, which are freely available at the Av-  
70 erage Daily Temperature Archive of the University of Dayton ([http://academic.udayton.](http://academic.udayton.edu/kissock/http/Weather/default.htm)  
71 [edu/kissock/http/Weather/default.htm](http://academic.udayton.edu/kissock/http/Weather/default.htm)), are used in addition. Note that some station  
72 time series suffer from a very small number ( $< 0.5\%$  in any case) of large outlier values  
73 which are first set to missing values and then are linearly interpolated (Jewson and Brix  
74 2005). Thereafter, for each DJF-season between 1997/98 and 2011/12 ( $n=15$ ) and for each  
75 reanalysis grid-box and station location, the HDD values are calculated as described in Eq.  
76 1 and represent the values to be predicted (the predictands).

77 The increase of Eurasian snow cover in October is calculated using daily satellite re-  
78 trievals from NOAA (Ramsay 1998), which are stored at the National Snow and Ice Data  
79 Center (<ftp://sidads.colorado.edu/pub/DATASETS/NOAA/G02156/24km/>). For each Oc-  
80 tober between 1997 and 2012, the daily snow cover over a spatial domain covering  $25 - 60^{\circ}N$   
81 and  $0 - 180^{\circ}E$  is calculated and a time series of  $d = 31$  days of square kilometer values is  
82 obtained (Cohen and Jones 2011). Then, the robust linear regression approach described in  
83 Street et al. (1988) is applied to this time series. For this specific problem, robust regression  
84 is preferable to ordinary regression since it avoids the fitted straight line from being ‘moved’  
85 towards outlier values that do occur in the daily snow cover time series. Consequently, ro-  
86 bust regression is resistant to outlier values and essentially removes one uncertainty source.  
87 The slope of the robust regression line (i.e. the regression coefficient) is then taken as index  
88 value (Cohen and Jones 2011), which represents the rate of advance of the snow cover extend  
89 in October (and not the mean value). This index value is available on November 1st and is

90 applied as the only predictor from which DJF-HDDs are predicted.

91 As a statistical forecasting method, simple linear regression is applied in a one-year out  
92 cross-validation framework (Michaelsen 1987). The objective of this approach is to construct  
93 a loop in which the  $i^{th}$  of  $n = 15$  DJF-HDD values (expressed in standardized anomalies)  
94 is forecasted from the regression parameters obtained from regressing the remaining  $n - 1$   
95 HDD values against its corresponding predictor values. In order to avoid artificial skill  
96 (von Storch and Zwiers 1999), the  $i^{th}$  of the  $n = 15$  DJF-HDD values is forecasted from  
97 the standardized anomalies of the remaining  $n - 1$  HDD values, i.e. the calculation of the  
98 mean and standard deviation is repeated in each step of the cross validation. Note that  
99 the temporal autocorrelation of the DJF-HDD time series was found to be spurious in those  
100 regions where the forecasting method is successful (see below for a definition of ‘successful’).  
101 This is in qualitative agreement with Madden (1977) and Kushnir et al. (2006) and justifies  
102 the assumption of data independence the cross-validation approach relies on.

103 To measure the forecast/out-of-sample skill of the proposed forecasting scheme, the time  
104 series of 15 independent predictions is compared to its observed counterpart. The first  
105 measure of forecast skill applied here is the root mean square error skill score (*rmse<sub>ss</sub>*)  
106 defined as follows (Jolliffe and Stephenson 2003):

$$rmse_{ss} = \left(1 - \frac{rmse}{rmse_{ref}}\right) \times 100 \quad (2)$$

107 where *rmse* is the root mean square error of the time series predicted by the statistical  
108 forecasting method and *rmse<sub>ref</sub>* is the *rmse* obtained from always predicting the climato-  
109 logical mean, the latter being 0, since standardized anomalies/zero-mean data are calculated  
110 in each step of the cross-validation. Note at this point that calculating the climatological  
111 mean upon 15 or even 10 values only is a standard procedure in weather derivative trading  
112 (Considine 2012). Thus, *rmse<sub>ss</sub>* gives the percentage with which the error committed by  
113 assuming ‘normal’ winters is decreased by using the statistical forecasting method proposed  
114 here.

115 In addition to *rmse*, the Pearson correlation coefficient ( $r$ ) between the observed and  
116 forecasted DJF-HDD values is applied and its significance is computed using a two-sided  
117 t-test (null hypothesis  $H_0 : r = 0$ ), which is a standard procedure for assessing the skill of  
118 seasonal forecasts (Kim et al. 2012).

119 The above mentioned one-year-out cross-validation is applied separately for each grid-  
120 box and station. Due to the massive repetition of the method to thousands of grid boxes  
121 covering the NH, significant local out-of-sample skill ( $\alpha_{local} = 0.05$ ) over a areal fraction  
122 of 5% would be expected even if the method had no predictive capability (assuming zero  
123 spatial autocorrelation). Consequently, it must be assured that the areal fraction obtained  
124 by the forecasting approach is very unlikely to be a product of chance. To this end, and  
125 similar to the method applied in DelSole and Shukla (2009), the cross-validation approach  
126 is repeated 1000 times, using the randomly re-shuffled predictor time series in each step.  
127 The 99th percentile of the resulting 1000 areal fractions of locally significant forecast skill  
128 arising from chance is then assumed to represent the critical value above which global/field  
129 significance ( $\alpha_{global}$ ) is obtained. Note that the spatial autocorrelation of the predictand is  
130 taken into account by this technique.

131 To additionally demonstrate that the results are not sensitive to possible trends in the  
132 observed DJF-HDD time series, the statistical forecasting approach was also applied to lin-  
133 early de-trended time series (the detrending was repeated in each step of the cross-validation  
134 in order to avoid artificial skill). The corresponding results were found to closely match  
135 those obtained from undetrended time series. Since the method-related uncertainties of  
136 de-trending short-time series are manyfold (Jewson and Brix 2005), and since one-year-out  
137 cross-validation disrupts the temporal sequence (i.e. the independent variable of the trend),  
138 the results for the raw time series will be shown in the next section.

### 139 3. Results

140 Fig. 1a shows the forecast skill in terms of *rmse* for the DJF-HDD values obtained  
141 from gridded ERA-Interim 2m air temperatures (see Eq. 1). Areas exhibiting an *rmse*  
142  $> 20\%$  are found for the following regions (in alphabetic order):

- 143 • Central America
- 144 • Eastern Mediterranean
- 145 • Greenland
- 146 • North Africa
- 147 • Norway
- 148 • Siberia
- 149 • Southeast Asia
- 150 • Southern Alaska
- 151 • **Southern U.S.A**

152 the latter being most relevant for weather derivative trading. Note that the areal fraction  
153 of locally significant correlations ( $\alpha_{local} = 0.05$ ) between observed and forecasted values —  
154 which are not shown since they roughly correspond to *rmse* values  $> 20\%$ — was found to  
155 be highly significant ( $\alpha_{global} = 0.01$ ), i.e. field/global significance was obtained.

156 Since HDD values for the major U.S. cities are directly traded at the CME GLOBEX  
157 electronic platform, the statistical forecasting and cross-validation procedure is explicitly ap-  
158 plied to the respective station data. In other words, the weather derivative product ‘heating  
159 degree days for U.S. cities’ is directly forecasted in this step of the study. As visualized  
160 in Fig. 1b, the station-specific results are in close agreement with those obtained from re-  
161 analysis data. With a *rmse* of up to 47% and an  $r$  of up to 0.85 for the case of the



162 DJF-HDD in Houston (Texas), considerable forecast skill is yielded for the southern U.S.  
163 cities. Note that  $r$  is significant at a test level of at least 1% for each of the 6 major cities of  
164 the southern United States (exception: Jacksonville, where  $r$  is significant at a test level of  
165 5%). An overview of the validation results for the Southern cities, including  $rmse_{ss}$ ,  $r$  and  
166 the  $p$  - value is given in Tab. 1. As an example, observed and forecasted time series are  
167 shown in Fig. 1c for Houston (Texas).

## 168 4. Discussion

169 The statistical forecasting method proposed in this study is expected to be of interest for  
170 both the buyers and sellers of the weather derivative market, in order to decide on whether  
171 to close a contract or not. For instance, on the basis of the forecast for DJF 2011/12 (see  
172 Fig. 1c), which was available on the 1st of November 2011 and turned out to be reasonably  
173 correct, an energy supplier fearing financial loss in case the standard deviation for DJF-  
174 HDD is largely negative might have insured against this risk. In contrast, on the basis of the  
175 forecast for DJF 2009/10, which also turned out to be reasonably correct, he probably would  
176 not have done so. In turn, if the forecast for DJF 2011/12 would have been available for  
177 the insurer, he/she probably would have raised the premium/price for the above mentioned  
178 insurance.

179 In swap contracts, the ‘fair strike’ (i.e. the strike for which the expected pay-off is zero) is  
180 usually defined as the unconditional/climatological mean of the observed time series of DJF-  
181 HDD (Jewson and Brix 2005). However, in those regions where DJF-HDD are predictable  
182 (see Fig. 1a), this definition should be reconsidered since the expected value for the DJF-  
183 HDD following a large (small) Eurasian snow cover increase in October is larger (smaller)  
184 than the unconditional mean. As a possible solution for this problem, the ‘fair strike’ could  
185 be defined as a the mean value conditioned to the index for October Eurasian snow cover  
186 increase. However, the sample size of this index is small to-date ( $n = 15$ , due to the limited

187 availability of daily snow cover data), which, presently, would be one argument against the  
188 calculation of conditional means.

189 Finally, it should be noted that the inter-annual variability of the boreal winter climate is  
190 not solely driven by Eurasian snow cover (Kushnir et al. 2006; Fletcher and Saunders 2006;  
191 Folland et al. 2012). Moreover, due to the short sample size available to date, possible non-  
192 stationarities in the described teleconnection cannot be currently assessed. Consequently,  
193 the proposed statistical forecasting scheme should be re-evaluated in the future.

## 194 5. Conclusions

195 This study has shown that DJF accumulated heating degree days over a large and highly  
196 significant fraction of the Northern Hemisphere extratropics can be skillfully predicted from  
197 Eurasian snow cover increase in the previous October. For this purpose, simple linear re-  
198 gression has been applied as a statistical forecasting method which has been validated in a  
199 one-year-out cross-validation framework. With a root mean square error skill score of up to  
200 47% and a Pearson correlation coefficient of up 0.85, the forecast skill at 6 major cities of the  
201 southern United States is significant at a test level of up to 0.1%, and, thus, is very unlikely  
202 a product of chance. Since the latter variables are directly traded as weather derivative  
203 products, it is important to communicate this expert knowledge to the respective sellers and  
204 buyers (i.e. from the energy sector) in order to guarantee fair trading.

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## REFERENCES

- 216 Baldwin, M. P., D. B. Stephenson, D. W. J. Thompson, T. J. Dunkerton, A. J. Charlton,  
217 and A. O'Neill, 2003: Stratospheric memory and skill of extended-range weather forecasts.  
218 *Science*, **301 (5633)**, 636–640, doi:{10.1126/science.1087143}.
- 219 Brands, S., R. Manzanas, J. M. Gutierrez, and J. Cohen, 2012: Seasonal Predictability of  
220 Wintertime Precipitation in Europe Using the Snow Advance Index. *J. Climate*, **25 (12)**,  
221 4023–4028, doi:{10.1175/JCLI-D-12-00083.1}.
- 222 Buyukalaca, O., H. Bulut, and T. Yilmaz, 2001: Analysis of variable-base heating and cooling  
223 degree-days for Turkey. *Applied Energy*, **69 (4)**, 269–283, doi:{10.1016/S0306-2619(01)  
224 00017-4}.
- 225 Cohen, J., M. Barlow, P. J. Kushner, and K. Saito, 2007: Stratosphere-troposphere coupling  
226 and links with Eurasian land surface variability. *J. Climate*, **20 (21)**, 5335–5343, doi:  
227 {10.1175/2007JCLI1725.1}.
- 228 Cohen, J. and J. Jones, 2011: A new index for more accurate winter predictions. *Geophys.*  
229 *Res. Lett.*, **38**, doi:{10.1029/2011GL049626}.
- 230 Cohen, J., K. Saito, and D. Entekhabi, 2001: The role of the Siberian high in North-  
231 ern Hemisphere climate variability. *Geophys. Res. Lett.*, **28 (2)**, 299–302, doi:{10.1029/  
232 2000GL011927}.
- 233 Cohen, J. L., J. C. Furtado, M. A. Barlow, V. A. Alexeev, and J. E. Cherry, 2012: Arctic  
234 warming, increasing snow cover and widespread boreal winter cooling. *Environ. Res. Lett.*,  
235 **7 (1)**, doi:{10.1088/1748-9326/7/1/014007}.

- 236 Considine, G., 2012: Introduction to weather derivatives. Tech. rep., Weather Derivatives  
237 Group, Aquila Energy, Chicago.
- 238 Dee, D. P., et al., 2011: The ERA-Interim reanalysis: configuration and performance of  
239 the data assimilation system. *Quart. J. Roy. Meteor. Soc.*, **137 (656, Part a)**, 553–597,  
240 doi:{10.1002/qj.828}.
- 241 DelSole, T. and J. Shukla, 2009: Artificial Skill due to Predictor Screening. *J. Climate*,  
242 **22 (2)**, 331–345, doi:{10.1175/2008JCLI2414.1}.
- 243 Fletcher, C. G., S. C. Hardiman, P. J. Kushner, and J. Cohen, 2009: The Dynamical Re-  
244 sponse to Snow Cover Perturbations in a Large Ensemble of Atmospheric GCM Integra-  
245 tions. *J. Climate*, **22 (5)**, 1208–1222, doi:{10.1175/2008JCLI2505.1}.
- 246 Fletcher, C. G. and M. A. Saunders, 2006: Winter North Atlantic Oscillation hindcast skill:  
247 1900-2001. *J. Climate*, **19 (22)**, 5762–5776, doi:{10.1175/JCLI3949.1}.
- 248 Folland, C. K., A. A. Scaife, J. Lindesay, and D. B. Stephenson, 2012: How potentially  
249 predictable is northern European winter climate a season ahead? *Int. J. Climatol.*, **32 (6)**,  
250 801–818, doi:{10.1002/joc.2314}.
- 251 Gong, G., D. Entekhabi, and J. Cohen, 2003: Modeled Northern Hemisphere winter climate  
252 response to realistic Siberian snow anomalies. *J. Climate*, **16 (23)**, 3917–3931, doi:{10.  
253 1175/1520-0442(2003)016(3917:MNHWCR)2.0.CO;2}.
- 254 Hardiman, S. C., P. J. Kushner, and J. Cohen, 2008: Investigating the ability of general  
255 circulation models to capture the effects of Eurasian snow cover on winter climate. *J.*  
256 *Geophys. Res.*, **113 (D21)**, doi:{10.1029/2008JD010623}.
- 257 Jewson, S. and A. Brix, 2005: *Weather derivative valuation. The meteorological, statistical,*  
258 *financial and mathematical foundations.* Cambridge University Press, Cambridge.
- 259 Jolliffe, I. and D. Stephenson, 2003: *Forecast Verification.* Wiley, Chichester.

- 260 Kim, H.-M., P. J. Webster, and J. A. Curry, 2012: Seasonal prediction skill of ECMWF  
261 System 4 and NCEP CFSv2 retrospective forecast for the Northern Hemisphere Winter.  
262 *Climate Dyn.*, **39** (12), 2957–2973, doi:{10.1007/s00382-012-1364-6}.
- 263 Kolstad, E. W. and A. J. Charlton-Perez, 2011: Observed and simulated precursors of  
264 stratospheric polar vortex anomalies in the Northern Hemisphere. *Climate Dyn.*, **37** (7-  
265 8), 1443–1456, doi:{10.1007/s00382-010-0919-7}.
- 266 Kushnir, Y., W. A. Robinson, P. Chang, and A. W. Robertson, 2006: The physical basis for  
267 predicting Atlantic sector seasonal-to-interannual climate variability. *J. Climate*, **19** (23),  
268 5949–5970, doi:{10.1175/JCLI3943.1}.
- 269 Madden, R., 1977: Estimates of autocorrelations and spectra of seasonal mean tempera-  
270 tures over North-America. *Mon. Wea. Rev.*, **105** (1), 9–18, doi:{10.1175/1520-0493(1977)  
271 105<0009:EOTAAS>2.0.CO;2}.
- 272 Michaelsen, J., 1987: Cross-validation in statistical climate forecast models. *J. Climate Appl.*  
273 *Meteor.*, **26** (11), 1589–1600, doi:{10.1175/1520-0450(1987)026<1589:CVISCF>2.0.CO;2}.
- 274 Peings, Y., D. Saint-Martin, and H. Douville, 2012: A Numerical Sensitivity Study of the  
275 Influence of Siberian Snow on the Northern Annular Mode. *J. Climate*, **25** (2), 592–607,  
276 doi:{10.1175/JCLI-D-11-00038.1}.
- 277 Quayle, R. and H. Diaz, 1980: Heating degree-day data applied to residential heating energy-  
278 consumption. *J. Appl. Meteorol.*, **19** (3), 241–246, doi:{10.1175/1520-0450(1980)019<0241:  
279 HDDDAT>2.0.CO;2}.
- 280 Ramsay, B. H., 1998: The interactive multisensor snow and ice mapping system. *Hydrol. Pro-*  
281 *cesses*, **12** (10-11), 1537–1546, doi:{10.1002/(SICI)1099-1085(199808/09)12:10/11<1537::  
282 AID-HYP679>3.0.CO;2-A}.

- 283 Sailor, D. and J. Munoz, 1997: Sensitivity of electricity and natural gas consumption to  
284 climate in the USA - Methodology and results for eight states. *Energy*, **22** (10), 987–998,  
285 doi:{10.1016/S0360-5442(97)00034-0}.
- 286 Smith, K. L., P. J. Kushner, and J. Cohen, 2011: The Role of Linear Interference in North-  
287 ern Annular Mode Variability Associated with Eurasian Snow Cover Extent. *J. Climate*,  
288 **24** (23), 6185–6202, doi:{10.1175/JCLI-D-11-00055.1}.
- 289 Street, J., R. Carroll, and D. Ruppert, 1988: A note on computing robust regression  
290 estimates via iteratively reweighted least-squares. *Am. Stat.*, **42** (2), 152–154, doi:  
291 {10.2307/2684491}.
- 292 Timmer, R. P. and P. J. Lamb, 2007: Relations between temperature and residential natural  
293 gas consumption in the central and eastern United States. *J. Appl. Meteorol. Climatol.*,  
294 **46** (11), 1993–2013, doi:{10.1175/2007JAMC1552.1}.
- 295 von Storch, H. and F. Zwiers, 1999: *Statistical Analysis in Climate Research*. Cambridge  
296 University Press, Cambridge.
- 297 Zeng, L., 2000: Weather derivatives and weather insurance: Concept, application, and analy-  
298 sis. *Bull. Amer. Meteor. Soc.*, **81** (9), 2075–2082, doi:{10.1175/1520-0477(2000)081<2075:  
299 WDAWIC>2.3.CO;2}.

## 300 List of Tables

- 301 1 Forecast skill in terms of the root mean square error skill score (*rmsess*, in  
302 %), the Pearson correlation coefficient (*r*) and the p-value of the Pearson  
303 correlation coefficient (*pval*, in %). Also given are the Weather Bureau Army  
304 Navy (WBAN) station numbers. 16



TABLE 1. Forecast skill in terms of the root mean square error skill score (*rmsess*, in %), the Pearson correlation coefficient (*r*) and the p-value of the Pearson correlation coefficient (*pval*, in %). Also given are the Weather Bureau Army Navy (WBAN) station numbers.

<i>Name</i>	<i>Station number</i>	<i>rmsess</i>	<i>r</i>	<i>pval in %</i>
Atlanta	13874	36.9	0.78	0.07
Dallas/Fort Worth	03927	31.8	0.73	0.19
Houston	12960	47.0	0.85	0.01
Little Rock	13963	24.3	0.65	0.82
Jacksonville	13889	19.9	0.60	1.78
Raleigh/Durham	13722	28.4	0.70	0.36

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306 1 **Panel (a):** One-month lead forecast skill for DJF accumulated heating degree  
307 days, calculated upon 2m air temperatures from ERA-Interim, in terms of the  
308 root mean square error skill score, the latter calculated with respect to the  
309 climatological mean. Results are from a one year-out cross-validation for the  
310 DJF-seasons between 1997/98 and 2011/12 (n=15). **Panel (b):** as (a) but  
311 using in-situ station data. The forecast skill for the six named stations is  
312 significant in terms of the p-value of the Pearson correlation coefficient (see  
313 Tab. 1 for more details. **Panel (c):** Forecasted vs. observed DJF-HDD for  
314 Houston (Texas).

18

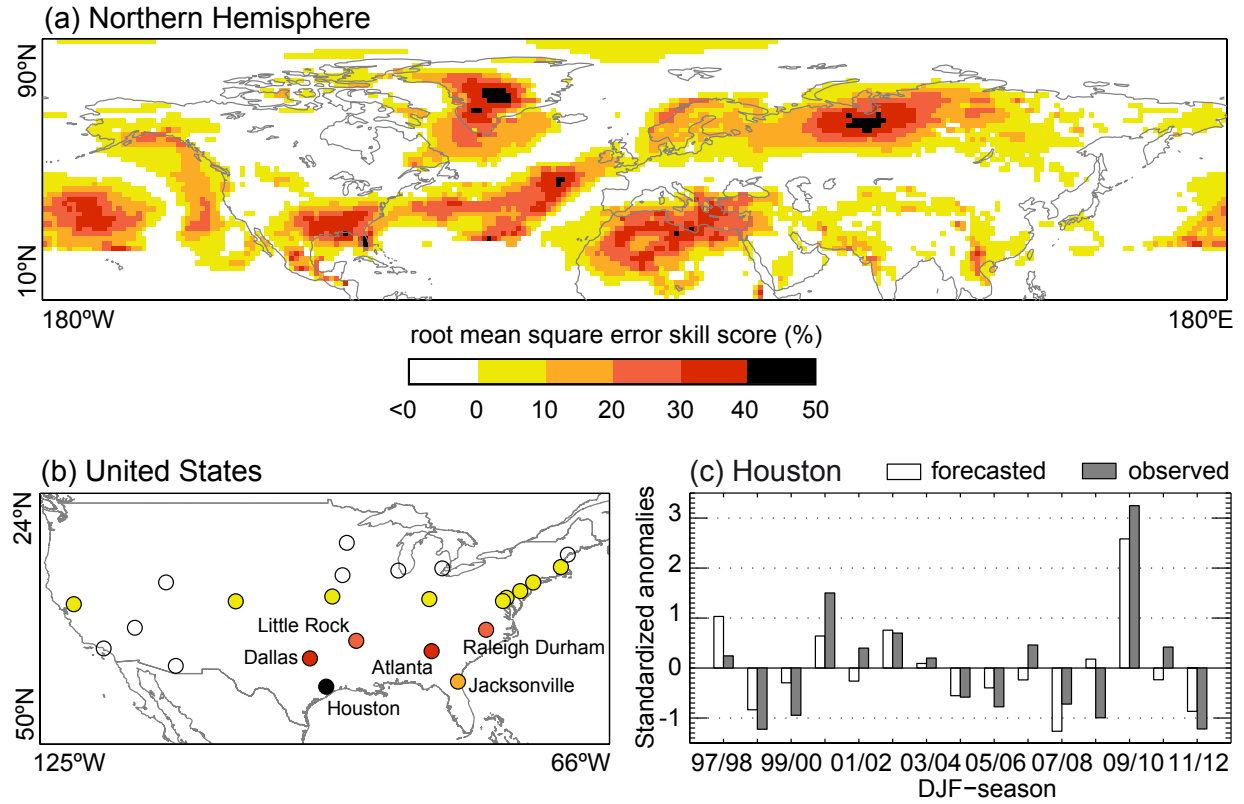


FIG. 1. **Panel (a)**: One-month lead forecast skill for DJF accumulated heating degree days, calculated upon 2m air temperatures from ERA-Interim, in terms of the root mean square error skill score, the latter calculated with respect to the climatological mean. Results are from a one year-out cross-validation for the DJF-seasons between 1997/98 and 2011/12 ( $n=15$ ). **Panel (b)**: as (a) but using in-situ station data. The forecast skill for the six named stations is significant in terms of the p-value of the Pearson correlation coefficient (see Tab. 1 for more details). **Panel (c)**: Forecasted vs. observed DJF-HDD for Houston (Texas).